

การศึกษาเชิงเปรียบเทียบของมรดกที่ได้จากสภาพการเกิดทางธรณีวิทยาต่าง ๆ



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สถาบันวิทยบริการ  
จุฬาลงกรณ์มหาวิทยาลัย

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COMPARATIVE STUDY OF EMERALDS FROM DIFFERENT  
GEOLOGICAL OCCURRENCES



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สถาบันวิทยบริการ  
จุฬาลงกรณ์มหาวิทยาลัย

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ฐิตินทรีย์ ปวโร : การศึกษาเชิงเปรียบเทียบของมรกตที่ได้จากสภาพการเกิดทาง ธรณีวิทยา  
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ประเทศไทยเป็นศูนย์กลางการเจริญในอัญมณี การผลิตเครื่องประดับและส่งออกสินค้าอัญมณีที่สำคัญ  
แห่งหนึ่งของโลก ประเทศไทยได้มีการนำเข้ามรกตก้อนหรือแบบที่เจียรระไนแล้วจากแหล่งต่าง ๆ ที่สำคัญ อาทิ  
โคลัมเบีย บราซิล แชนเบีย มาดากัสการ์ และไนจีเรีย การทราบแหล่งที่มาของมรกตจะมีผลต่อราคาและเป็น  
ประโยชน์ต่อวงการตลาดอัญมณี ดังนั้นจุดประสงค์ของการศึกษาค้นคว้าครั้งนี้ เพื่อเปรียบเทียบสมบัติต่าง ๆ ลักษณะ  
ภายใน การดูดกลืนรังสี และองค์ประกอบทางเคมีของมรกตที่ได้จากสภาพการเกิดทางธรณีวิทยาต่าง ๆ  
การวิจัยครั้งนี้ได้ทำการศึกษาดูตัวอย่างจากแหล่งมรกตที่เป็นที่นิยมในตลาดอัญมณีจำนวน 149  
ตัวอย่าง ซึ่งสามารถแบ่งเป็น 2 ประเภท ตามสภาพการเกิดทางธรณีวิทยาที่แตกต่างกัน คือ แบบที่ 1 มรกตที่เกิด  
ไม่สัมพันธ์กับหินซีสต์ (มรกตจากโคลัมเบีย และไนจีเรีย) และ แบบที่ 2 มรกตที่เกิดสัมพันธ์กับหินซีสต์ (มรกต  
จากบราซิล ได้แก่ แหล่ง Canaiba/Socoto, Itabira และ Santa Terezinha, มรกตจากมาดากัสการ์ และ  
แชนเบีย)

ลักษณะบางอย่างของมรกตสามารถบ่งชี้ถึงแหล่งกำเนิดได้ พบว่าค่าดัชนีหักเหและค่าความ  
ถ่วงจำเพาะสัมพันธ์กับธาตุอัลคาไลน์ โดยมรกตที่มีสภาพการเกิดแบบที่ 1 มีค่าดัชนีหักเหและค่าความ  
ถ่วงจำเพาะปานกลาง ถึงสูงนั้น จะมีธาตุอัลคาไลน์ที่สูงด้วย ในทางตรงกันข้าม มรกตที่มีสภาพการเกิดแบบที่ 2  
มีค่าดัชนีหักเหและค่าความถ่วงจำเพาะที่ต่ำ จะมีธาตุอัลคาไลน์ที่ต่ำ เช่นกัน มรกตที่มาจากแชนเบียและบราซิล  
(Canaiba/Socoto) พบว่ามีปริมาณธาตุโซเดียมและซีเซียมสูง ส่วนมรกตจากมาดากัสการ์ พบว่ามีปริมาณ  
โปแตสเซียมสูงมาก มรกตจากโคลัมเบีย พบว่ามีปริมาณเหล็กต่ำมาก ลักษณะมลทินภายในบางชนิดสามารถใช้  
จำแนกมรกตจากแหล่งต่าง ๆ ได้ เช่น ลักษณะรอบนอกคล้ายพื้นปลาของมลทินของไหลชนิดสามสถานะที่พบใน  
มรกตจากโคลัมเบียและไนจีเรีย พบว่ามรกตจากโคลัมเบียจะมีสัดส่วนของเหลว/ก๊าซ น้อยกว่าในมรกตจาก  
ไนจีเรีย เป็นต้น

สำหรับลักษณะทางสเปกตรัม ในส่วนของการดูดกลืนช่วงรังสี UV-Vis-NIR พบว่ามรกตส่วนใหญ่แสดง  
ช่วงการดูดกลืนของ โครเมียม/วานาเดียม และเหล็ก ยกเว้นมรกตจากโคลัมเบียที่ไม่แสดงการดูดกลืนของเหล็ก  
จากการศึกษาสเปกตรัมของน้ำในมรกตช่วง NIR สามารถหาความสัมพันธ์ของโมเลกุลน้ำในโครงสร้างของมรกต  
(ชนิด 1 และ 2) กับลักษณะการเกิดของมรกตในแต่ละแหล่งได้ คือ มรกตที่มีสภาพการเกิดแบบที่ 1 อย่าง  
ไนจีเรียจะแสดงสเปกตรัมของน้ำชนิดที่ 1 เด่นชัด ส่วนมรกตที่มีสภาพการเกิดแบบที่ 2 โดยเฉพาะมรกตจาก  
Santa Terezinha จะแสดงสเปกตรัมของน้ำชนิดที่ 2 เด่นชัด

ภาควิชา.....ธรณีวิทยา.....ลายมือชื่ออธิบดี.....  
สาขาวิชา.....โลกศาสตร์.....ลายมือชื่ออาจารย์ที่ปรึกษา.....  
ปีการศึกษา.....2547.....ลายมือชื่ออาจารย์ที่ปรึกษาร่วม.....

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KEY WORD: emerald / beryl / Brazil / Colombia / Madagascar / Zambia / Nigeria

THITINTHAREE PAVARO : COMPARATIVE STUDY OF EMERALDS FROM DIFFERENT GEOLOGICAL OCCURRENCES THESIS ADVISOR : ASSOC. PROF. VISUT PISUTHA-ARNOND, PH.D., THESIS CO-ADVISOR : WILAWAN ATICHAT, M.SC., 196 P., ISBN : 974-17-7069-3

Thailand is known as one of an important center for precious stones cutting, jewelry manufacturing and exporting. Rough or cutting emeralds have been imported to Thailand from foreign countries especially from Colombia, Brazil, Zambia, Madagascar and Nigeria. Determination of origin of emerald can be useful as value added factors in the gem marketing. This study aims at comparing the gemmological properties, internal characteristic, absorption spectra and chemical composition of emeralds from the different geological occurrences.

Totally 149 emerald samples from various types of important geological occurrence were acquired for this study. They can be divided into two different major categories: non-schist-related type I deposits (Nigeria and Colombia) and schist-related type II deposits (Carnaiba/Socoto, Itabira and Santa Terezinha in Brazil, Madagascar and Zambia).

Some characteristics are locality specific and can be used for origin determination. In addition, refractive index (RI) and specific gravity (SG) are related to alkali contents in emerald. The high to medium RI and SG values are found in emeralds from type II deposit which is high in alkali contents. In contrast the low RI and SG values are found in emeralds from type I deposit which is low in alkali contents. The emeralds from Zambia and Carnaiba/Socoto contain relatively high contents of Li and Cs, while the emeralds from Madagascar show exceptionally high K content. The Colombian emeralds show consistently very low iron content. Some inclusion features can be used to distinguish emeralds from different locality such as emeralds from Colombia and Nigeria contain different jagged outline in 3-phase inclusions in which those from Colombia have lower liquid/vapor ratios.

For the spectroscopic properties, UV-Vis-NIR absorption spectra of emeralds from most localities show a combination of  $\text{Cr}^{3+}/\text{V}^{3+}$  and  $\text{Fe}^{2+,3+}$ , except the spectrum of Colombian emerald has no  $\text{Fe}^{2+,3+}$  absorptions. NIR absorption spectra of water molecules (type I and II) in emerald's structure show relationship to alkali contents. In addition, emeralds from Nigeria of type I deposit is dominated by the type I water. In contrast, emeralds from Santa Terezinha in Brazil of type II deposit is dominated by the type II water.

Department.....Geology.....Student's signature.....

Field of Study..... Earth Science.....Advisor's signature.....

Academic year.....2004.....Co-advisor's signature.....

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# CHAPTER I

## INTRODUCTION

### 1.1 General Overview

Emerald is an intense green variety of beryl which is well-known as one of the most beautiful coloured gems. Other varieties of beryl are blue known as aquamarine, red as bixbite, yellow as heliodor and completely colourless as goshenite. Emerald is always a treasure for the appeal of attractive green colour calm, and the rich history of emerald combined with its beauty and rarity has made it the most valuable.

Emerald, as well as beryl, is composed of silicon, aluminum oxygen and beryllium which come from rocks of the continental crust. The elements that give the emerald its green colour are chromium and vanadium which concentrate below in upper mantle. The resulting folds and faults mobilized fluid that move along fractures. These fluids can dissolve and transport the necessary elements that have been brought together for emerald formation. As a result, In contrast to other varieties of beryl, emerald can only form in certain unique geological settings, and thus emerald generally does not occur in association with other varieties of beryl.

Emerald deposits are known from 5 continents namely Africa, South America, Asia, Europe and Australia. Emeralds can be found in several geological environments which are normally related to all potential source rocks of its component elements. The important production of emeralds are from South America, such as Colombia and Brazil, and from Africa, such as Zambia, Madagascar and Nigeria.

The first know emerald mine was the Cleopatra's mine in Egypt where it is generally accepted as the world's oldest source of emeralds. Egypt was the only significant source of emeralds until the 1500's, when the Spanish invaded the Americas (Ward, 1993). Up to that time, South America had been using emeralds in ornaments and ceremonies. The emeralds originated from Colombia and Brazil, were by far much larger, more transparent, and greener than those from Egypt. Since the 16<sup>th</sup> century, large quantities of Colombian emeralds have entered the European market. In very fine qualities, Colombian emeralds can retail for over \$20,000 a carat (Newman, 1995).

Thailand is known as one of an important center for precious stones cutting, jewelry manufacturing and exporting. The most important gemstones are ruby, sapphire and emerald. Rough emeralds have been imported from foreign countries especially from Colombia, Zambia and Brazil and cut in Thailand. Loose emeralds were then exported mainly to Europe, USA and Japan which according to the Thailand export statistics in 2003 the value was approximately 222.54 million baths or account to about 0.002% of gem and jewelry exporting revenue (<http://www.git.or.th>, Thailand Export Statistics).

Colombian emerald is by far the most desirable because of its splendor and clearer than those from the other deposits. Consequently, the emeralds from this source are considered the most valuable in the gem trades whereas most Brazilian emeralds contain large amount of inclusions, thus, are less desirable. It is therefore justifiable to study the well known deposits because emeralds in various countries of origins are different in their values.

## 1.2 Rationale and the Purpose of Study:

The country of origin of an emerald is one important aspect in gem identification. Many customers occasionally request a gem laboratory to issue the country of origin of an emerald in a gem identification report beside routine gem identification. Hence in order to be able to distinguish emeralds from different sources, it is necessary to collect and compare all available physical, optical and chemical data of emeralds from various sources. The assumption adopted in this study is that emeralds from different geologic occurrences should inherit some different gemological properties, internal characteristics as well as the chemical fingerprint that are able to give a clue to the country of origin of that emerald sample. Therefore the purpose of this research is 1) to study the gemmological properties, such as colour, refractive index, specific gravity, 2) to observe their internal characteristics, and 3) to compare the absorption spectra and trace element compositions of emeralds from different geological occurrences. The emerald samples used in this study were from Colombia, Brazil, Zambia, Madagascar and Nigeria that were obtained from a reliable person. The samples cover all major

types of emerald deposits and are among the most important emeralds commercially available in the gem markets today.

### 1.3 The method of study

Firstly, literature survey of the previous works was carried out on all technical aspects of emerald. Then some representative samples of emerald from various sources were acquired from a reliable person. The samples used in this study were obtained from 5 different countries of origin, i.e. Colombia, Brazil, Zambia, Madagascar and Nigeria. The physical and optical properties (i.e. refractive index (IR), birefringence, specific gravity (SG) and luminescence under short wave and long wave) were determined on polished slabs cut perpendicular to the c axis. After that the observation of internal characteristics, such as minerals and liquid inclusions, was undertaken together with photomicrographing. Measurement of the absorption spectra of the samples were then carried out using a UV-Vis-NIR Spectrophotometer and a Fourier Transform Infra-red Spectrophotometer (FTIR) at The Gem and Jewelry Institute of Thailand (GIT). Some selected samples from each locality were subsequently analyzed for their trace elements contents ( Li, Be, B, Na, Mg, Al, Si, P, K, Ca, Ti, V, Cr, Mn, Fe and Cs) using a Laser Abrasion Induced Coupled Plasma Mass Spectrometry (LA-ICP-MS) at the Macquarie University in Sydney, Australia. Results were then presented, discussed and concluded in the thesis report. The method of study can be summarized as follows (Figure 1.1)

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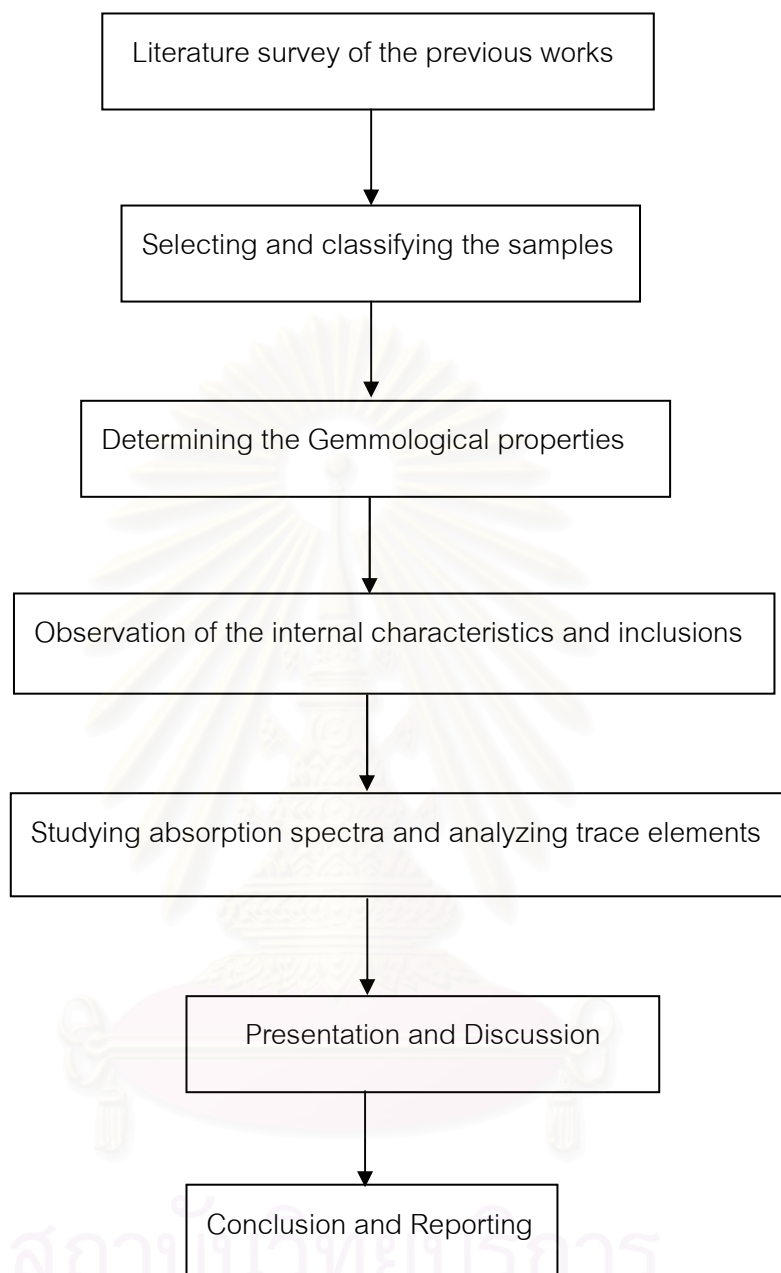


Figure 1.1: Flowchart of the method of study

## CHAPTER II

### GEOLOGY AND OCCURRENCES

An emerald is composed mainly of silicon, aluminum and oxygen. The forth primary component is beryllium, which is rare in the Earth's upper crust (~1.5 ppm). Therefore emerald is not a common mineral that comes from rocks of the continental crust. The elements give emerald its colour are chromium and vanadium that can occur in sedimentary rocks, particular in black shales. These elements are also found in basalts of the oceanic crust, in peridotites and dunites of Earth's upper mantle (Schwarz et al., 2002).

Specifically geologic and geochemical conditions are required for chromium and/or vanadium to encounter beryllium. There are a few deposits in which the circulation processes inside one geologic environment are sufficient for emerald formation, such as the black shales of Colombia (see later section). But in general the source rocks must be first open to permit the circulation and mobilization of elements. Then the action of tectonic plate, the result of folds and faults could mobilize fluids along fractures. As those fluids moved they could dissolve and transport the elements necessary for emerald formation. Emeralds can crystallize from a fluid in schists and gneisses or miarolitic cavities or quartz lenses, faults, breccias, lenticular vugs, druses and isolated crystal pocket in geologic structure such as veins along fracture zones (Schwarz et al., 2002).

#### 2.1 Classification of deposits of study emerald:

Emerald deposits are known from 5 continents namely Africa, South America, Asia, Europe and Australia. The important production of emerald is however from South America, such as Colombia and Brazil, and from Africa, such as Zambia, Madagascar and Nigeria.

Emeralds can be found in several geological environments which are normally related to all potential source rocks of its component elements. In general two main

types of emerald deposits can be distinguished in this study; type I deposits are non-schist-related and type II deposits are schist-related. Both types are mainly controlled by tectonic structures, such as faults and shear-zones (Schwarz, et al, 2002).

The type I deposits (non-schist-related emerald mineralization) can be subdivided into 2 sub-types:

- Ia; pegmatite without schist; e.g., from Kaduna Plateau of Nigeria (NI) and
- Ib; black shales with veins and breccias; e.g., from Cordillera Oriental of Colombia (CO).

The type II deposits (schist-related emerald mineralization) can be further subdivided into 2 sub-types:

- IIa; schist without pegmatite; e.g., from Santa Terezinha (STA) of Brazil and
- IIb; schist with pegmatite; e.g., from Carnaiba/Socoto (CAR) and Itabira, (ITA) of Brazil, from Mananjary of Madagascar (MA) and from Ndola Rural district of Zambia (ZA). A summary of all the emerald deposits are shown in Table 2.1

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Table 2.1: Summary of geological occurrences of emeralds from Brazil, Colombia, Madagascar, Zambia and Nigeria.

Type I Deposits (Non-schist-related)		Type II Deposits (Schist-related)				
Type Ia (Pegmatite without Schist)	Type Ib (Black Shales with Viens and Breccias)	Type IIa (Schist without Pegmatite)	Type IIb (Pegmatite with Schist); emerald occur between contact zone of country rock and pegmatite veins			
Emerald crystallized in vugs of granitic - rock, which intruded to the country rock.	Host rock is black shales.	Host rock is talc- chlorite- amphibole- magnetite schist.	Host rock is phlogopite- carbonate - talc schist.	Host rock is amphibolite.	Host rock is serpentinite.	Host rock is talc.
Kaduna Plateau, Nigeria	Cordillera Oriental, Colombia	Santa Terezinha (Brazil)	Mananjary, Madagascar	Ndola Rural, Zambia	Carnaiba/ Socoto (Brazil)	Itabira (Brazil)

\* The detail of each deposit is described below.



## 2.2 Type I: Non-schist-related emerald deposits;

### *2.2.1 Ia: Emeralds in granite and pegmatite vugs without schist seams, e.g. Kaduna and Plateau States of Nigeria*

The emerald mineralization of the Kaduna and Plateau States in Nigeria is associated with two periods of magmatism and the accompanying intrusion of alkali granites: the Pan-African orogenesis (600 - 450 million years ago) and the Mesozoic orogenesis (190-144 million years ago). These deposits also contain significant elements, such as tin, niobium, tantalum and zinc (Schwarz, 1996).

To the east of the Nigeria River and north of the Benue River, there is an important pegmatite belt which is apparently not related to major granite intrusions. The pegmatites occur in an area bounded to the north by Kafanchan, to the south by the Afu Hills, to the east by the Jos Plateau and to the west by Nassarawa (Figure 2.1). These pegmatites vary considerably in their mineralogy. The bulk of those which are metal-bearing are complex albitized pegmatites with important gem potential.

The emerald occurs in small pegmatitic pockets in association with quartz, feldspar and topaz (Figures 2. 2 and 2.3). These pegmatite pockets, which can be up to 8 cm in size, are found at the contact of a granite with basement rocks and represent cavities created by gas loss from the cooling magma. The temperatures of crystallization of Nigerian emeralds were reported to be about 400 - 500 °c (Schwarz, 1996).

In the Pan-African pegmatites, emerald is associated with aquamarine, tourmaline and beryl. In the Mesozoic pegmatites, the emerald occurs in pegmatite cavities or in phlogopite-alkali-feldspar-granites with quartz, blue topaz and beryl/aquamarine (Schwarz et al., 2002).

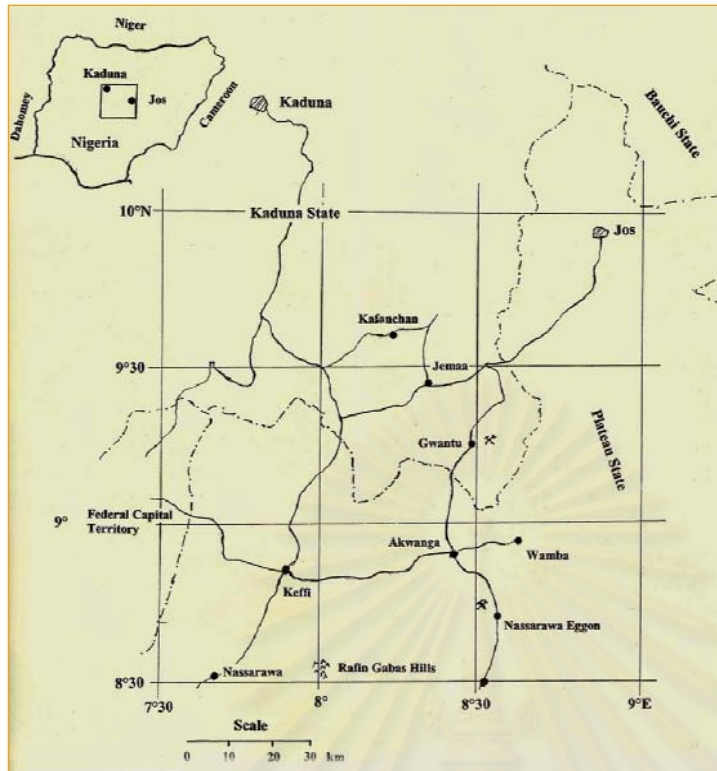


Figure 2.1: Location map with the emerald occurrences at the Kaduna and Plateau States in Central Nigeria (Schwarz, 1996).

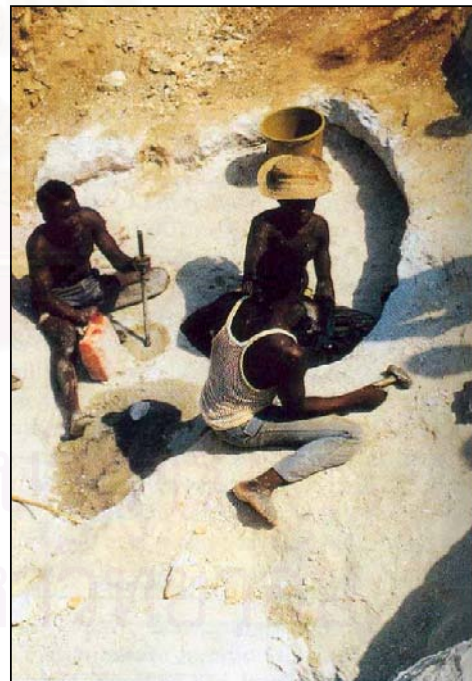
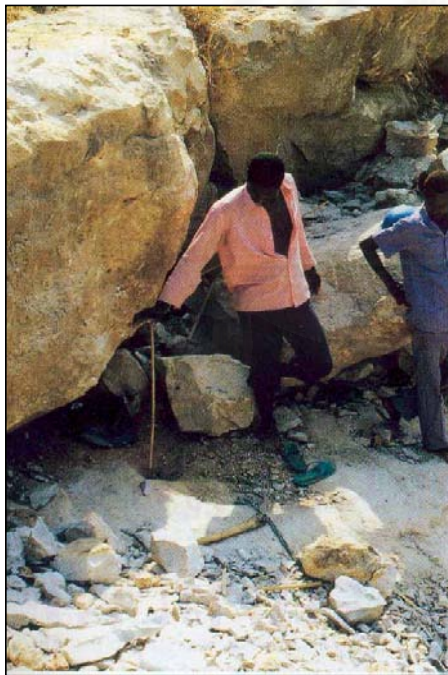


Figure 2.2: (left) and Figure 2.3: (right) Searching and drilled holes prepared for blasting to expose in the zone of granites (Schwarz, 1996).

According to Schwarz, et al. in 2002, the emeralds in Nigeria were probably formed between magmatic and hydrothermal stages. The alkali-granites were enriched with beryllium and fluorine. The beryllium might have been derived from the micas and feldspars of the granites and the chromium could have come from the mafic rocks. The magmatic origin of the mineralizing solution was indicated by the oxygen and hydrogen isotopic compositions of the emerald. Geologic structures and the type of mineralization were characterized by the interaction of solution with volcanic rocks.

### ***2.2.2 Ib: Emeralds in veins and breccias in black shales: e.g. Cordillera Oriental of Colombia***

Colombian emeralds are significant in both quality and quantity. The deposits were exploited by the native before the appearance of Europeans in the 16<sup>th</sup> century and have been worked more or less afterward. The occurrences of emerald are so widespread in this area that new deposits have been found in modern times and the probability is still great to be discovered in the future (Sinkankas, 1981). The new and famous deposits are Muzo and Chivor mines (Figure 2.4), where renowned worldwide for the quality and quantity of their emeralds (Bosshart, 1991). There is also a large number of smaller deposits. These emerald occurrences form two narrow zones extending near and along the two multiphase fault boundaries of the Eastern Cordillera. These are the same original boundaries of the Cretaceous sedimentary basins. The eastern zone contains the mining districts of Chivor, the western zone is the districts of Muzo and Coscuez.

The Colombian emeralds occur in calcite veins intensely folded and invaded running vertically through dark shale (Figure 2.5) of early Cretaceous age (about 120 to 130 million years old). Emerald forms as an accessory mineral in hydrothermal albite-carbonate-pyrite veins (Keller, 1981). In all mining district, the emerald-bearing veins are crosscutting the strata and are spatially associated with breccias and albites. The dark rock series was chalky and argillaceous of marine origin and is therefore grey to black in colour (Figures 2.6).



Geochemical studies have indicated the absence of magmatic activity related to the origin of the emeralds. The composition of water from the structural channels and oxygen in Colombian emerald is consistent with basinal brines that have interacted with local evaporate deposits. These brines were likely reacted with organic matter in the black shales, releasing organically trapped beryllium, chromium and vanadium into the solution (Schwarz, 1998).

Since the early Tertiary period, the black shale series have been lifted, folded, faulted and mineralized. Many emerald appear in places, where large scale tectonic faults once cut into the sediments (Bosshart, 1991). In the course of sodium metasomatic interaction which probably took place at relatively low temperature of about 300°C, hydrothermal fluids leached the beryllium, chromium and vanadium from the black shales (Schwarz and Giuliani, 2002).

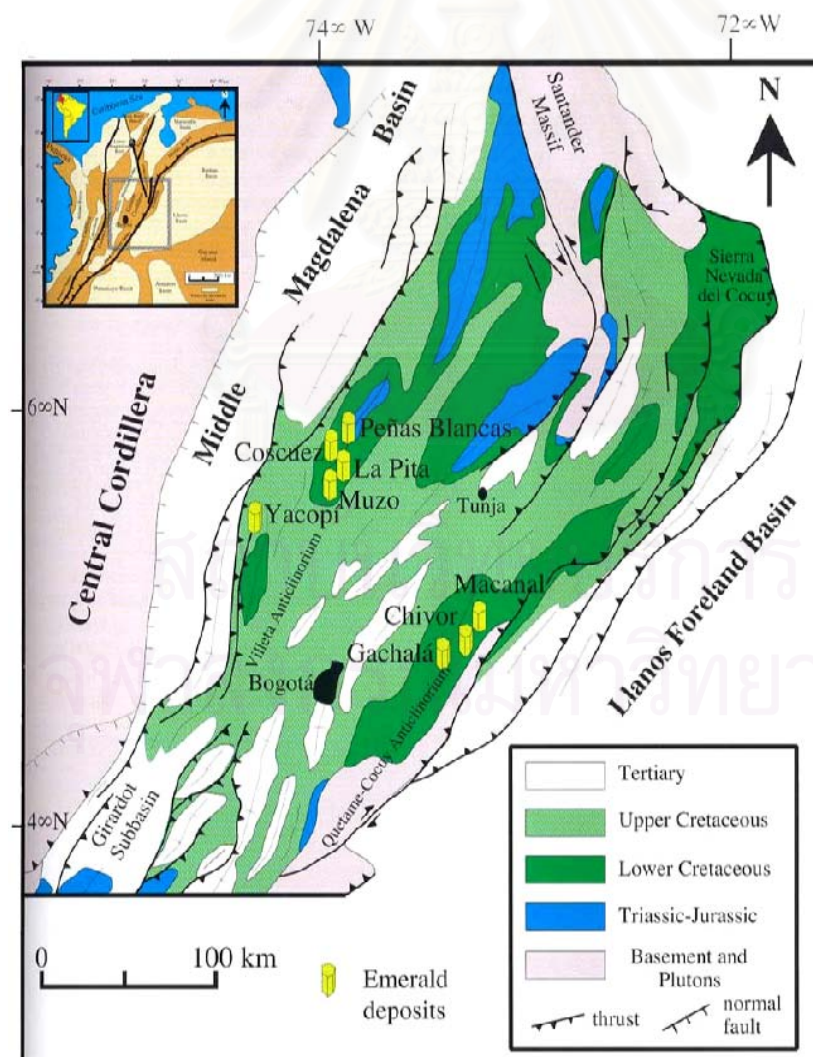


Figure 2.4: The geology in the area of Colombian emerald deposits showing the important mines (Branquet et al., 1999).

Figure 2.5: In Muzo, the Colombian emeralds occur in calcite veins vertically crosscutting in dark, carbonaceous shales with thin layer of limestones (Schwarz, 1998).

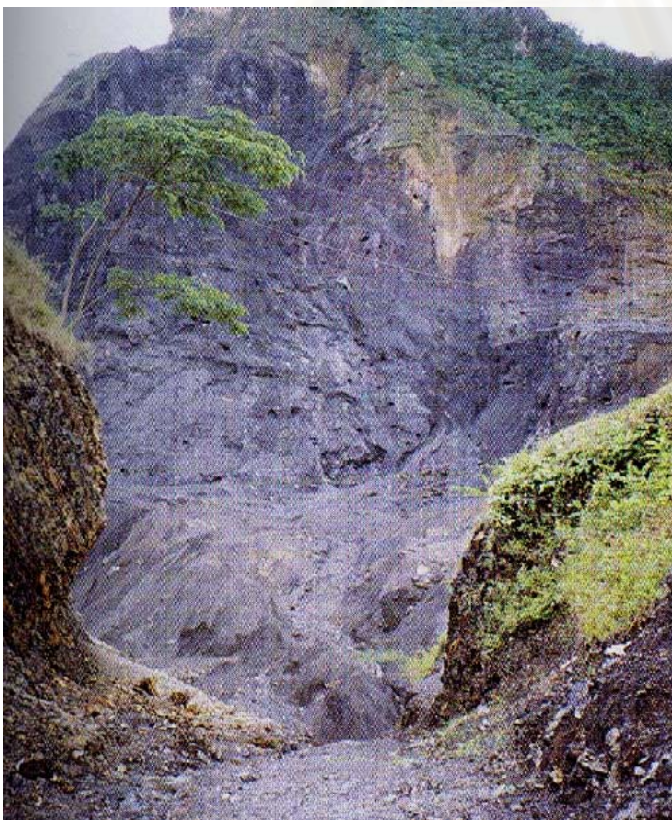
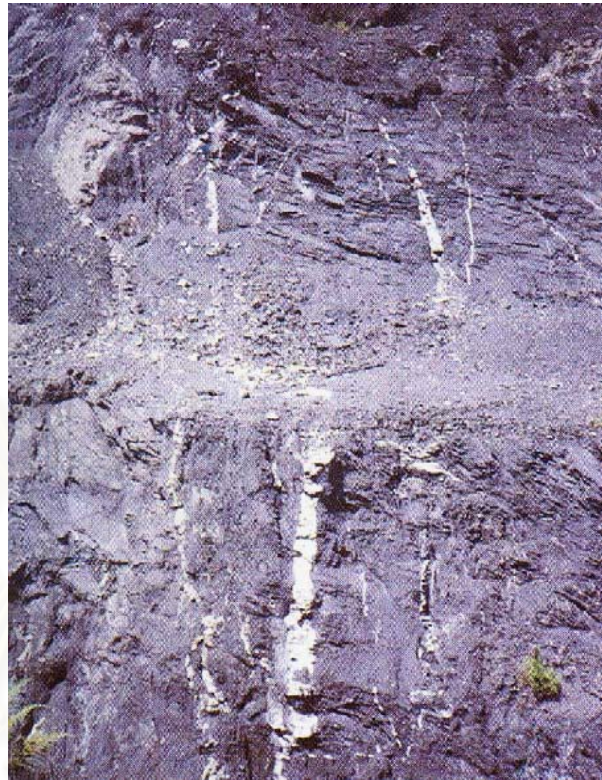


Figure 2.6: The Coscuez mine; Cordillera Oriental (Ward, 1993).



## 2.3 Type II: Schist-related emerald deposits;

### 2.3.1 IIa: Emeralds in schist without pegmatites: e.g. Santa Terezinha deposits in Goiás State of Brazil

The Santa Terezinha deposits are situated at 230 km northwest of Brasília (Figure 2.7) and about 275 km north of Goiânia, the capital of the state of Goiás. This region is a part of the Brazilian Shield. In this region, the Middle Precambrian rocks (belong to the Araxá Group) consist mainly of mica schists and quartzites, varying in thickness from a few hundred to almost 2,000 meters.

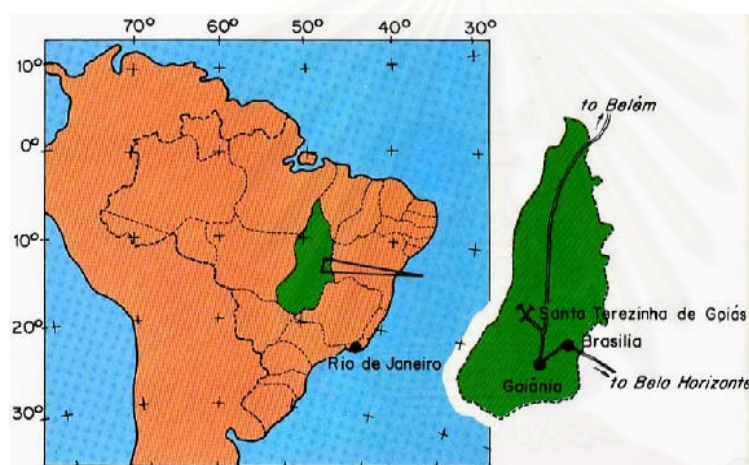


Figure 2.7: The location map of Santa Terezinha deposits at Goiás State, Brazil (Cassedanne and Sauer, 1984).

The emeralds are disseminated in phlogopite schists and phlogopitized carbonate-talc-schists. In the deposit, faults thrusts and shear zones also controlled the infiltration of hydrothermal fluids. Emerald-rich zones are encountered in the core of sheared folds and along foliation planes. Two mineralization types were distinguishable: (Schwarz, 2002).

- 1) A carbonate-rich ore with dolomite, phlogopite, talc, quartz, chlorite, tremolite, spinel, pyrite and emerald;
- 2) A phlogopite-rich ore with quartz, carbonate, chlorite, talc, pyrite and emerald.

The emeralds grew in the talc schist (Figure 2.8) and incorporated inclusions of chromite, pyrite, and talc from the schist. Later tectonic folding caused fracturing of some of the pre-formed emeralds. The mineralized schist provides a high concentration of emerald crystals, for example, more than one kilogram of emeralds in six tons of mineralized schist. In general, the deposit is of the mica-oligoclase-beryl type as defined by Smirnov (1977). The beryllium most likely came from mica (phlogopite/biotite)-rich rocks while the chromium came from the host talc-schists (Schwarz and Giuliani, 2002).

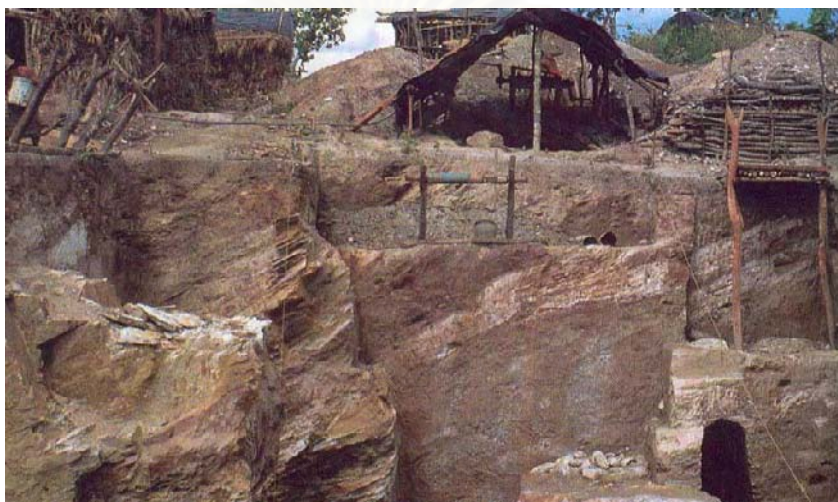


Figure 2.8: The pits of the emerald-bearing talc schist are excavated to remove the ore at the part of the lateritic plateau in Santa Terezinha (Cassedanne and Sauer, 1984).

### ***2.3.2 IIb: Emeralds in pegmatites with schist seams: e.g. Madagascar, Zambia and Brazil (Itabira, Carnaiba and Socoto)***

Most schist-related emerald deposits are associated with granitic rocks, greisen, aplites, pegmatites and quartz veins. It normally shows a dark, clearly schistose zone developed at the contact between pegmatitic and mafic rocks (serpentinite, talc-schist or amphibolite).

These deposits are found in the Precambrian or Paleozoic volcano-sedimentary series. At first, pegmatites and aplites form in the roof region of plutons. The black wall schist around the pegmatites and aplites is composed of phlogopite and plagioclase, derived by exchange of elements (alkali-metasomatism) between granitic rock and the



neighboring serpentinites, actinolite/tremolite schists or talc schists. In some deposits, greisen formation and tectonic deformation are characteristic for the contact zone and hydrothermal fluids to access the alteration zones where emeralds later formed.

The beryllium could come from the decomposition of beryl, feldspar, mica and from phenakite while the chromium may come from mafic rocks such as serpentinite and talc schist (Schwarz and Giuliani, 2002).

### *IIb (1): Mananjary region of Madagascar*

The fourth largest island in the world is Madagascar, which is 1,580 km (1,000 miles) long and 580 km (365 miles) across at its widest. Two major groups of rocks cover the island: ancient Precambrian basement rock and younger sediments. The basement rocks, which cover most of the area, consist mainly of gneisses and schists overlain over wide areas by crystalline limestones and schist-quartzites. Numerous granite intrusions contributed pegmatite mineralization. The pegmatites contain various minerals and are commonly complex, including those suitable for gems and mineral specimens. Both common and gem beryl occur in many bodies (Murdock, 1963 cited in Schwarz, 1994).

Beryl is one of the important gem minerals (more than 50 minerals) found on Madagascar (Chikayama, 1989 cited in Schwarz, 1994). According to Lacroix, 1922 (cited in Schwarz, 1994) provided the first detailed survey account of occurrences of both primary and secondary (eg. alluvial) gem quality beryl there. The primary emerald occurrences are in the vicinity of Mananjary (Figure 2.9).

At the beginning of the 1990s, the worked area around Mananjary comprised at least 50 km<sup>2</sup> (Thomas, 1993 cited in Schwarz, 1994). The Morafeno Mining is performed by independent mine and a number of small mechanized operations (Figures 2.10). The mining have been exploited commercially-rich emerald mineralization that is embedded in mica schist and amphibole schists. Most crystals range from one to three carats, but are low quality crystals.



Figure 2.9: Map showing three commercially mining areas of the emerald deposits in Mananjary, Morafeno (Schwarz, 1994).



Figure 2.10: The overview of Morafeno mining operations shows some deposits. (Schwarz, 1994).

*Ilb (2): Ndola Rural district of Zambia*

The main emerald-producing area in Zambia comprises two deposits, Kamakanga and Kafubu (both in Ndola Rural district), which lie within a few kilometers of each other. This emerald field is located between the Ndola Rural district and Kitwe district of northern Zambia (Figure 2.11), approximately 32 km south-west of Kitwe and 40 km west-northwest of Luansya, near the entrance of the Miku River into the Kafubu River (Bank, 1974 cited in Koivula, 1982). Both Miku and Kafubu areas produce excellent, gem-quality emeralds; some crystals weigh well over 100 ct (Koivula, 1982).

The Zambian emerald deposits occur within rocks of the Muva Supergroup -- which are persistent bands of talc-chlorite-amphibole (tremolite/actinolite)-magnetite schist. They are intruded by pegmatites which occur as feldspar-quartz-muscovite bodies or as minor quartz-tourmaline veins. Contact aureoles of pegmatitic veins with the ultramafic schists are usually altered to phlogopite-biotite-tourmaline aggregates and the minor veins are often concordant to the foliation of the country rocks. The emeralds are found in the contact zones, and sometimes within the quartz-tourmaline veins (Figure 2.12). Muva Supergroup forms the youngest part of pre-Katangan basement immediately southwest of the well known Copper belt of Zambia.

The chromium was probably derived from magnetite in the talc-magnetite schists, as the magnetite in these rocks shows to contain a low % chromium (Bank, 1974 cited in Koivula, 1982). The emeralds are found in biotite-phlogopite schists in which dark brown to black-tourmaline also occurs. Other rocks associated with the emerald- and-tourmaline-bearing mica schists are talc-magnetite schist with secondary quartz veining (Bank, 1974; Sinkankas, 1981 cited in Koivula, 1982).



Figure 2.11: The map of the emerald bearing area in Zambia (Seifert et al, 2004).



Figure 2.12: Emerald associated with quartz and tourmaline veins in talc schist. (Moroz and Eliezri, 1998).



*Ilb (3): Carnaiba/Socoto deposits in Bahia State of Brazil*

Carnaiba and Socoto deposits are in the district of Bahia. The geological situation in these regions are marked by the intrusion of granite batholiths into the Serra de Jacobina unit which is composed of schists, quartzites and serpentinites (Schwarz and Giuliani, 2002).

The rocks of the Serra de Jacobina were subdivided by Couto et al. in 1978 (cited in Schwarz and Eidt, 1989), into the Jacobina Group – which consists of an ultrabasic basement unit (with an age of 2.4-2.7 billion years) which, together with the granite, is important for the formation of the emerald occurrence) and the Itapicuru Complex in Figure 2.13 (Griffon et al., 1967; Barbosa, 1973; Santana et al., 1980; Santana et al., 1981 cited in Schwarz and Eidt, 1989).

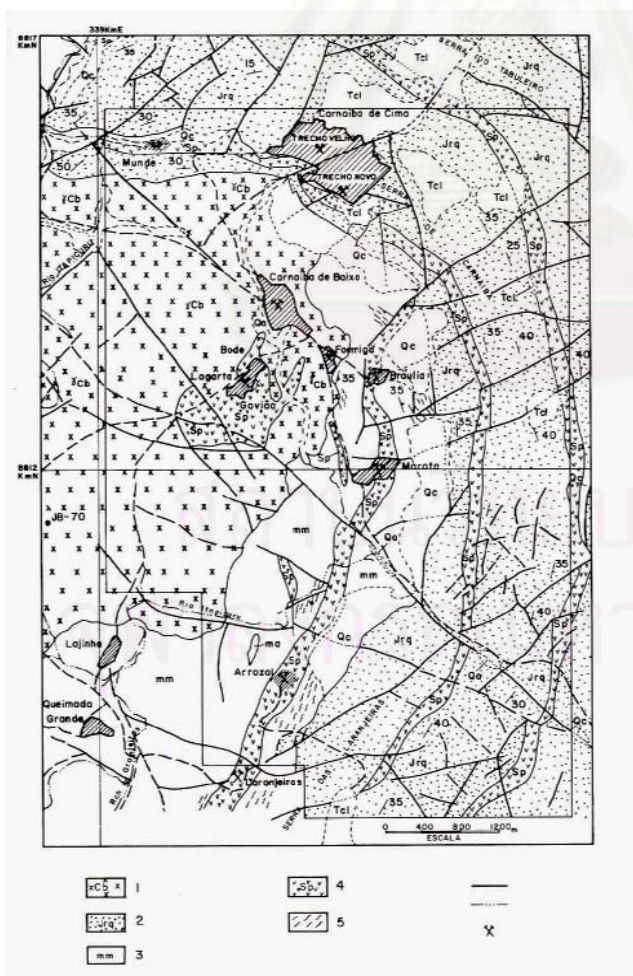


Figure 2.13: Geological map of the Carnaiba region (Schwarz and Eidt, 1989)

The Carnaiba/Socoto deposits were related to the zones of the strongest tectonic activities, which were characterized by the association of mafic-ultramafic rocks or their metamorphic derivatives (Cr/Fe source rocks) with pegmatites (provide Be) (Schwarz et al., 1990). It was the Cr-rich serpentinites and peridotites that were intruded by pegmatites and phlogopite or biotite/ phlogopite were in direct contact with the pegmatite and quartz veins. These micas are oriented parallel to the vein walls and show a lepidoblastic texture which gives the rock a schist-like appearance. The rock is therefore called mica-schist and is the emerald host rock. The emeralds occur either in the forms of aggregates or as single crystals, (Griffon et al, 1967, Barbosa, 1973, Santana et al, 1980, Santana et al, 1981 Couto and Almeida, 1982, Moreina and Santana, 1982 and Couto, 1985 cited in Schwarz and Eidt, 1989).

In these deposits, it is believed that metasomatic processes accompanying the intrusion of pegmatite bodies into the ultramafic rocks are responsible for the formation of emerald. The temperature of geochemical processes resulting in mineral association range of 600 - 800 °c (Griffon et al., 1967 cited in Schwarz and Eidt, 1989).

A mining attempt usually begins with a limited open pit and is worked with uncomplicated hand tools (Figure 2.14). These quarries are known in Brazil as garimpo, as they are worked by independent miners as called 'garimpeiros'.

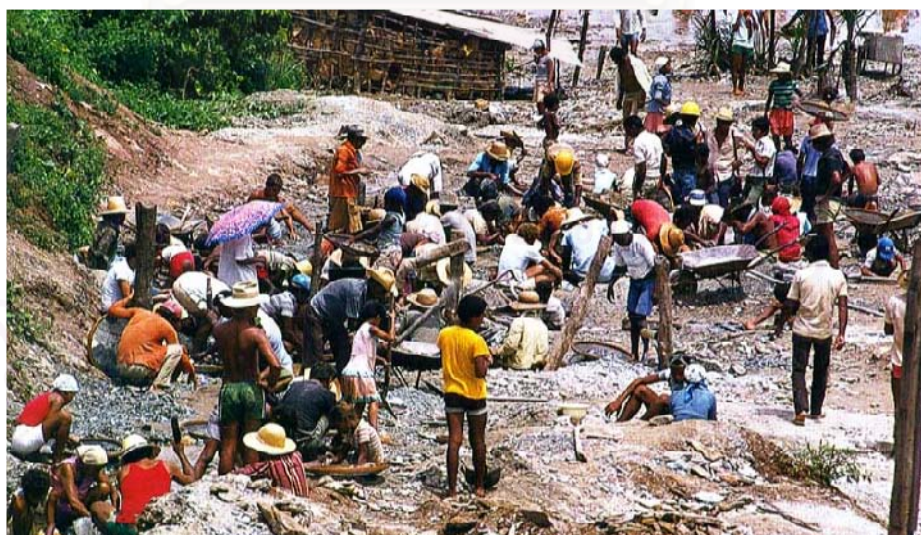


Figure 2.14: The 'garimpeiros' who are working by hand with pick, hammer, chisel and crowbar (Schwarz, 2002).

*Ilb (4): Itabira deposits in Minas Gerais State of Brazil*

The significant emerald deposits of Itabira township are Capoeirana and Belmont mines (Figure 2.15). The Belmont mine lies 13 km SE of the town of Itabira and 120 km E of Belo Horizonte, the state capital of Minas Gerais State (Schwarz and Giuliani, 2002). The Capoeirana mine was discovered only about 10 km SE of Belmont mine in Figures 2.16 (Epstein, 1989).

According to Schwarz and Giuliani (2002), the stratigraphy of the Itabira region is characterized by two rock series of Precambrian age: the crystalline basement (so-called 'Série pré Minas' after Pfulg, 1968) and the overlying meta-sediments (so-called 'Supergrupo Minas'). The basement rocks are composed mainly of paragneiss and poly-metamorphic migmatites.

In the area of the emerald occurrence, a belt of schists dominates, stretching in a north to north-easterly direction. The width of the belt varies between 750 and 1200 meters. Leuco-gneisses occur symmetrically on both sides of this belt. The schist belt, together with the mafic rocks were strongly folded. Emeralds occur in black biotite/phlogopite schists in green chlorite schists or in kaolin masses (altered pegmatite). Crystals of lower quality are also found in quartz masses (Schorsch, 1973; Schorsch and Guimarães, 1976; Schorsch et al., 1982 cited in Hänni, Schwarz and Fischer, 1987).

The average emerald-content of the biotite schist is 165s/ton. The both mines are the richest emerald occurrence in Brazil so far as emerald-content of the parent rock in relation to the average quality of the emeralds is concerned (Hänni, 1987).



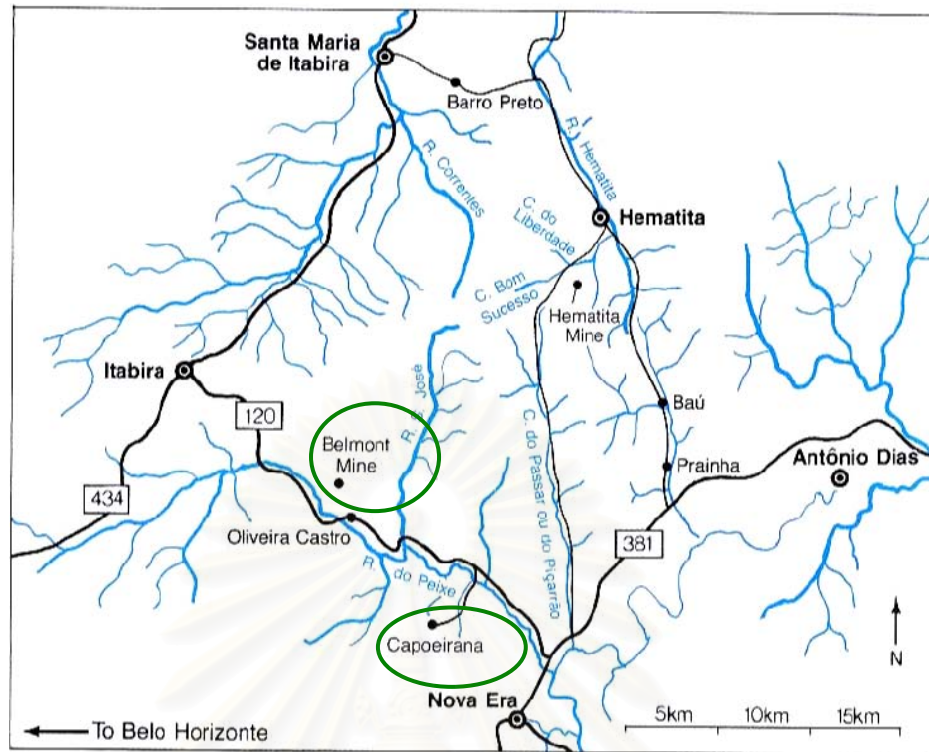


Figure 2.15: The map of two important emerald deposits—the Capoeirana and the Belmont mines in Itabira township, Brazil (Epstein, 1989).



Figure 2.16: The present working of the Capoeirana emerald deposits (Schwarz, 1998).

## CHAPTER III

### GENERAL OBSERVATION

#### 3.1 Crystal habits and Structure

Emerald which is a variety of beryl belongs to the highest symmetry class of the hexagonal crystal system. The emerald crystal looks the same from both the top and the bottom, as the opposite poles are identical. Hexagonal crystals can exhibit habits ranging from acicular through prismatic to thin tabular. The c-axis is the principal axis of symmetry and is six-fold, that is the crystal can rotate in six increments of  $60^\circ$ . Emerald crystals are more modest, as they only form more or less prismatic crystals shown in Figure 3.1. Very few tabular emerald crystals exist that can be rather complex with maximum number of crystal faces. This applies less to the prism faces than to the terminations, which can be bound by numerous faces of the most varied forms, mainly bipyramidal forms (Vrba, 1881 cited in Hochleitner, 2002)..

Emerald often forms in tectonically active areas. As a result the crystals are broken and commonly contain many fractures. Broken emerald crystals are sometimes found frozen in calcite (Figure 3.2: B). The individual pieces are slightly separated from one another during period of tectonic activity. Generally, the most complex emerald crystals are from localities in which crystal form in vugs or cavities.



Figure 3.1: Hexagonal prism (left) and dihexagonal prism (right) of emerald crystals (Vrba, 1881 cited in Hochleitner, 2002).

Figure 3.2: Specimens of emeralds with dark organic matter (A) and its specimen associated with calcite (B) from Muzo, Colombia (Photo by Pavaro, T.).

The primary elements of emerald are beryllium, aluminum, oxygen and silicon, that has the general formula:  $\text{Be}_3\text{Al}_2(\text{Si}_6\text{O}_{18})$ . The dominant features in the structure consists of rings of silicon atoms, each surrounded by four oxygen atom in tetrahedral arrangement. These rings forming hollow columns parallel to the *c*-axis of the crystal. Between the rings lie the aluminium and beryllium atoms, each aluminium atom is surrounded and coordinated with an octahedral group of six oxygen atoms, while each beryllium atom is surrounded by four oxygen atoms on a distorted tetrahedron. The linked silicon-oxygen of neighbourings  $\text{Si}_6\text{O}_{18}$  rings both laterally and vertically in the beryl structure is called cyclosilicate (Figure 3.3). The structure is thus like a honeycomb without atomic centre nearer than  $2.55 \text{ \AA}$  to a centre of the open channels that forms parallel to the *c*-axis (Deer, 1992 cited in Hochleitner, 2002). Other elements not included in the general formula, such as potassium, sodium and cesium, can fit into these hollow channels. In addition, molecules such as the water or carbon dioxide commonly reported in beryl can also fit into these spaces. The colour of emerald is determined by foreign elements that are built into the lattice. The important ones are chromium and vanadium (Hochleitner, 2002).

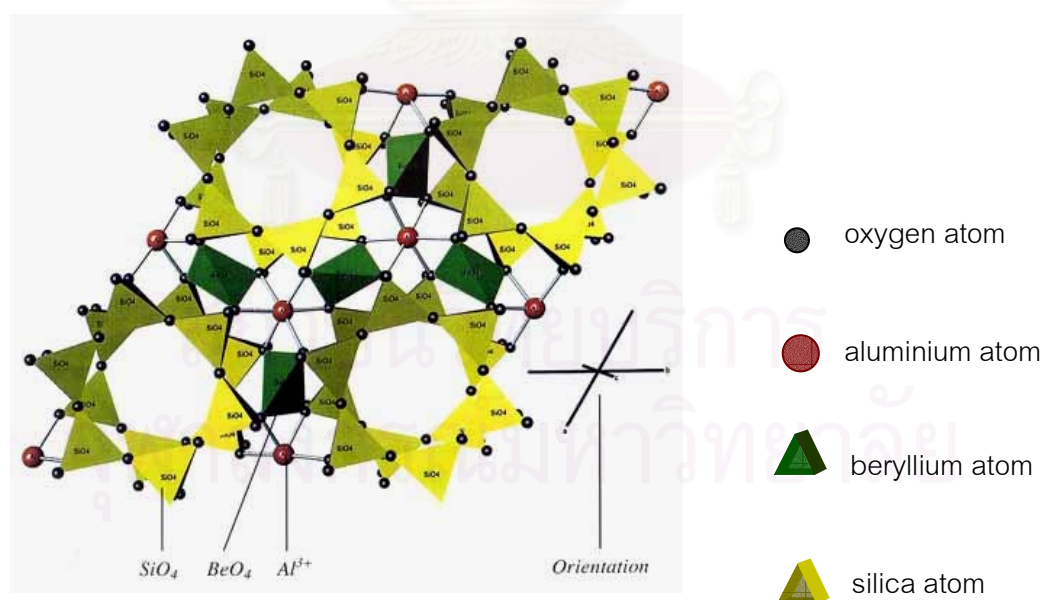


Figure 3.3: The beryl structure composed of ring-shaped arrangement of  $\text{SiO}_4$  tetrahedra (Hochleitner, 2002).



Two types of water may occur, the “type I” water molecule occurs alone, and is oriented in the hollow channels with its oriented by  $90^\circ$  to the c- axis of the beryl. The “type II” water molecule is associated with nearby alkali ion on the molecule dipole and lies with its symmetry axis parallel to the hexagonal c-axis direction (Figures 3.4a, 3.4b and 3.4c) (Wood and Nassau,1976).

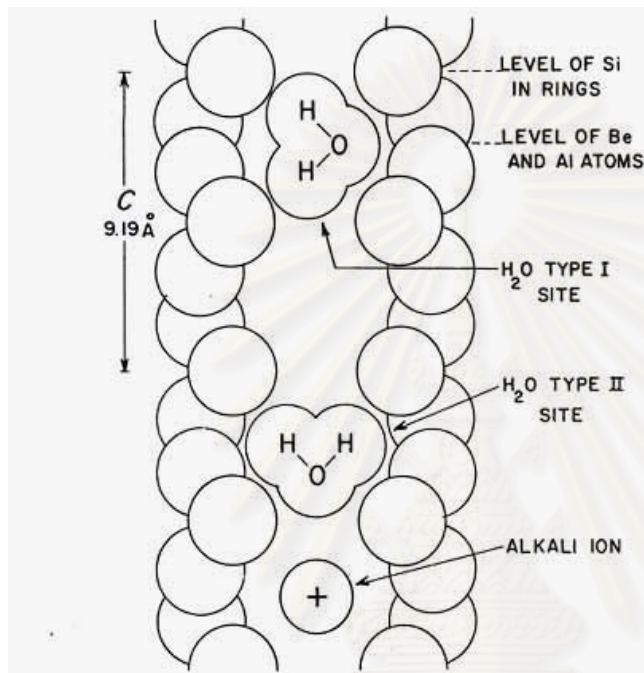


Figure 3.4: Side view of a beryl structure, only the oxygen atoms of the wall being shown, with possible positions and orientations of alkali ions and both types of water molecules located in channels sites, (Sinkankas, 1989).

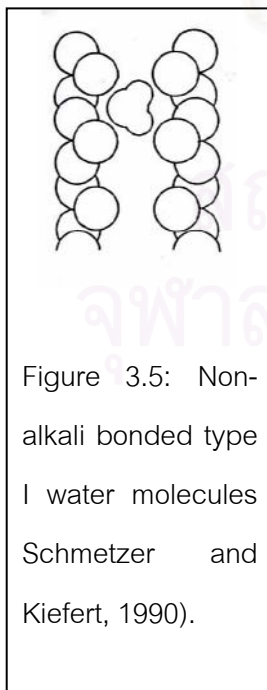


Figure 3.5: Non-alkali bonded type I water molecules Schmetzer and Kiefert, 1990).

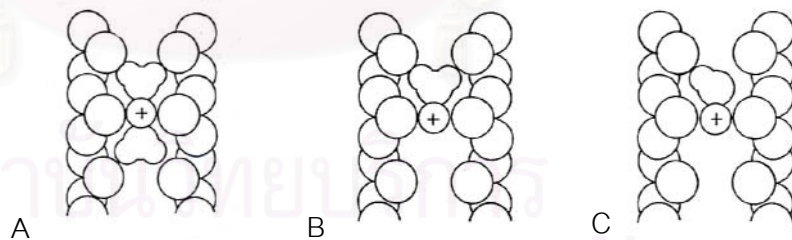


Figure 3.6: Alkali bonded type II water molecules; (A) type IIa water molecules showing in sequence  $H_2O-Na-H_2O$ , (B) type IIb water molecules showing  $H_2O-Na-X$  and (C) type IIc water molecules showing  $OH-Na-X$  (hydroxyl groups), with X representing vacancies in channel sites of the beryl structure (Schmetzer and Kiefert, 1990).

### 3.2 Visual appearance

All emerald samples used in this study consists of 148 polished and one cabochon stones from 5 countries of origin, namely Brazil (29 samples), Colombia (30 samples), Madagascar (30 samples), Zambia (30 samples) and Nigeria (30 samples). The samples are mostly transparent to semi-transparent and vitreous luster. Their weights are from approximately 0.3 to 2.8 carats. They are slightly different hue and tone, mostly ranges of colour are between bluish green to green (Figure 3.5).

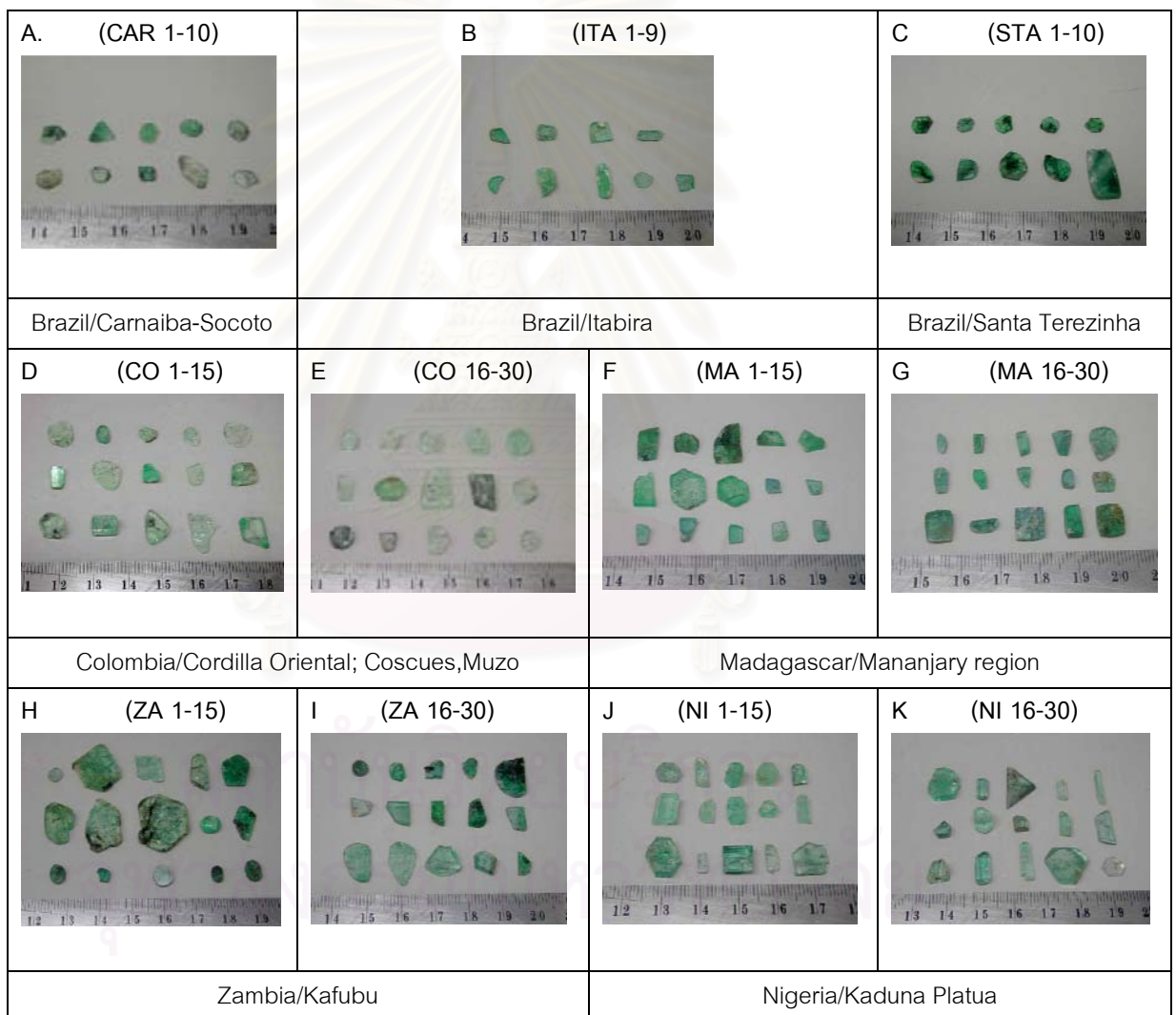


Figure 3.7: Showing all samples in this study; 29 samples of Brazilian emeralds (A to C), 30 samples of Colombian emeralds (D to E), 30 samples of emeralds from Madagascar (F to G), 30 samples of Zambian emeralds (H to I) and 30 samples of Nigerian emeralds (J to K).

The colours of 149 emerald samples are compared with the standard gem colour of Gemological Institute of America (GIA GemSets) under the daylight lamp. The lamp's temperature is 6500 K. The colour code comprises hue, tone and saturation. Hue is its colour. Tone is the lightness to darkness of the hue, the higher number is the darker tone (scale 0 -10). Saturation is the strength and purity of the hue, the higher number is the stronger hue (scale 1-6). In this study, the range of colour (depending on the depth of samples) or colour code of emeralds from each area are reported by the following sequence, tone / saturate / hue. In general, Nigerian emeralds are very light to medium light / very slightly grayish to moderately strong / very slightly bluish green to green. The Colombian emeralds are very light to medium / very slightly grayish to moderately strong / very slightly bluish green to green. The Brazilian emeralds are very light to dark / very slightly grayish to moderately strong / very slightly bluish green to green. The emeralds from Madagascar are light to medium / very slightly grayish to strong / very slightly bluish green to bluish green. Finally, the emeralds from Zambia are very light to medium / very slightly grayish to strong / very slightly bluish green to bluish green. The colour codes of all the samples are listed in Appendix I (Figure I.1 and Table I.2).

### 3.3 Physical and Optical properties

Beryl is one of the durable minerals. For example, large numbers of beryl crystals have been found to retain traces of crystal faces despite prolonged chemical attack and abrasion with other stones. Fracture surfaces are usually brilliant, smooth and conchoidally curved like freshly broken glass.

The specific gravity (SG) and optical properties, such as refractive index (RI), birefringence and fluorescence were examined. A summary of the results are shown in Table 3.1 and the results of all samples are given in Appendix I (Table I.3).

High optical values are found in emeralds from (Santa Terezinha) Brazil. Medium to high optical values are found in emeralds from Madagascar, Zambia, (Carnaiba/Socoto and Itabira) Brazil. Low to medium optical values are found in

emeralds from Colombia and Nigeria. High specific gravity values are found in emeralds from Brazil, Madagascar and Zambia whereas emeralds from Colombia and Nigeria are low.

Table 3.1: Summary of the gemological properties of emeralds from Brazil, Madagascar, Zambia, Nigeria and Colombia.

Sample	RI		Birefringence	SG	Fluorescence	
	e-ray	o-ray			LW	SW
Nigeria/NI (Type Ia)	1.560 – 1.570	1.569 – 1.582	0.007 – 0.016	2.601 – 2.687	Inert	Inert
Colombia/CO (Type Ib)	1.561 – 1.572	1.571 – 1.582	0.007 – 0.018	2.612 – 2.731	Inert	Inert
Brazil / STA (Santa Terezinha) (Type IIa)	1.579 – 1.586	1.592– 1.597	0.010 – 0.014	2.723 – 2.812	Inert	Inert
Madagascar/MA (Type IIb)	1.570 – 1.584	1.581 – 1.598	0.008 – 0.023	2.661 – 2.785	Inert	Inert
Zambia/ZA (Type IIb)	1.573 – 1.590	1.580 – 1.597	0.006 – 0.016	2.582 – 2.782	Inert	Inert
Brazil / CAR (Carnaiba/Socoto) (Type IIb)	1.571– 1.581	1.581 – 1.591	0.009– 0.014	2.705 – 2.769	Inert	Inert
Brazil / ITA (Itabira) (Type IIb)	1.571 – 1.580	1.581 – 1.591	0.006 – 0.011	2.690 – 2.727	Inert	Inert



### 3.4 Internal Characteristics

Certain inclusions, especially in emerald, are distinctive and sometimes can even serve to identify a particular deposit from which they came. Inclusions also provide evidence as to the geochemical environment in which the emerald crystals grew. In general inclusions can be divided into two types based on the time of formation as follows:

1. Primary inclusions are minerals or fluids which formed before or contemporaneously with the host emerald and were enveloped by the emerald crystal as it grew, such as some fluid inclusion; tube and foreign mineral inclusions; mica, quartz, pyrite, etc.
2. Secondary inclusions are those which developed afterwards; such as iron stained in fissures, micas developed to hematite.

In this study, a GIA binocular microscope with magnifications up to 70 times (70x) was used to examine internal and external characteristics of emeralds from different country of origins. Among the common internal features in those emeralds are foreign crystal inclusions, liquid inclusions and growth structures. Some mineral inclusions were further identified their species by using a Renishaw Laser Raman spectroscope System 1000 attached with a Leica microscope at the GIT. The inclusions were also studied under objective lenses with x5 x20 and x50 magnifications in both transmitted and reflected light microscope. The identification of a mineral inclusion by Raman spectroscopy was done by matching the pattern of Raman peak shift with those of the Renishaw and GIT databases (Figures 3.8 and 3.9).

Williams et al. (1997) reported the database of the Raman Shift peaks of common inclusions which are summarized in Appendix I (Table I.4). Some inclusions were photo-micrographed by a Nikon digital camera model Coolpix 950 attached to a gem microscope or by the Leica microscope attached to the Laser Raman spectroscope.



Figure 3.8: A binocular microscope at the Gem and Jewelry Institute of Thailand (GIT)



Figure 3.9: A Renishaw Laser Raman Spectroscopy at the Gem and Jewelry Institute of Thailand (GIT).

#### 3.4.1 Nigeria:

##### *Kaduna Platua*

Some of 30 samples of Nigerian emerald examined in this study contain mica inclusions (Figure 3.10 and Figure 3.11 with the Raman spectrum). One sample exhibits opaque-black platelets inclusions of ilmenite as identified by Laser Raman (Figure 3.12). Raman spectrum shows rutile inclusions in Figure 3.13. Raman spectrum of carbonate daughter crystals in multi-phase inclusions is shown in Figure 3.14. Strong relief, colourless to grey mineral inclusions of fluorite are shown in Figure 3.15. Various types of fluid inclusions, such as 2 phase (l,g) , 3 phase (s,l,g) or multi-phase inclusions (Figures 3.16 and 3.17), which look like the 'classic' fluid inclusions observed in Colombian emeralds in Figure 3.18. Growth phenomena, such as internal growth lines

and coloured zoning can be observed by looking in the direction of c-axis (Figure 3.19). Elongated, strong relief negative crystal can be observed in Figure 3.20.



Figure 3.10: Reddish brown to dark brown biotite inclusions in cluster (left) NI10, 70x and book-like appearance (right), d; NI12, 100x.

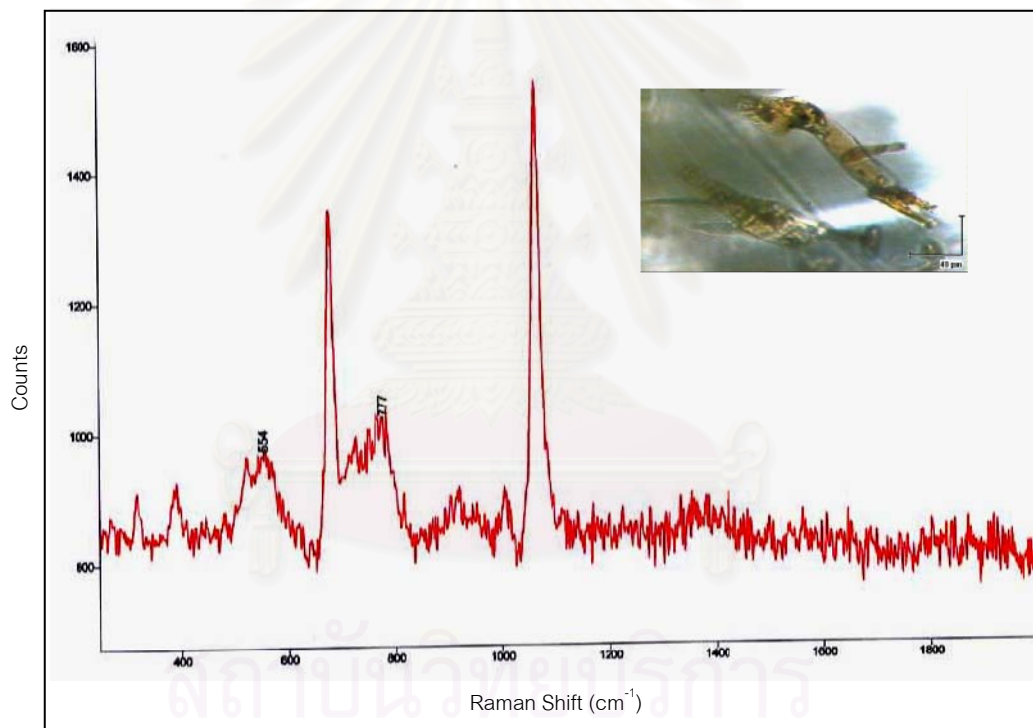


Figure 3.11: Raman spectrum of irregularly shaped biotite inclusions in NI13 (peaks at 554 and 777  $\text{cm}^{-1}$ ).

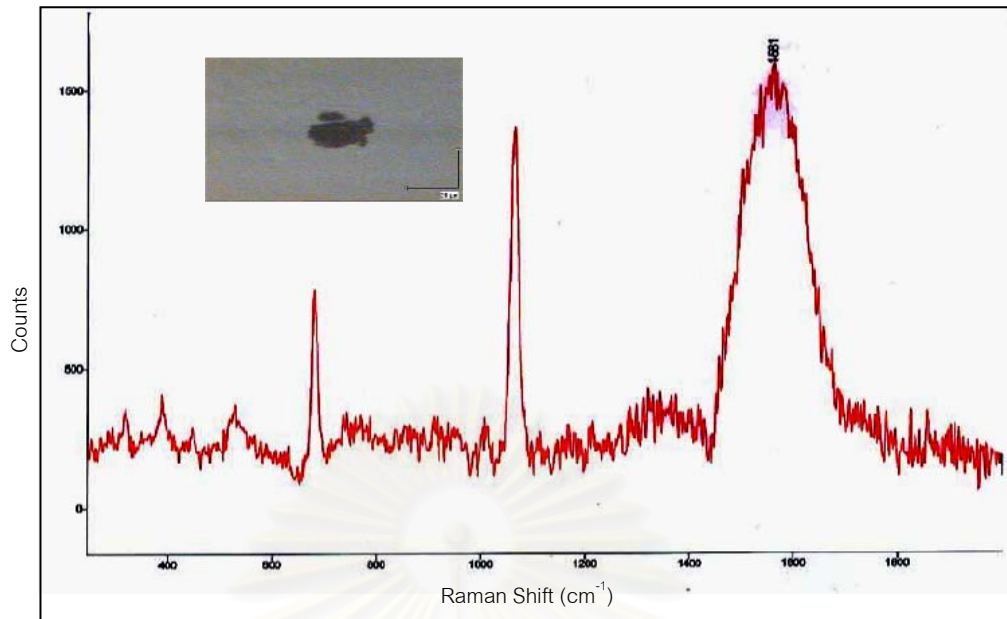


Figure 3.12: Raman spectrum of ilmenite inclusions, which is irregularly shaped grains with metallic luster (NI24, peaks at 398, 681, 1377 and 1581  $\text{cm}^{-1}$ ).

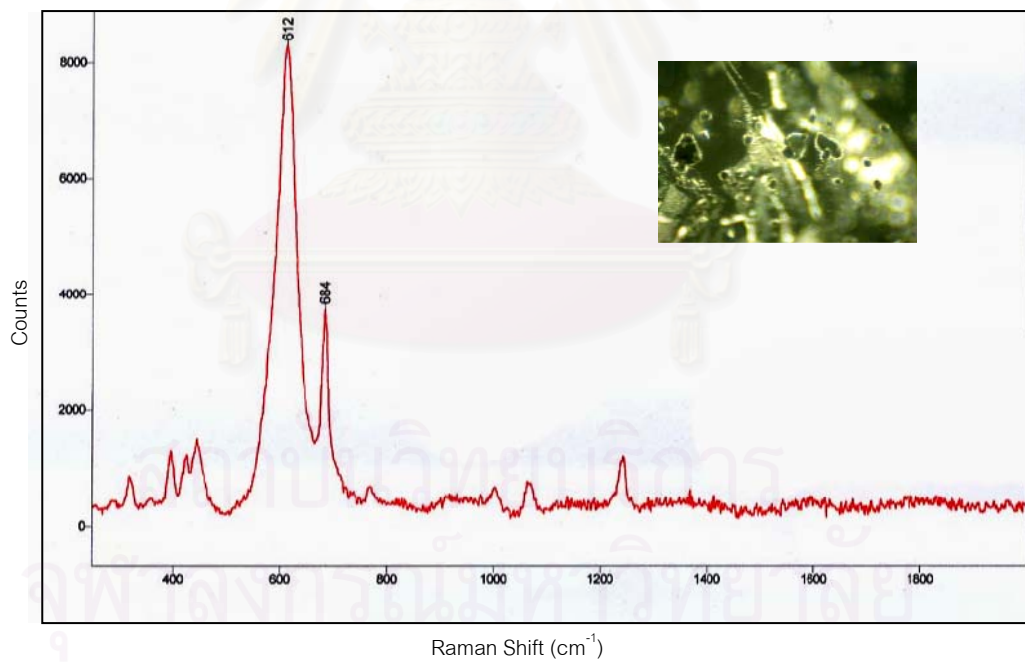


Figure 3.13: Raman spectrum of rutile inclusions, showing strong relief with metallic luster (NI30, peaks at 241, 448 and 612  $\text{cm}^{-1}$ ).

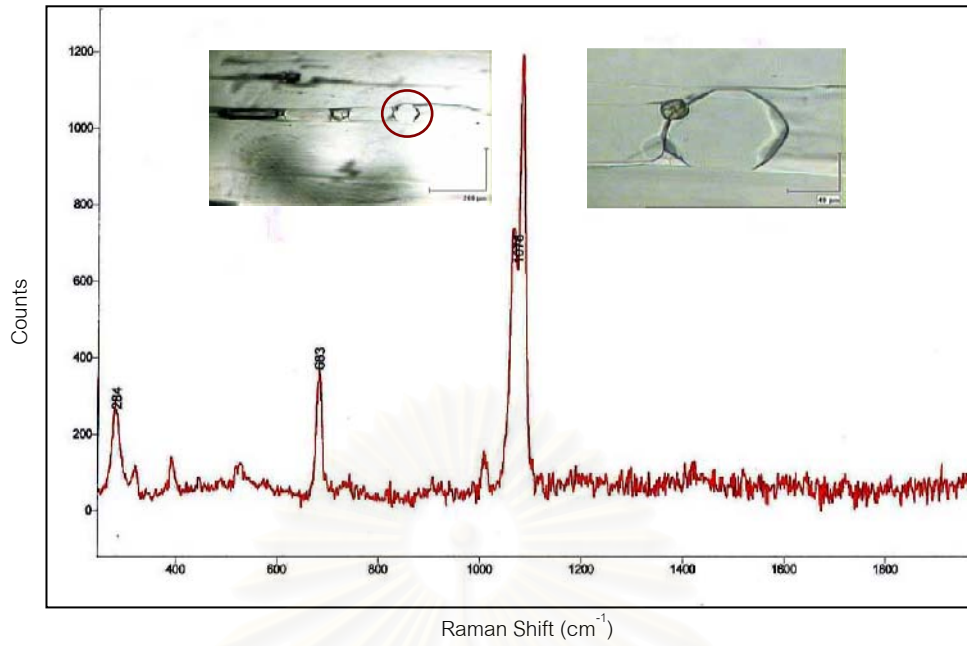


Figure 3.14: Raman spectrum of carbonate daughter crystal in multi-phase inclusions (NI19, peaks at 284, 683 and 1076  $\text{cm}^{-1}$ ).

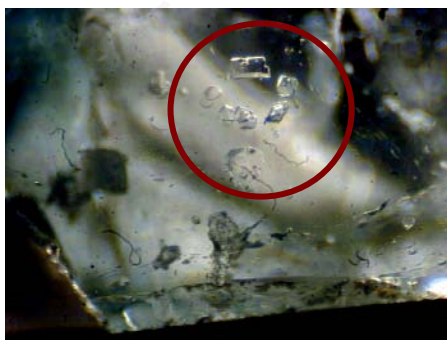
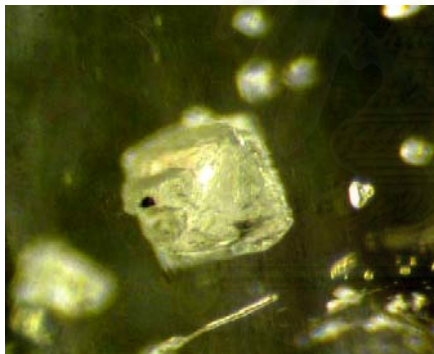


Figure 3.15: Strong relief, colourless to grey perfectly developed octahedral crystals (top) and irregularly shaped mineral inclusions (bottom) of probable fluorite (NI21, 120X).



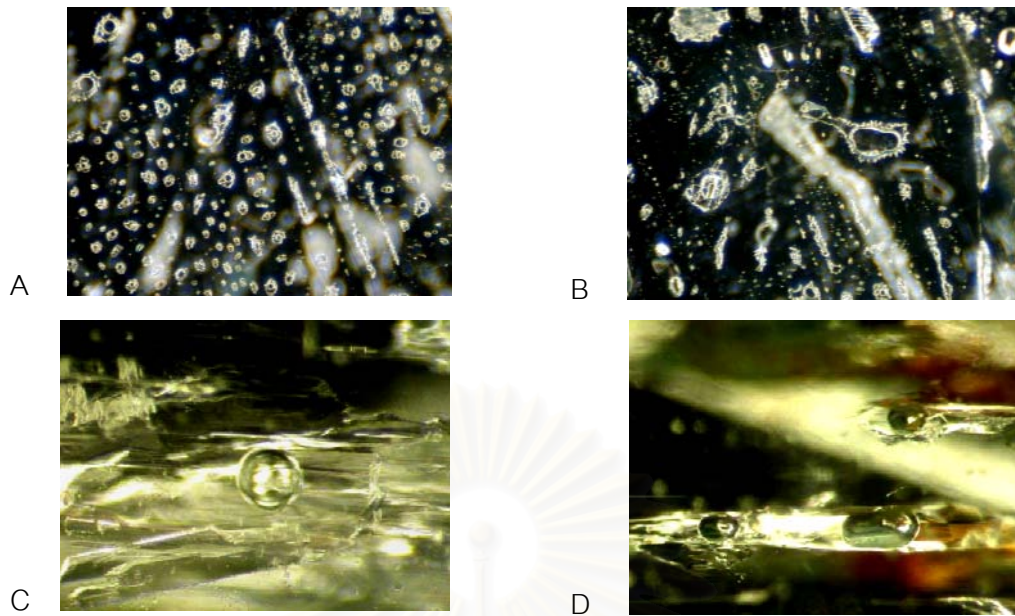


Figure 3.16: Secondary cavities in healing fissures showing rounded, irregularly shaped or 'jagged' outlines in (NI15, 70x); (A and B). And 2-phase fluid inclusion of the 'l,g' type (NI16, 120x); (C) and 3-phase fluid inclusions of the 's,l,g' type (NI10, 100x); (D).

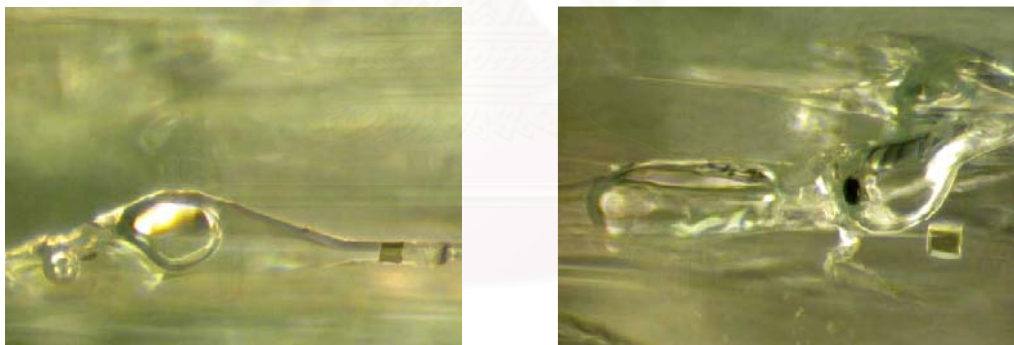


Figure 3.17: Both of primary multi-phase fluid inclusions of the '2s,l,g' type with rhombohedral (may be carbonate) and brown to very dark brown crystal (NI10, 120x).

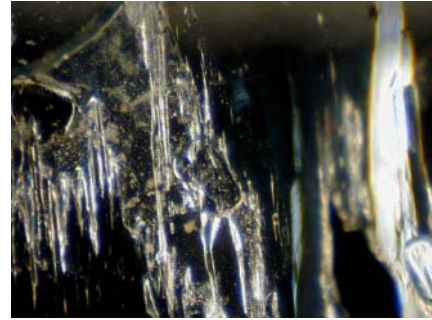
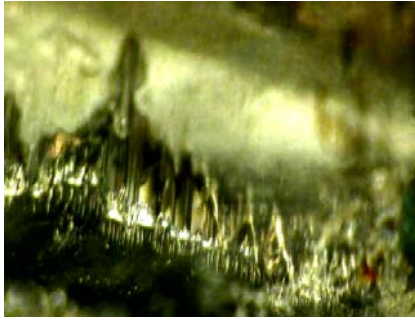


Figure 3.18: The cavities may show elongated (partly jagged) outlines. Mostly cavities often more or less flattened, and the cavity walls seem to be uneven (NI10 and NI12, 70x).



Figure 3.19: Elongate shape of isolated negative crystal in NI10, 100x.

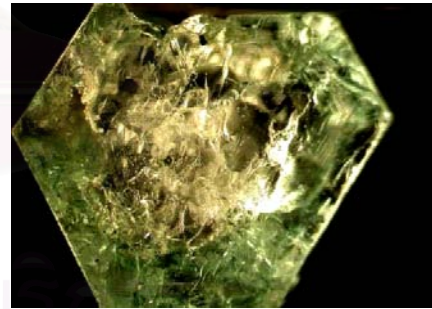
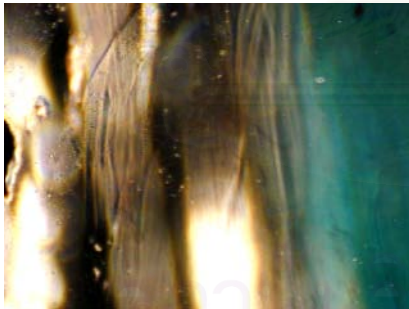


Figure 3.20: Pronounced hexagonal-concentric growth structures, parallel to the prism, and the basal faces (internal growth lines in NI12, 70X) and (colour zoning in NI9, 40X).



### 3.4.2 Colombia:

#### *Cordillera Oriental, Muzo, Coscues and Chivor mines*

The common mineral inclusions of Colombian emerald found in this study include pyrite (Figure 3.21), albite, carbonate and organic matters, which can be identified by Raman spectra (Figures 3.22 to 3.25). Most samples show compact aggregate of black tiny organic matters and carbonate inclusions (Figure 3.26). Elongated or jagged outlines of 2, 3 or multi-phase inclusions are characteristic in Colombian emerald (Figure 3.27). Primary elongated cavities parallel to the emerald's c-axis (growth tube) and strongly developed growth structures are shown in Figures 3.28 and 3.29.

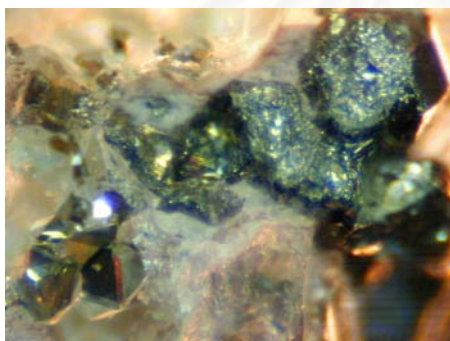


Figure 3.21: Pyrite crystals with golden-yellow metallic luster exposed on the surface (CO3, 60X).

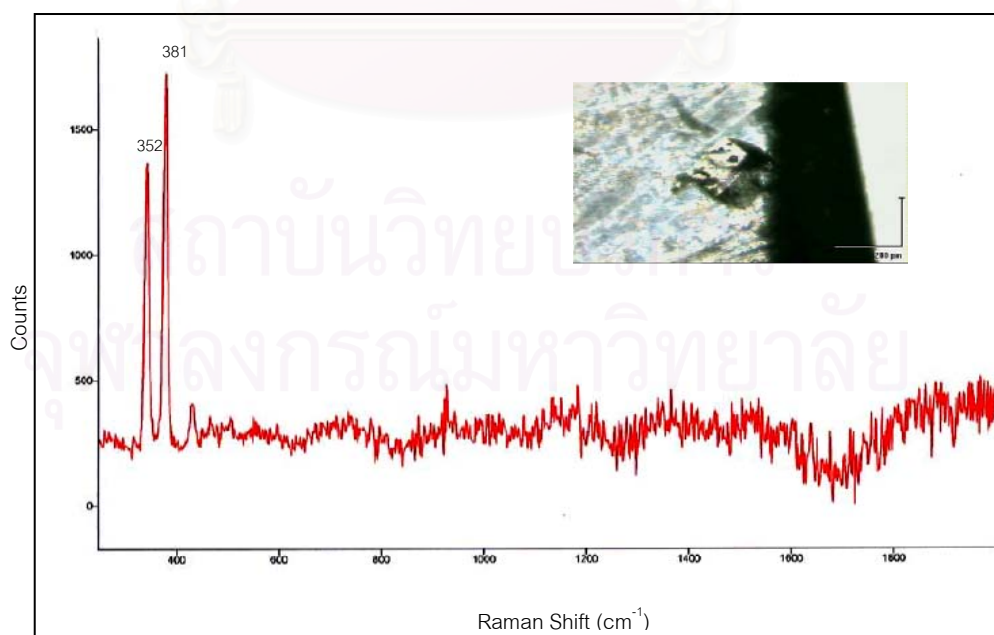


Figure 3.22: Raman spectrum of pyrite inclusion (CO9, peaks at 352, 381 and 443  $\text{cm}^{-1}$ ).

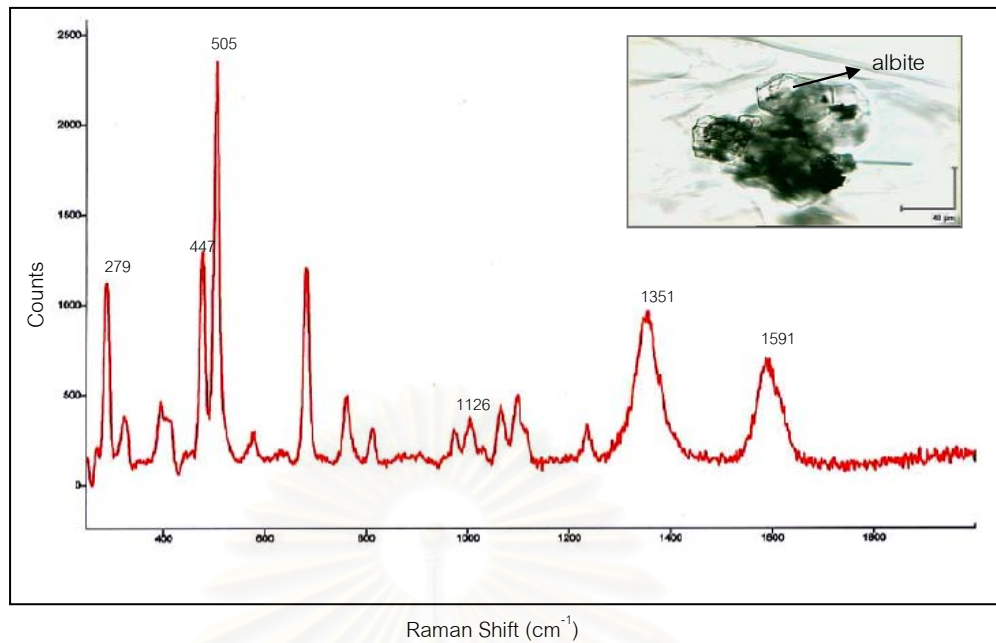


Figure 3.23: Raman spectrum of albite (colourless to greyish) inclusions associated with black organic matters in CO<sub>2</sub>; (albite: peaks at 279 and 1126) and (organic peaks at 447, 505, 1351 and 1591 cm<sup>-1</sup>).

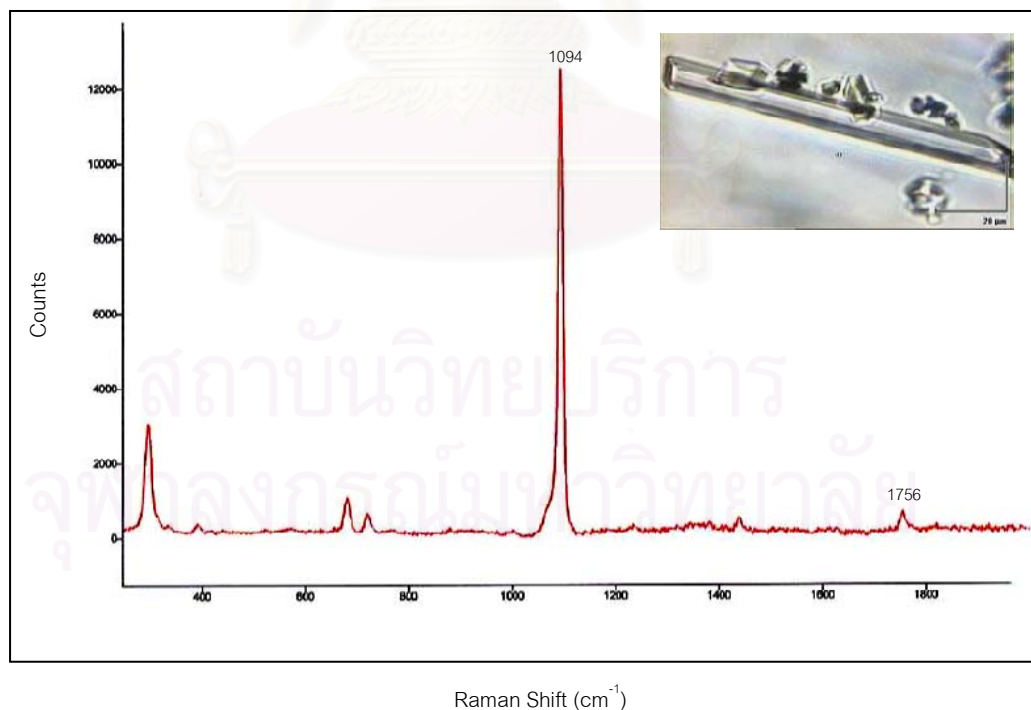


Figure 3.24: Raman spectrum of translucent elongated carbonate crystals (strong relief), (CO<sub>4</sub>, peaks at 1094 and 1756 cm<sup>-1</sup>).

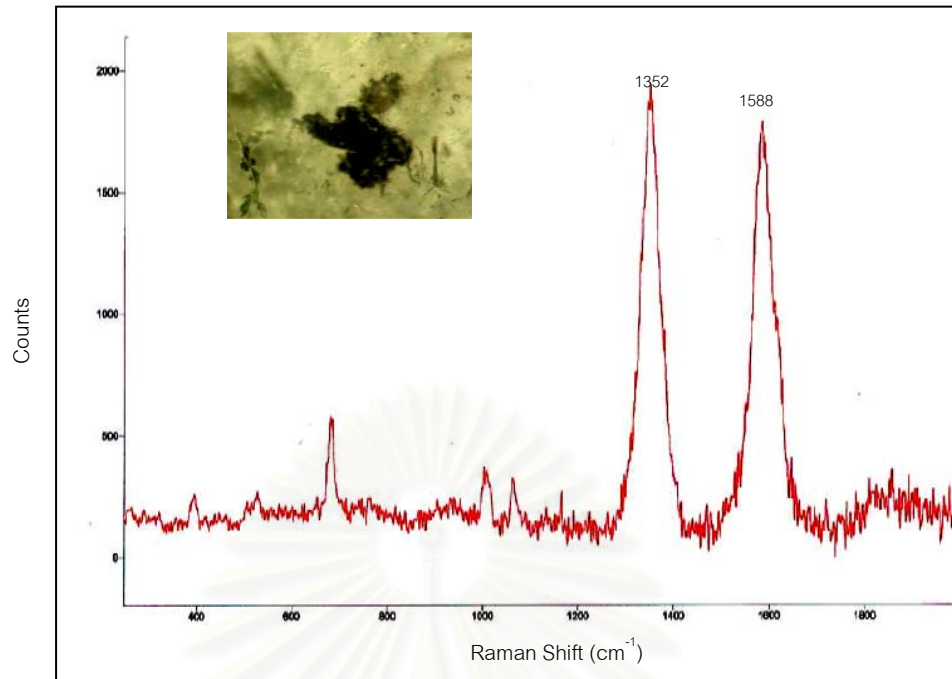


Figure 3.25: Raman spectrum of compact aggregated black organic matter (strong relief), (CO21, peaks at 1352 and 1588  $\text{cm}^{-1}$ ).

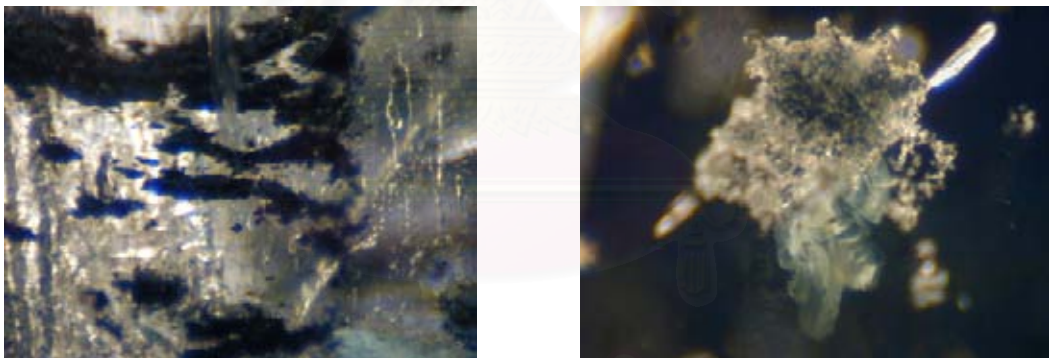


Figure 3.26: Compact aggregated black organic (left); CO-26, 50x and strong relief of gray agglomerate carbonate inclusions (right); CO13, 70x.



Figure 3.27: Primary cavities with a 3-phase inclusion, which composed of s,l,g (Top) and the multi-phase inclusions containing a rounded vapor bubble, a square shaped crystal and a black crystals (2s,l,g); (bottom) showing jagged outline in CO3.

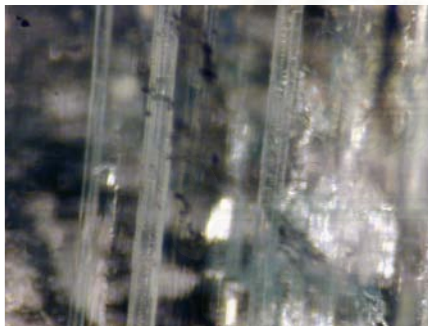
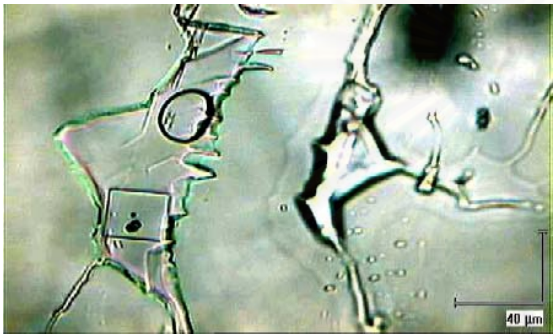


Figure 3.28: Elongated large growth tubes (CO26, 70x).

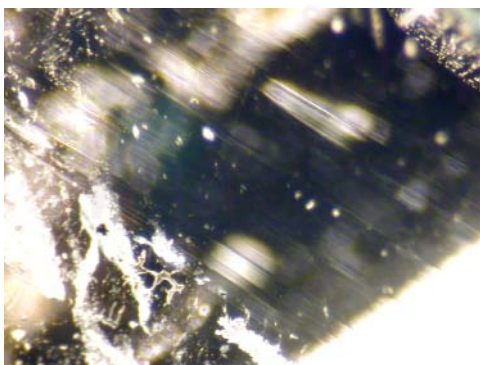


Figure 3.29: Well developed growth structures perpendicular to c-axis (CO7, 40x).



### 3.4.3 Santa Terezinha deposits of Goias State in Brazil:

The most important inclusions in 10 emerald samples from Santa Terezinha deposits of Goias State are pyrite, carbonate and chromite (Figures 3.30 to 3.32), which some of these inclusions were confirmed by the Laser Raman shown in Figures 3.33 to 3.34. Raman spectra of low relief quartz and talc inclusions (transparent, colourless and irregularly-shaped) are shown in Figures 3.35 and 3.36. Among the other inclusions numerous black-opaque inclusions of spinel were also observed and in general formed more or less compact aggregates (Figures 3.37 and its Raman spectrum in Figure 3.38). Light brown, irregularly-shaped mica inclusions are randomly distributed in the host crystal as shown in Figure 3.39 together with its Raman spectrum. 'Dust-particles' and secondary inclusions of iron stain in fissures and fractures can be seen in Figures 3.40 and interference colour of oil-filling in Figure 3.41.

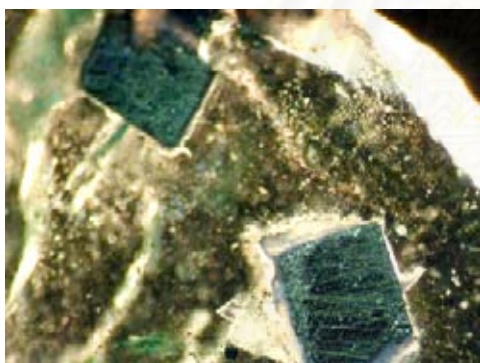


Figure 3.30: Numerous large pyrite crystals (partly cubes cut at surface) with many faces showing a typical yellowish metallic luster (STA5, 50x).

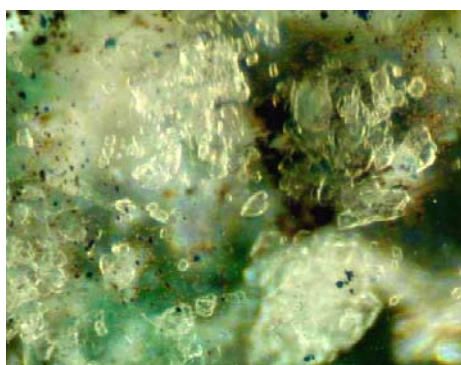


Figure 3.31: Colourless to brownish, transparent carbonate crystals with strong relief (STA10, 70x).



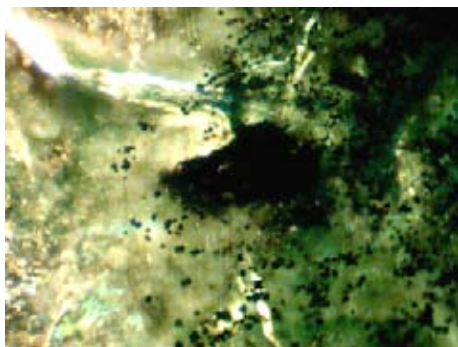


Figure 3.32: Chromite is present as black rounded crystals or in octahedrons. The large individual crystals are mostly isolated, and some form irregular clouds or films (STA1, 70x).

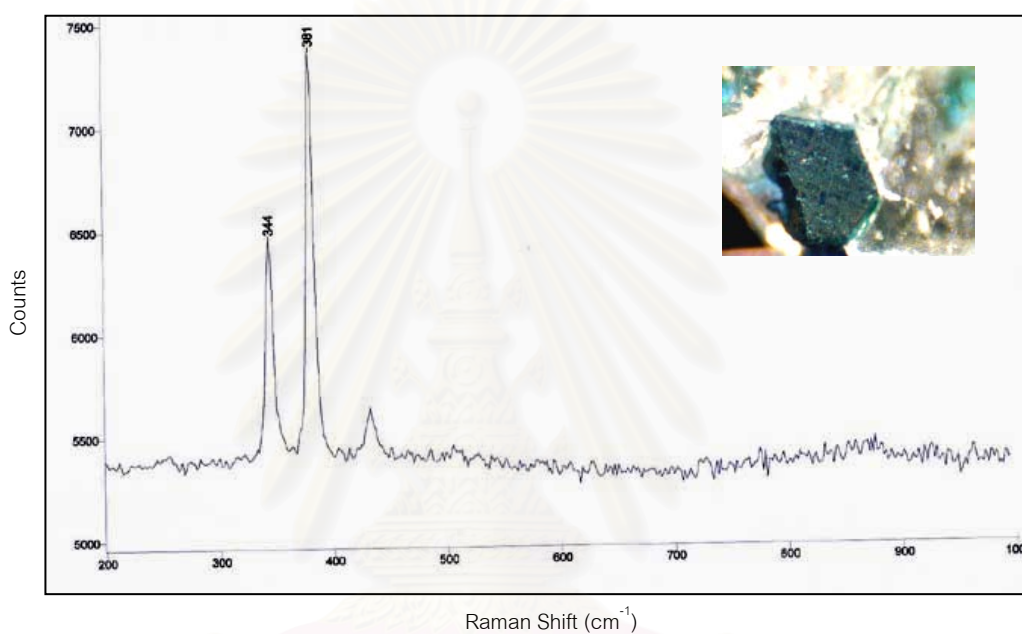


Figure 3.33: Raman spectrum of a pyrite inclusion which is common in emerald of the Santa Terezinha deposits (STA2; peaks at 344 and 381 cm<sup>-1</sup>).

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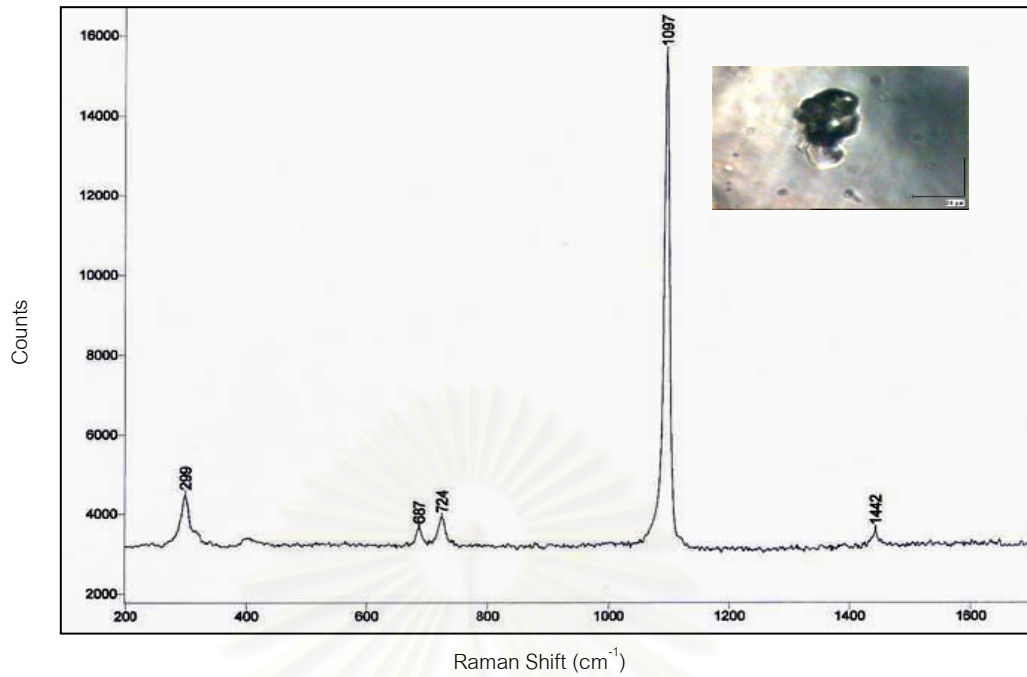


Figure 3.34: Raman spectrum of a whitish to grayish-transparent to translucent carbonate inclusions (strong relief) that may be calcite or dolomite (STA7; peaks at 724, 1097 and 1442 cm<sup>-1</sup>).

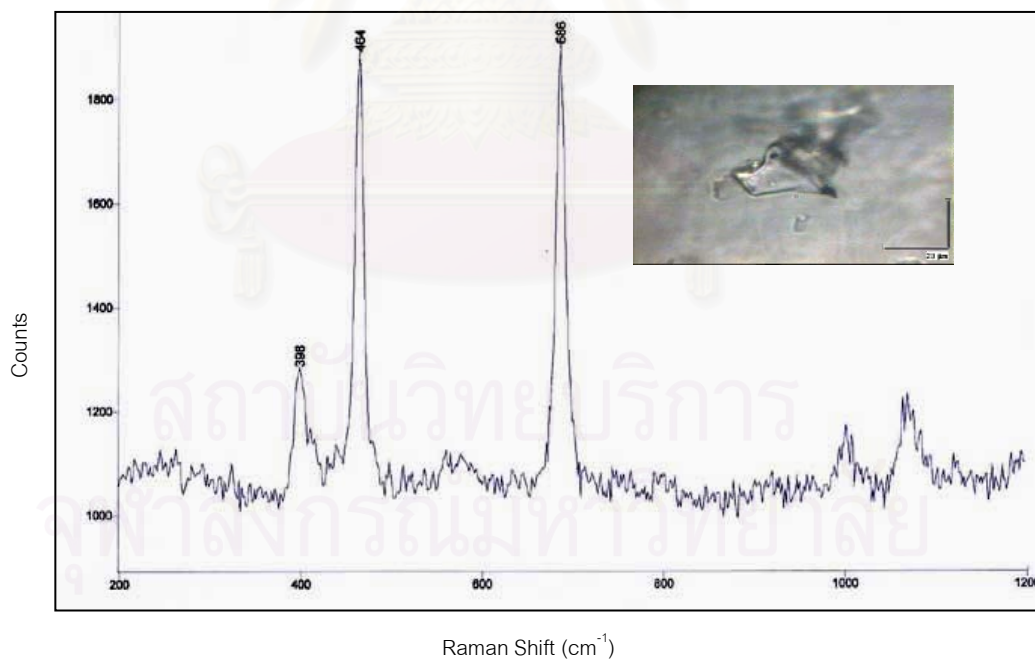


Figure 3.35: Raman spectrum of whitish transparent irregularly-shaped quartz inclusions (low relief); (STA10; peaks at 398 and 464 cm<sup>-1</sup>).

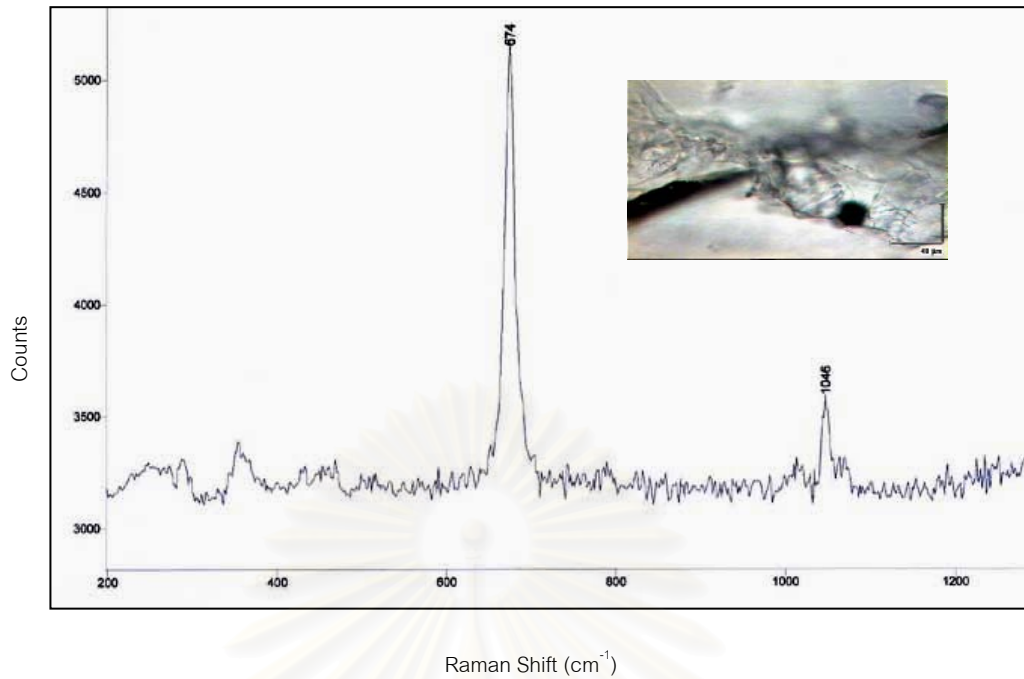


Figure 3.36: Raman spectrum of isolated colourless-transparent talc platelets. They show a faint relief and are randomly distributed in the host crystal (STA9; peaks at 674 and 1046  $\text{cm}^{-1}$ ).

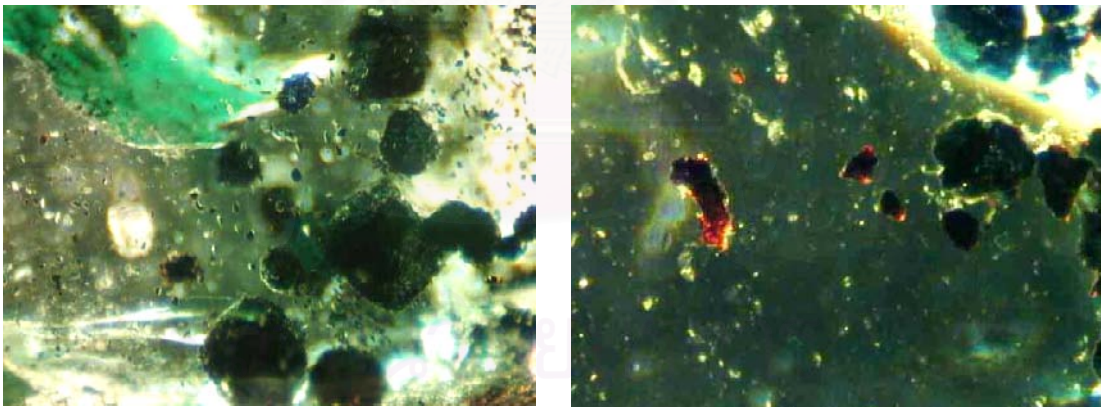


Figure 3.37: On the left, Numerous black-opaque spinel crystals, generally formed more or less compact aggregates (STA9, 70x) and on the right, small black grains of opaque Fe-Cr spinel, which are commonly dispersed along (healed) fissures (STA6, 70x).

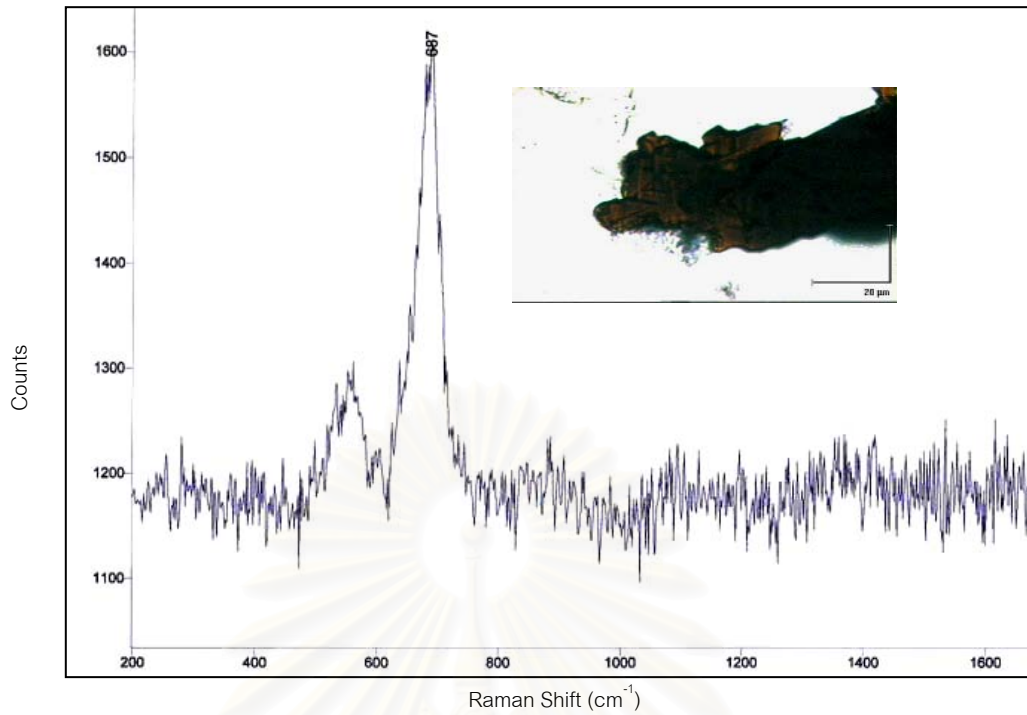


Figure 3.38: Raman spectrum of a compact aggregate crystal, possibly Cr-Fe spinel inclusions with iron stain (STA9; peaks at about 555 and 687 cm<sup>-1</sup>).

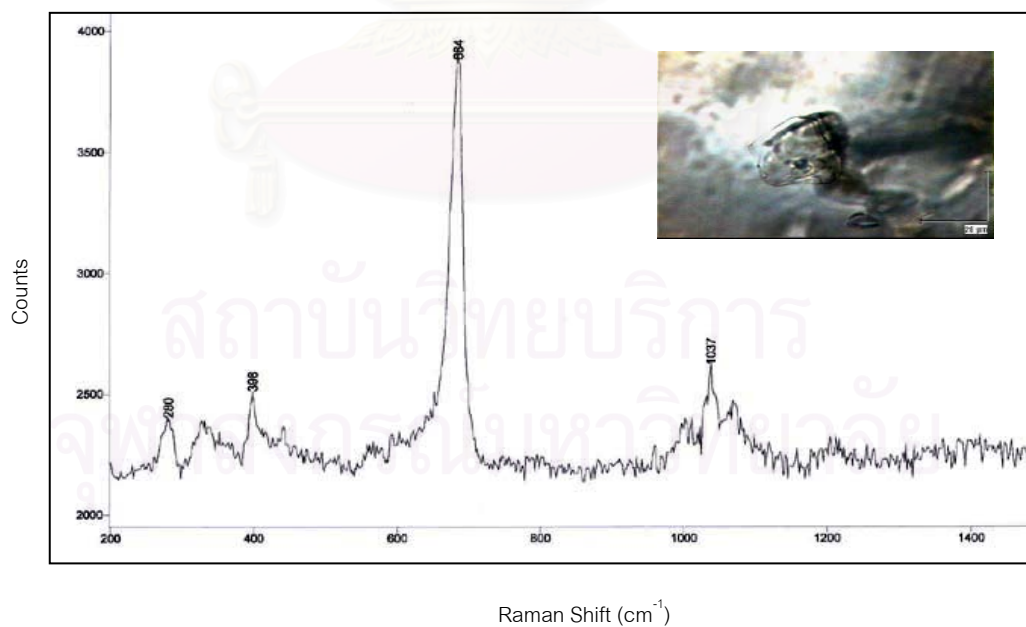


Figure 3.39: Raman spectrum of a light brownish transparent flake of mica, may be biotite or phlogopite inclusions in Brazilian emerald from Santa Terezinha de Goiás (STA5; peaks at 280, 359, 398 and 1037 cm<sup>-1</sup>).



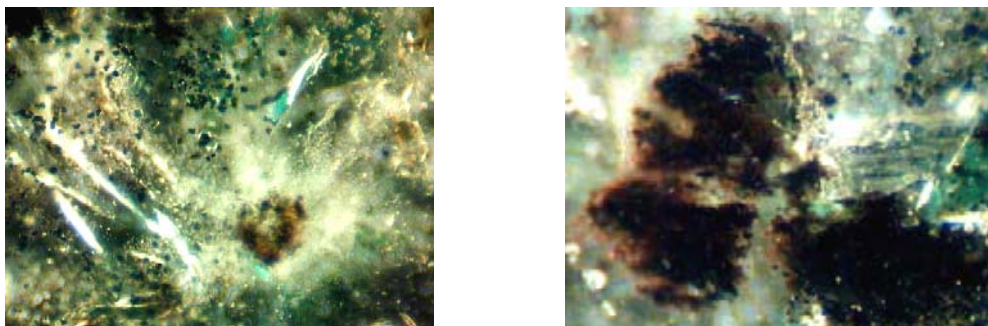


Figure 3.40 (Both pictures) Most of fluid inclusions are so small that they look like 'dust-particles' and secondary inclusions of iron stain in fissures and fractures (STA1, 3; 70x).

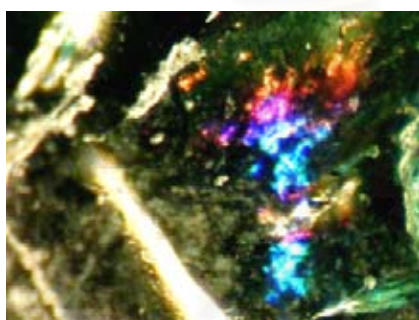


Figure 3.41 Fracture-filling with oil showing interference colour (STA3, 60x).

#### 3.4.4 Madagascar:

##### *Mananjary deposits of Madagascar*

The most common mineral inclusions in 30 emerald samples from the Mananjary deposits are quartz, which can be divided, similar to those observed by Schwarz (1994), in 4 categories on the basis of their morphology and manner of occurrences such as, 1) Transparent, colourless, elongated, prismatic crystals, often associated with primary fluid inclusions in Figures 3.42, 2) Rounded or spherical grains, not associated with fluid inclusions or 3) Irregularly rounded crystals, some of which had corroded rough surface; in Figures 3.43 and 4) Isolated crystals, which were baby crystals of former fluid-filled cavities can be shown with Raman spectrum in Figures 3.44.



Mica (biotite/phlogopite) inclusions normally can be observed from worldwide occurrences related to mica schists (Zambia, Madagascar and Brazil). In these samples, mica inclusions appear as dark brown transparent crystals which may have an hexagonal outline or prismatic shape (Figures 3.45 and 3.46 with the Raman spectrum). The amphibole inclusions can be seen in emeralds due to host rocks (mica schist) grade into amphibole schist. These mineral inclusions are green and show a long prismatic habit of actinolite which was identified by Laser Raman (Figure 3.47). Compact aggregates composed of amphibole and talc crystals are also observed in the border region of the sample (Figure 3.48).

Low relief well-developed rhombohedral carbonate inclusion is also noted (Figures 3.49) and in Figure 3.50 shows Raman spectra of daughter carbonate crystals in CO<sub>2</sub>-rich fluid inclusions. Large numbers of primary fluid inclusions occur as elongated cavities which look like 'rain effect' in Figures 3.51. Some fluid inclusions (negative crystals) show high relief in Figure 3.52. A sample in Figure 3.53 exhibits the healed fissures filled with foreign substances (may be oil) and some samples show coloured zones with primary fluid inclusions oriented perpendicular to c-axis (Figure 3.54)

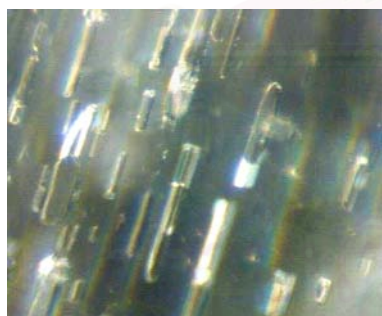


Figure 3.42: Elongated, prismatic quartz inclusions, often associated with primary fluid inclusions (MA-17, 100X).

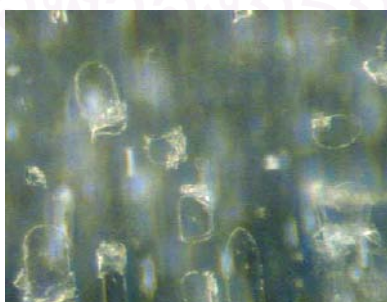


Figure 3.43: Colourless transparent, irregularly rounded quartz inclusions, some of which had corroded rough surface (MA-18, 90X).

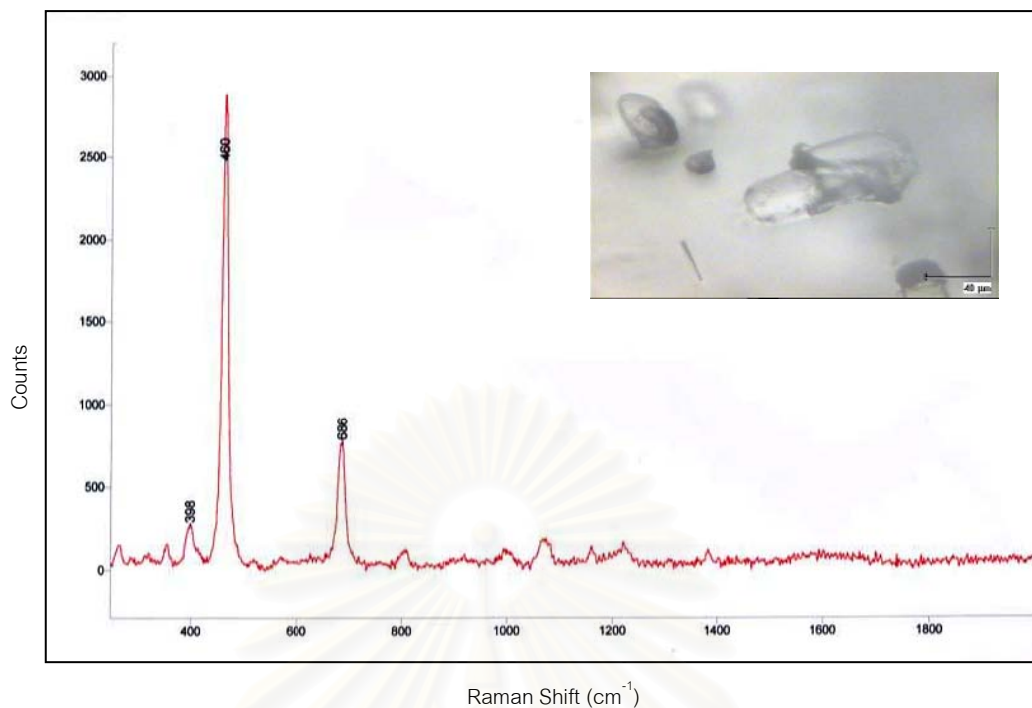


Figure 3.44: Raman spectrum of isolated elongated, prismatic colourless quartz inclusion, often associated with primary fluid inclusions (MA-11; peaks at 398 and 460 cm<sup>-1</sup>).

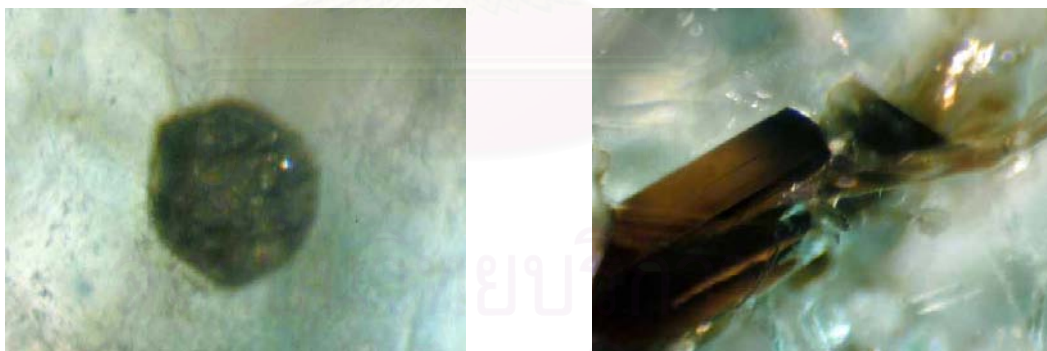


Figure 3.45: Perfect (pseudo-), hexagonal outline of a dark brown mica inclusion on the left (MA2, 120x) and dark brown elongate crystals or 'books' of mica inclusions on the right (MA4, 60X).

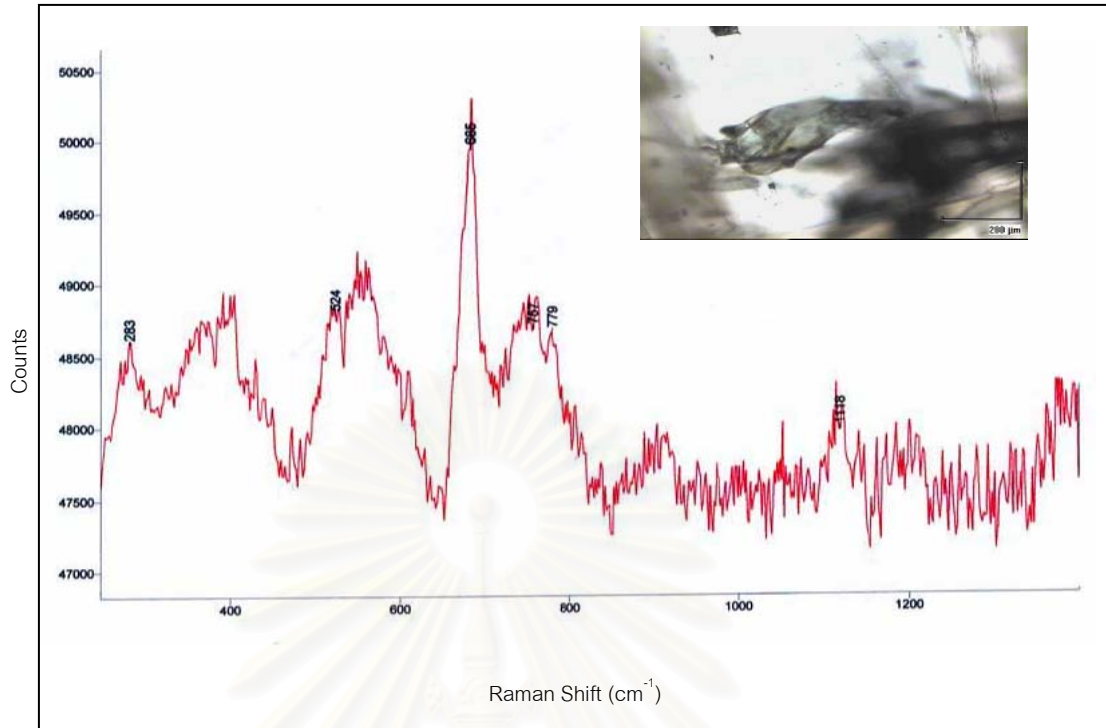


Figure 3.46: Raman spectrum of brown mica crystals are irregularly rounded and randomly orientated in the emerald host (MA1; peaks at 283, 406, 524, 557, 563, 767 and 779  $\text{cm}^{-1}$ ).

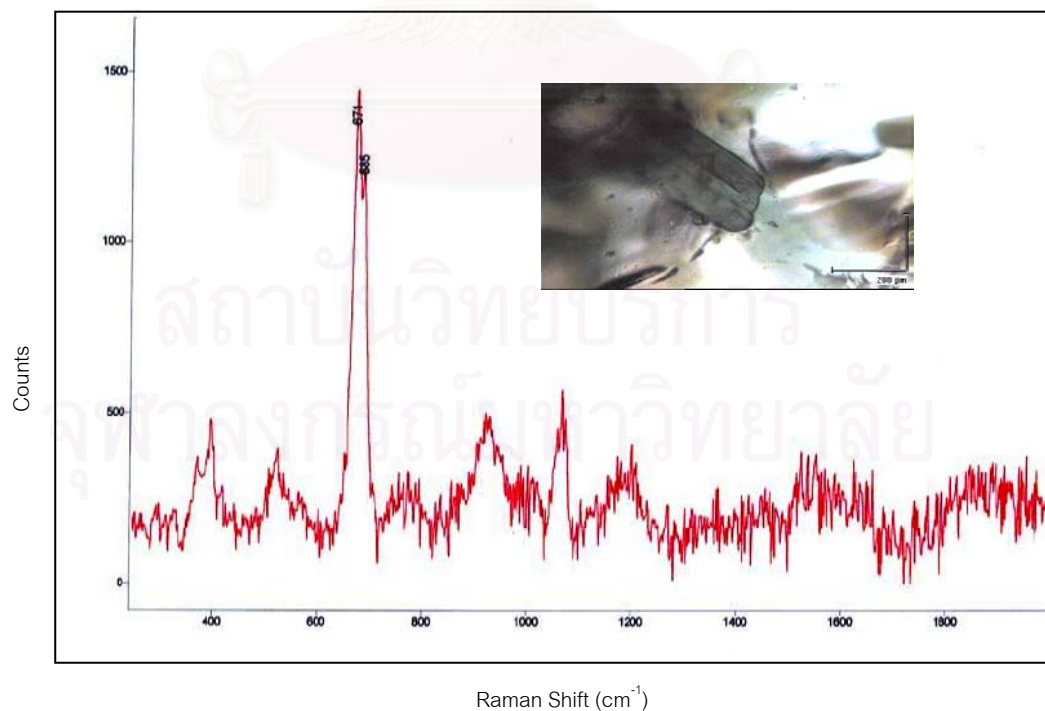


Figure 3.47: Raman spectrum of long prismatic, transparent green amphibole (actinolite) inclusions (MA3; peaks at 383 and 671  $\text{cm}^{-1}$ ).

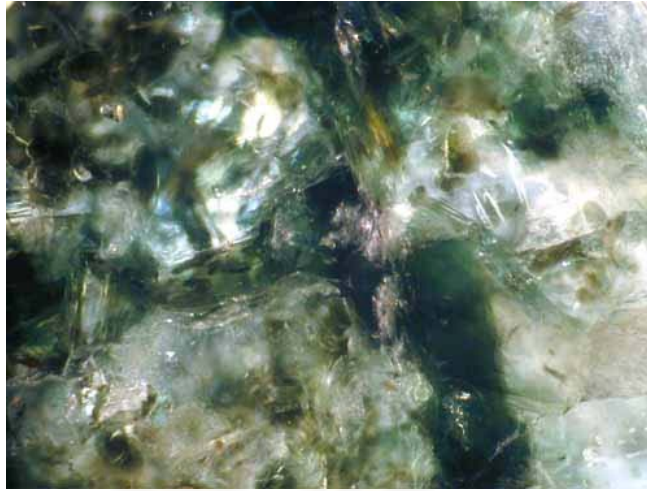


Figure 3.48: Compact aggregates composed of green amphibole (actinolite) and talc (colourless to whitish mineral inclusions) at the rim of the sample (MA3, 100x).

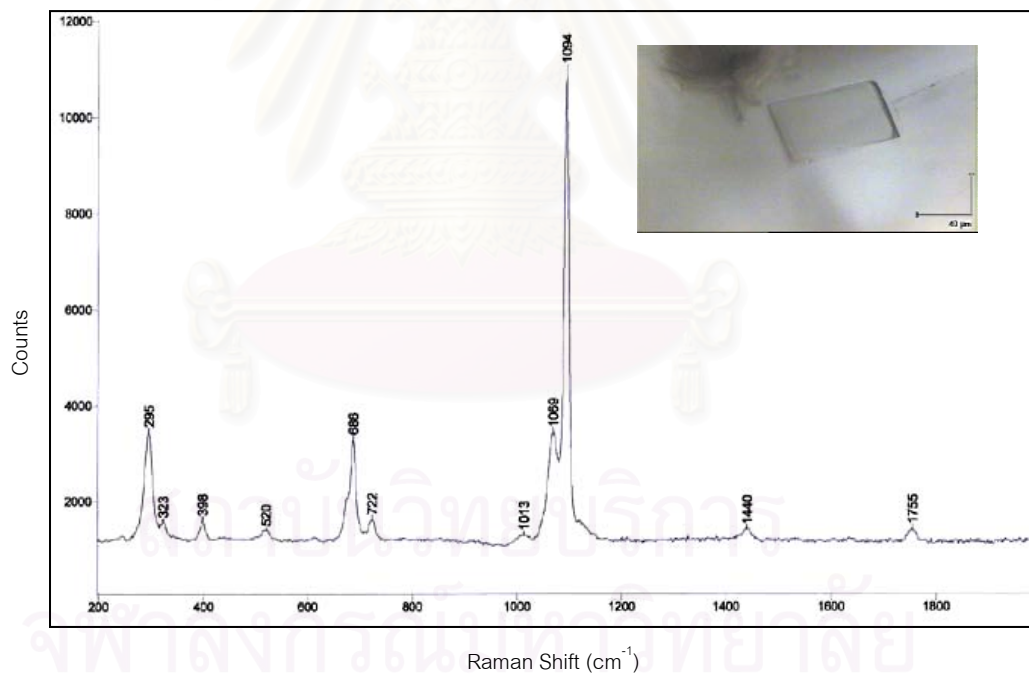


Figure 3.49: Raman spectrum of rhombohedral light brown carbonate crystal can be seen in some of the Mananjary emeralds (MA3; peaks at 295, 722, 1069, 1094, 1440 and 1755  $\text{cm}^{-1}$ ).



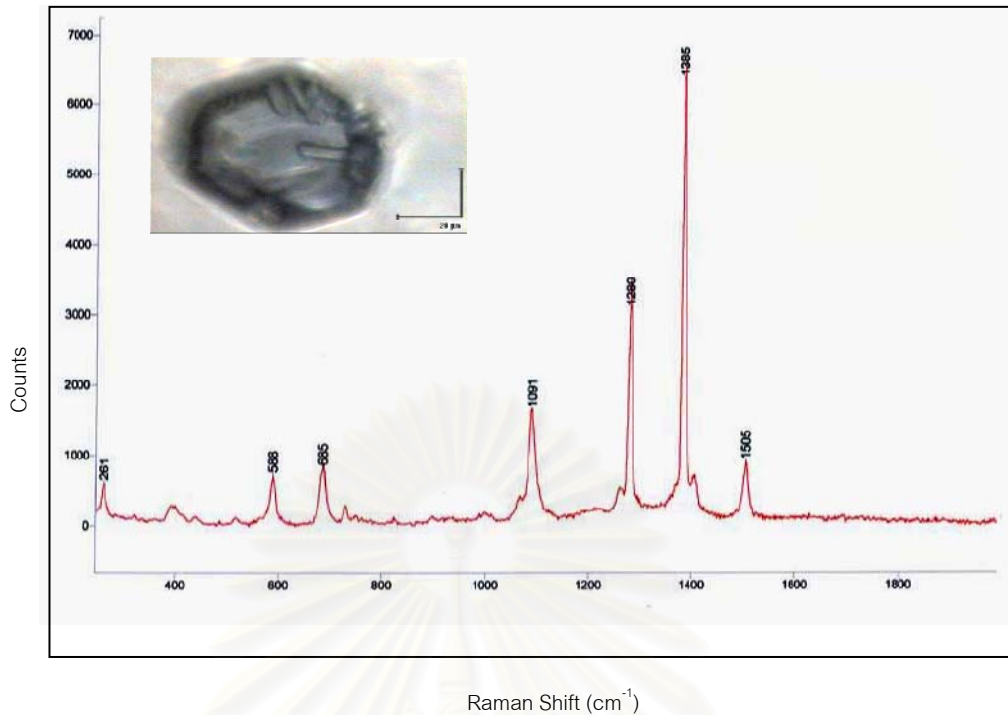


Figure 3.50: Raman spectra of several daughter carbonate crystals and CO<sub>2</sub> in negative crystal inclusions have been identified by Laser Raman spectroscopy in sample MA20 (carbonate peaks at 261, 1091 and 1505 cm<sup>-1</sup> and CO<sub>2</sub> at 1280 and 1385 cm<sup>-1</sup>).

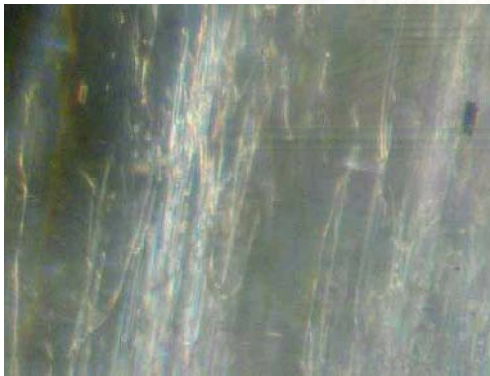


Figure 3.51: Some samples showing primary fluid inclusions occurred as elongated cavities, look like 'rain effect' (MA24, 90x).



Figure 3.52: Well developed rectangular outlines of high relief negative crystal (MA20, 120x).



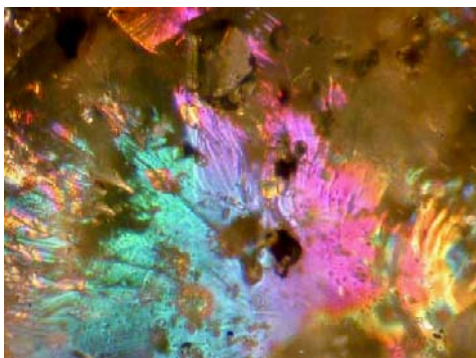


Figure 3.53: Oil-filling in fractures showing interference colour (MA20, 50x).

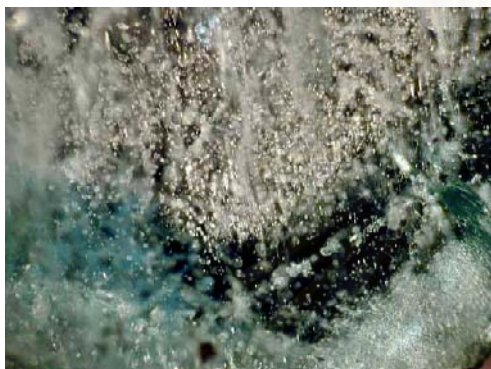


Figure 3.54: Syngenetic inclusions formed along growth and coloured zones, perpendicular with c-axis (MA8, 30x).

### 3.4.5 Zambia:

#### *Kafubu area in Ndola Rural district*

The most common mineral inclusions in 30 samples of Zambian emerald are micas, which include phlogopite/biotite platelets, euhedral and irregularly rounded transparent, brown crystals (Figures 3.55 and Figure 3.56 with the Raman spectrum). The next common mineral inclusions are green, elongated prismatic actinolite inclusions and their Raman spectra shown in Figures 3.57 and 3.58. Raman spectroscopy can identify the following daughter crystals in multi-phase fluid inclusions; colourless – whitish to grayish, transparent to translucent carbonate crystals (high relief) in multi-phase (11,2s,1g) inclusions (Figure 3.59), low relief, colourless-transparent quartz crystals and orange hematite platelets (Figures 3.60 and 3.61), and both hematite and quartz multi-phase (Figure 3.62). Other internal features are unhealed fissures showing a frosted appearance (Figure 3.63) and strongly developed growth structures and coloured zoning parallel to the emerald's prism faces (Figure 3.64). In Figure 3.65 shows

brown foreign substances filled in fissures exhibiting dendritic pattern. And unhealed fissures show mirror-like reflection effect which are, partly, relative planer to the run more or less parallel to the emerald's c-axis (Figure 3.66). Growth tubes containing iron stains and primary fluids are shown in Figure 3.67.

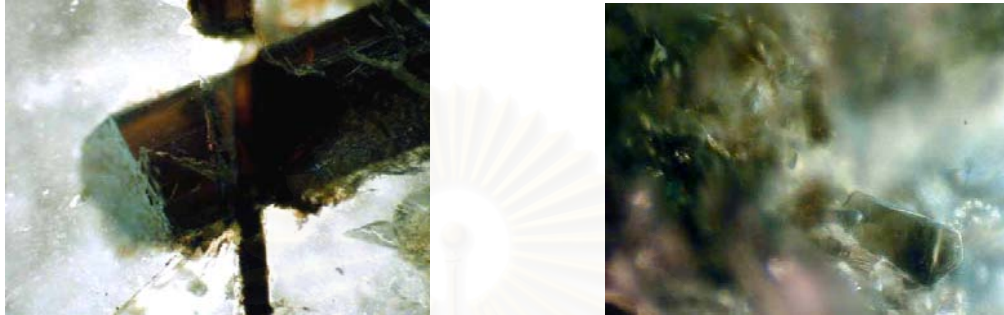


Figure 3.55: 'Booklets' of mica (left); (ZA13, 120x) and cluster of brown irregularly rounded mica platelets of biotite/phlogopite (right), which do not show any crystallographic orientation in the emerald host; (ZA30, 100x).

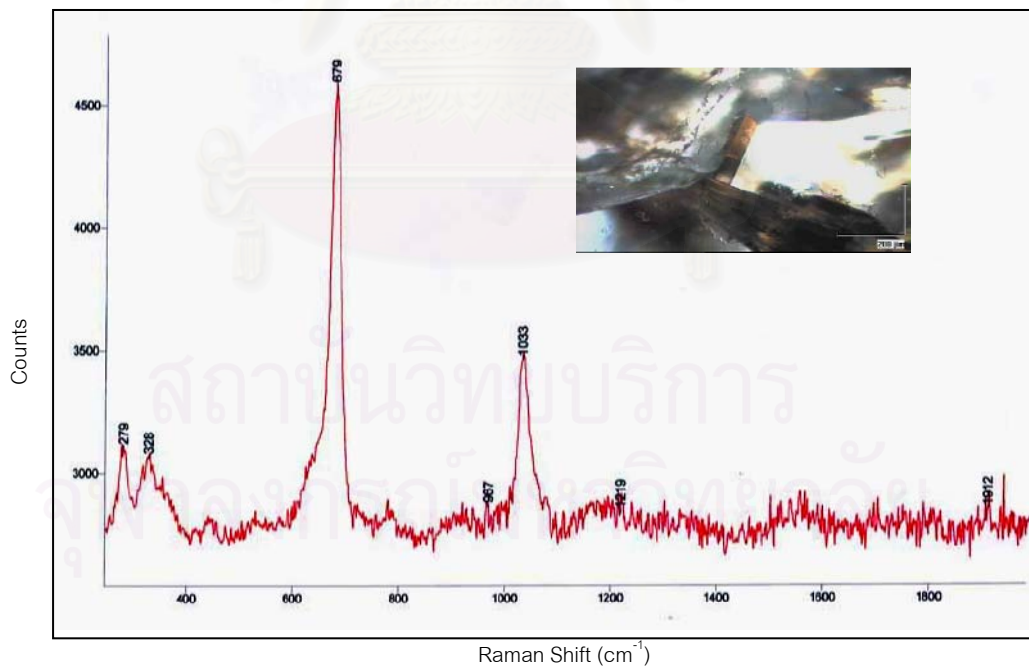


Figure 3.56: Raman spectrum of isolated mica platelets (ZA20; peaks at 279, 328 and 679 cm<sup>-1</sup>).

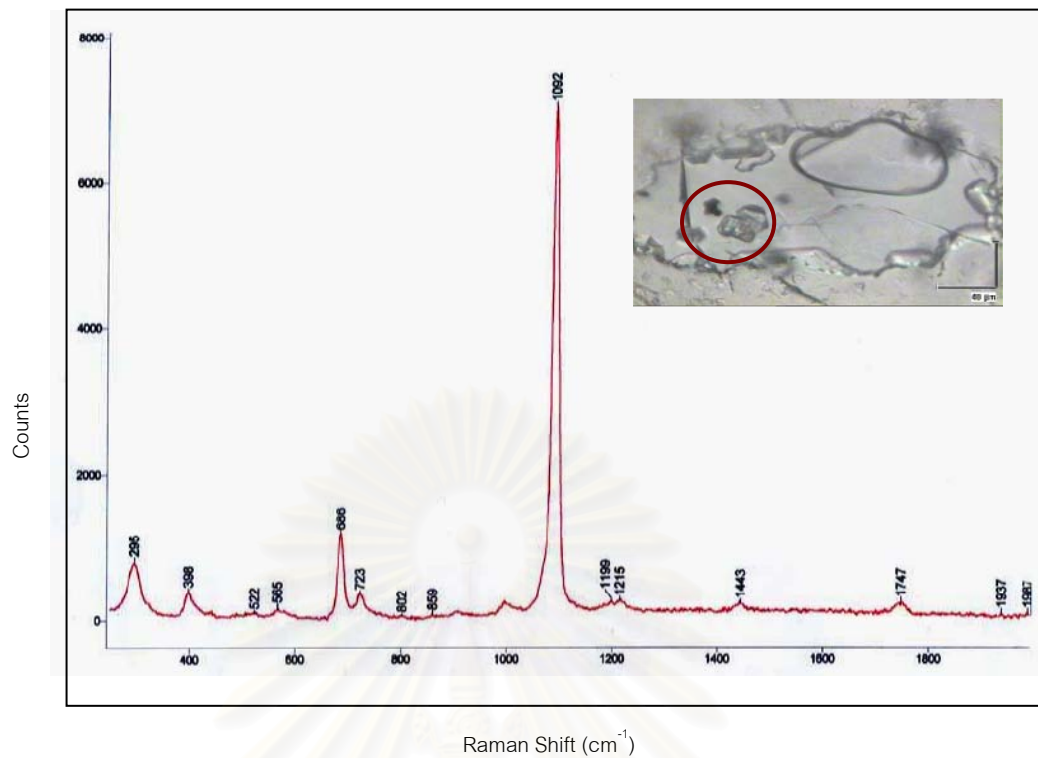


Figure 3.57: Raman spectrum of carbonate crystals (high relief) in a multi-phase (1,2s,1g) inclusion (ZA23; peaks at 565, 723, 802 and 1092 cm<sup>-1</sup>).



Figure 3.58: elongated prismatic actinolite inclusions in both pictures (high relief); (ZA17, 120x).

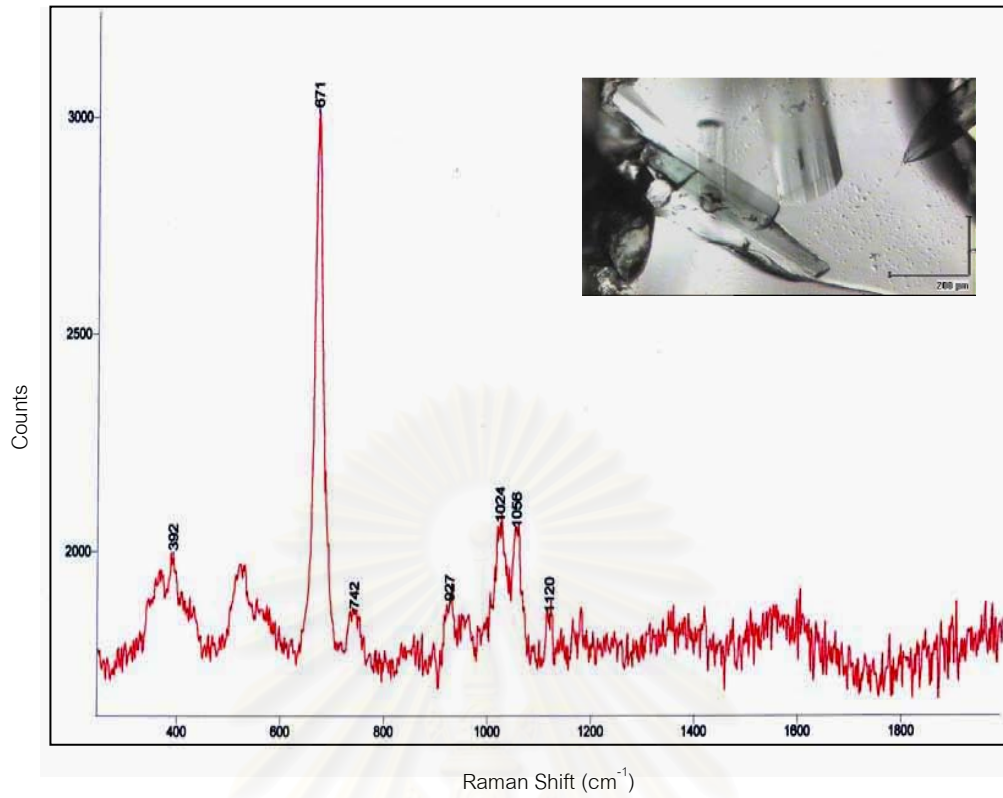


Figure 3.59: Raman spectrum of actinolite inclusions in Zambian emerald (ZA17; peaks at 383, 671, 742 and 1058  $\text{cm}^{-1}$ ).

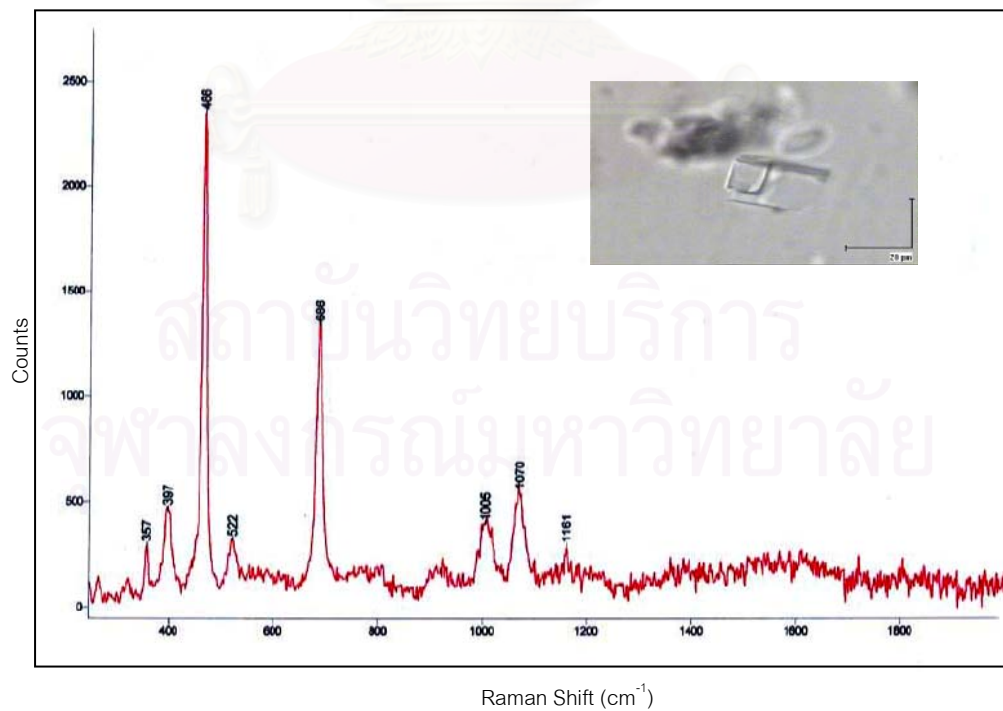


Figure 3.60: Raman spectrum of transparent, colourless quartz inclusions in low relief (ZA25; peaks at 357, 397, 466 and 1161  $\text{cm}^{-1}$ ).



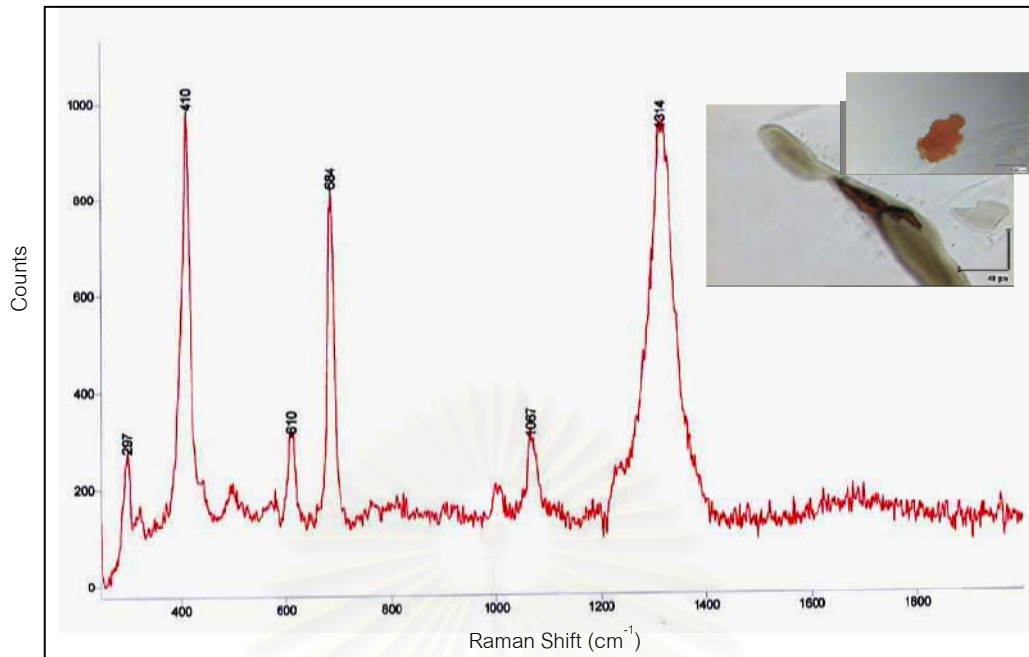


Figure 3.61: Raman spectra of mica altered to hematite (bottom) in Za21; (mica showing peaks at 297 and 1067  $\text{cm}^{-1}$  and hematite showing peaks at 410, 610 and 1314  $\text{cm}^{-1}$ ), high relief of orange hematite platelets (top).

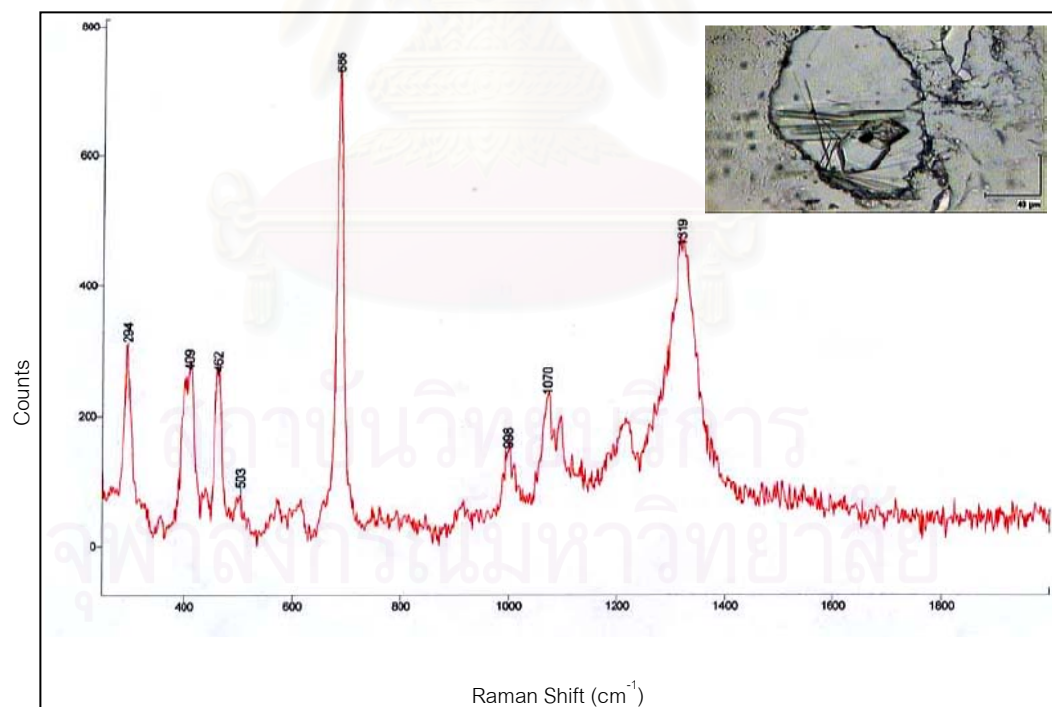


Figure 3.62: Raman spectra of sample ZA23 showing a combination of hematite (peaks at 294, 409, 503, 1070 and 1319  $\text{cm}^{-1}$ ) and quartz (peaks at 462  $\text{cm}^{-1}$ ) in a multi-phase inclusion.



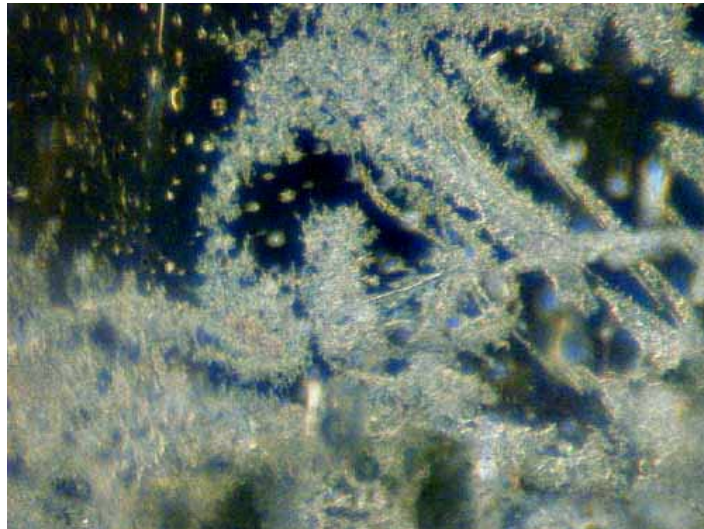


Figure 3.63: Unhealed fissures showing a frosted appearance (ZA22, 120x).

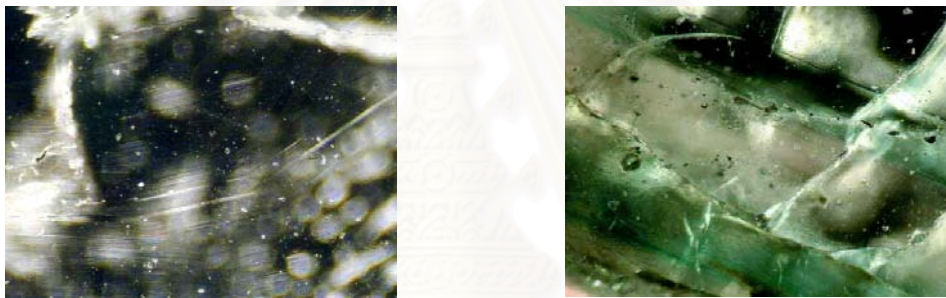


Figure 3.64: Strongly developed growth structures (ZA4, 70x) and coloured zoning parallel to the emerald's prism faces (ZA23, 70x).

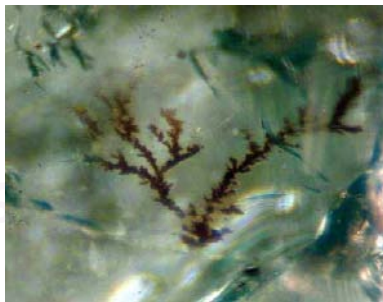


Figure 3.65: Unhealed fissures, partly with foreign material (ZA5, 100x).



Figure 3.66: Reflection from unhealed fissures showing mirror-like effect (ZA1, 70x).

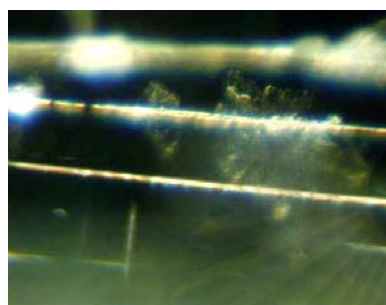


Figure 3.67: Growth tubes containing iron stain and fluids (left); (ZA29, 120x) and primary fluid inclusions (2-phase) showing a high relief gas bubble (right); (ZA29, 140x).

#### 3.4.6 Canaiba-Socoto Deposits in Bahia State in Brazil:

The most common mineral inclusions found in 10 emerald samples from the Canaiba-Scototo deposits of Bahia State were brown irregular-shaped mica platelets and greenish-brown chlorite, which were randomly distributed in the emerald host crystals (see Figures 3.68 and 3.69). Raman spectra confirmed a mica inclusion of probably biotitic or phlogopitic composition (Figures 3.70) and a chlorite inclusion (Figure 3.71). Growth phenomena (internal growth lines and coloured zoning) can be observed by looking in the direction of c-axis (Figure 3.72). Iron stain and thin channel-like growth were found in some samples (Figure 3.73). Most of fluid inclusions are very small, that look-like 'dust-particles' and other 'dust-particles' form 'flake-like' agglomerations (Figure 3.74).

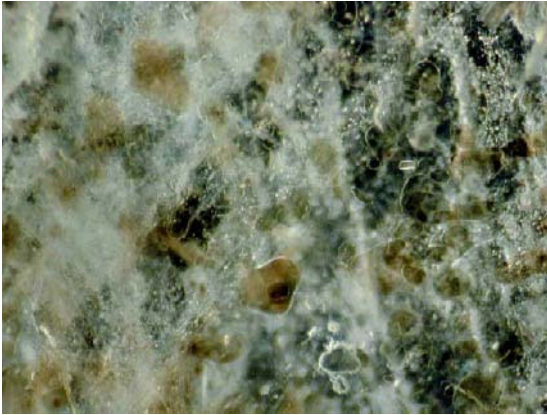


Figure 3.68: Numerous brown booklets and crystals of biotite inclusions (CAR6, 70x).

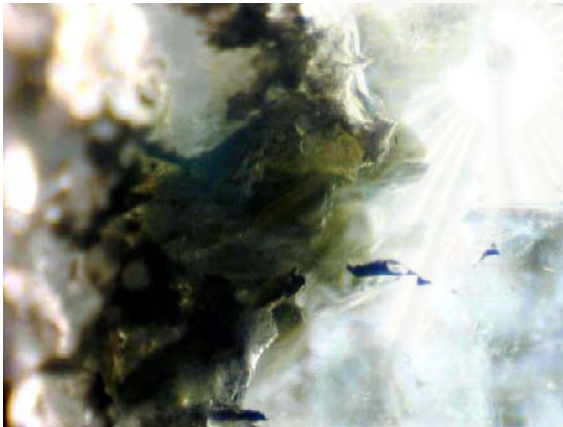


Figure 3.69: Typical chlorite inclusions showing irregularly shaped and greenish brown crystals (CAR8, 70x).

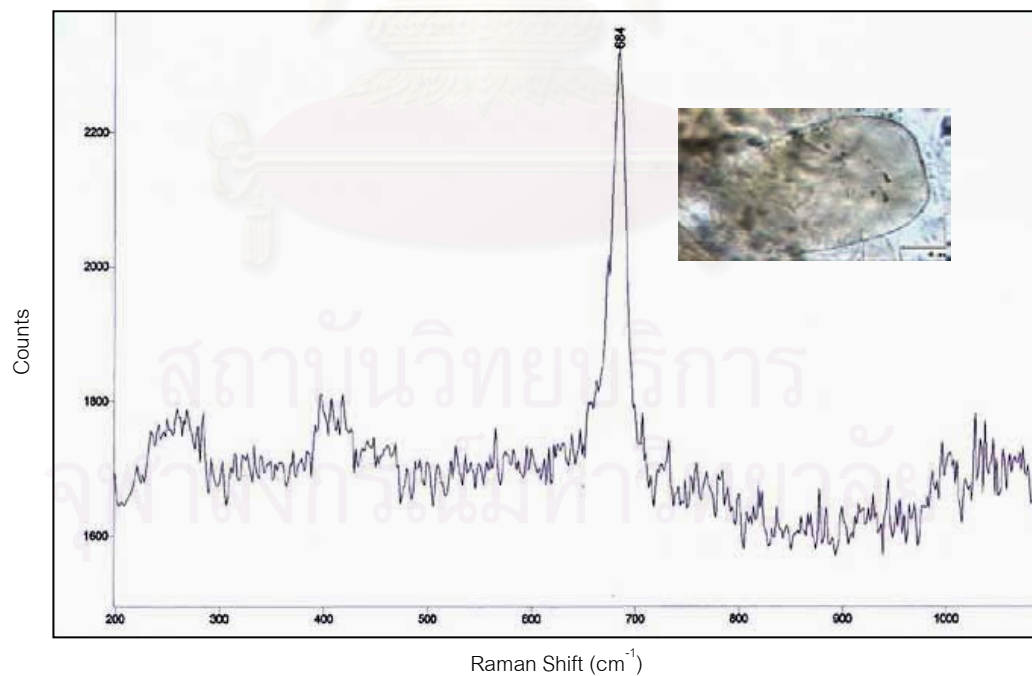


Figure 3.70: Raman spectrum of a biotite inclusion in a Brazilian emerald from the Caniaba-Socoto deposits (CAR6; biotite peaks at 406, 684  $\text{cm}^{-1}$ ).



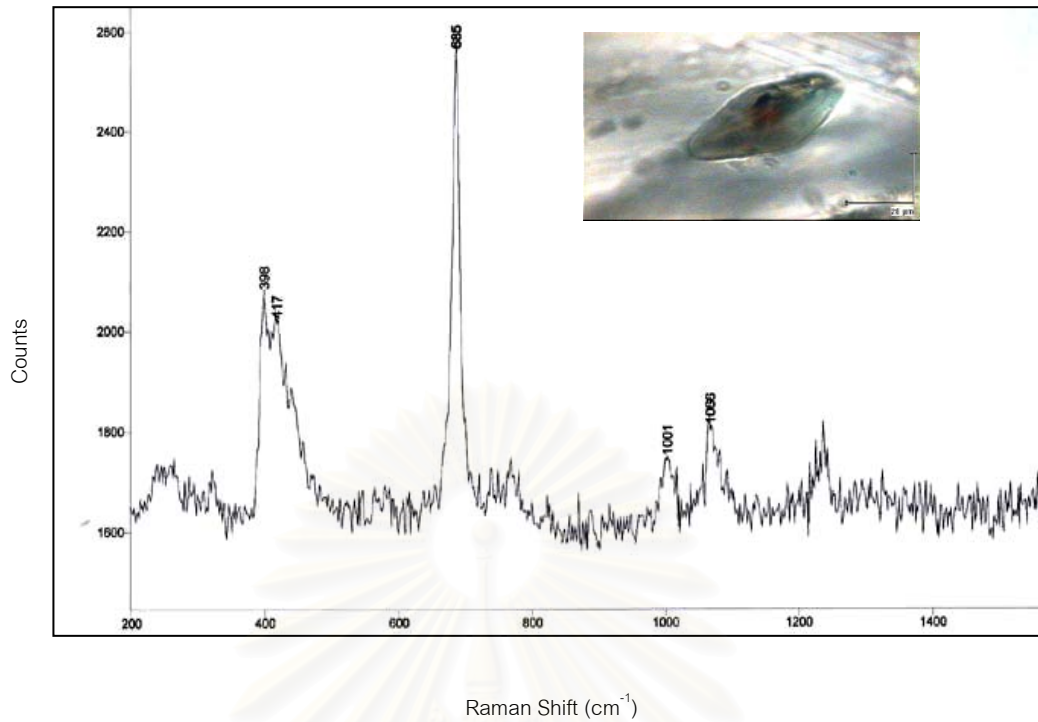


Figure 3.71: Raman spectrum of a chlorite inclusion associated with the biotite inclusion in a Brazilian emerald from Caniaba-Socoto deposits (CAR10; chlorite peaks at 398, 685, 1011 and 1088  $\text{cm}^{-1}$ ).

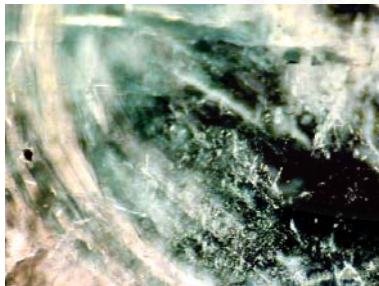


Figure 3.72: Well developed growth structures and coloured zoning (parallel to the emerald's prism face) (CAR5, 60x).

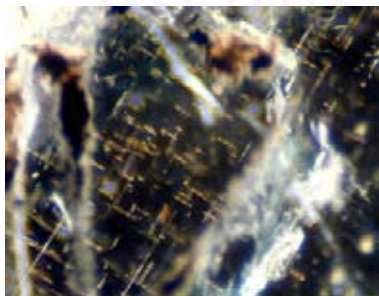


Figure 3.73: Small fluid inclusions look like 'dust-particles' and some of these healed with iron stain (CAR9, 70x).

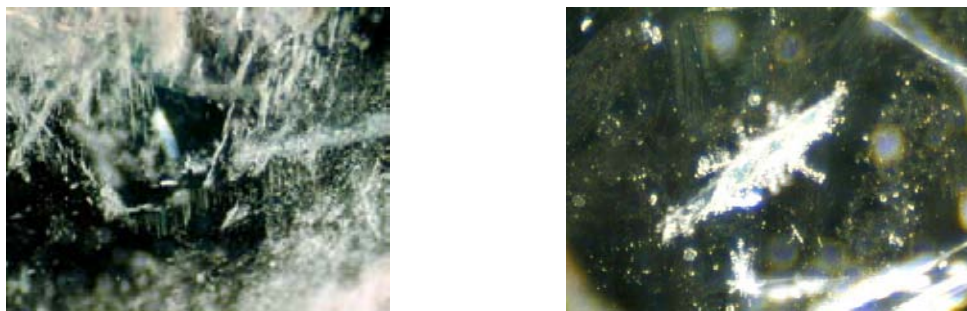


Figure 3.74: On the left, 'particles' showing sugar grain like' structures (CAR4, 70x) and on the right healed fissures have a quite compact aspect, other 'dust-particles' form 'flake-like' agglomerations (CAR7, 70x).

#### 3.4.7 Itabira deposits in Minas Gerais State of Brazil:

Under the optical microscope, most inclusions of 10 emerald samples from the Belmont mine at Itabira deposit of Minas Gerais state are biotite micas, on one hand, exhibit deep brown colour and thick translucent to opaque platelets (Figure 3.75). Mica inclusions, on the other hand, are usually strongly round and Raman spectra of these mica inclusions (biotite and phlogopite) are shown in Figure 3.76. Quartz inclusions were found as isolated, colourless-transparent and irregularly-shaped crystals with weak relief as shown in Figure 3.77 together with its Raman spectrum. A peak of carbon dioxide, within negative crystal is shown by Raman spectra (figure 3.78). There are numerous primary fluid inclusions in the form of growth tubes. These fluid inclusions are partly associated with quartz crystals dispersed over fissure planes (Figure 3.79). Two and three phase fluid inclusions were commonly found in short barrel-shaped primary negative crystals (Figure 3.80). In general, tiny growth tubes (primary cavities/fluid inclusions) and very small fluid inclusions can be observed in some samples (Figure 3.81).



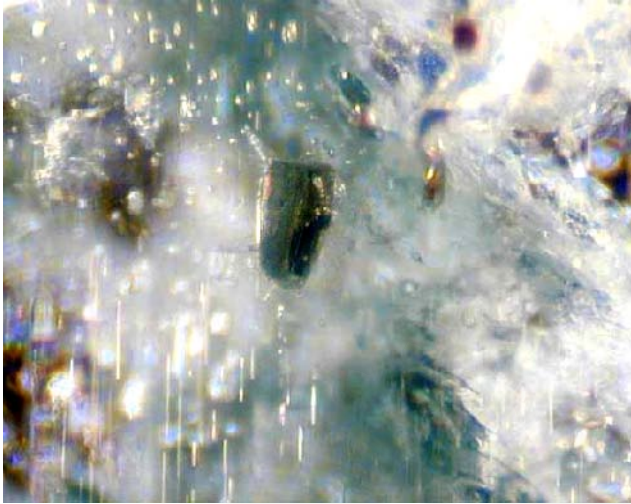


Figure 3.75: High relief dark brown biotite inclusions which are randomly distributed in the host crystal (ITA10, 60x).

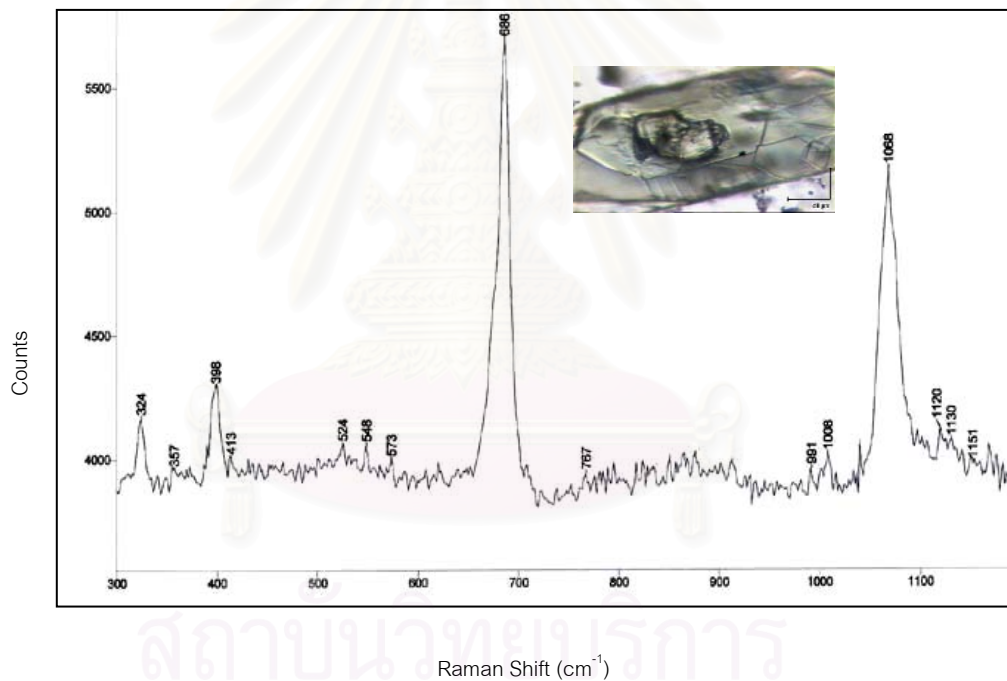


Figure 3.76: Raman spectrum of transparent flakes of micas may be biotite/phlogopite inclusions in Brazilian emerald from Itabira (ITA1; peaks at 357, 688, 413, 548, 787 and 1088 cm<sup>-1</sup>).

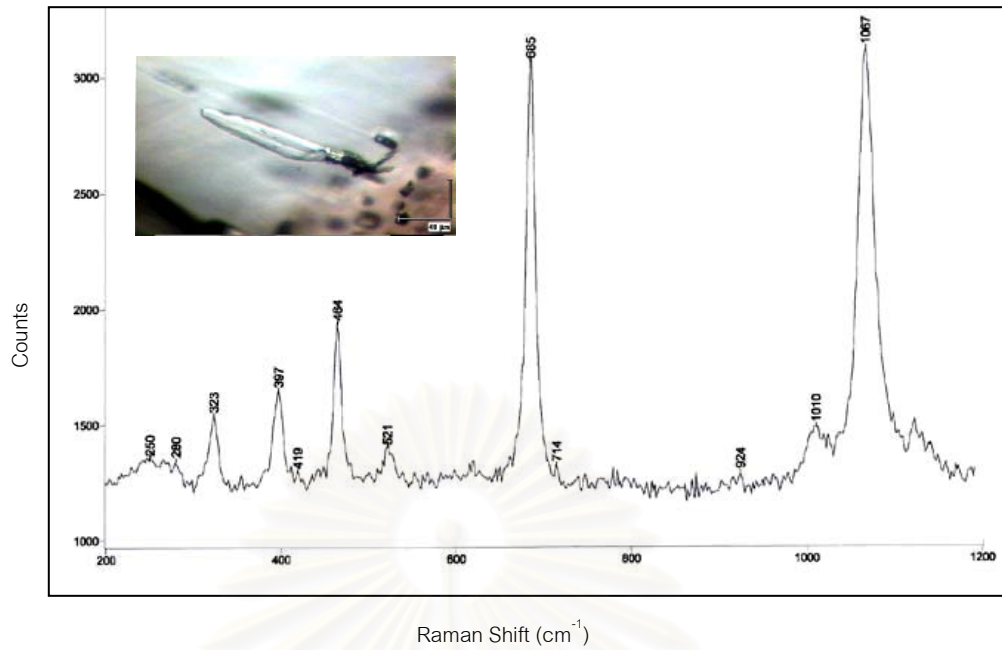


Figure 3.77: Raman spectrum of a colourless-transparent irregularly-shaped quartz inclusion (ITA10; peaks at  $464 \text{ cm}^{-1}$ ).

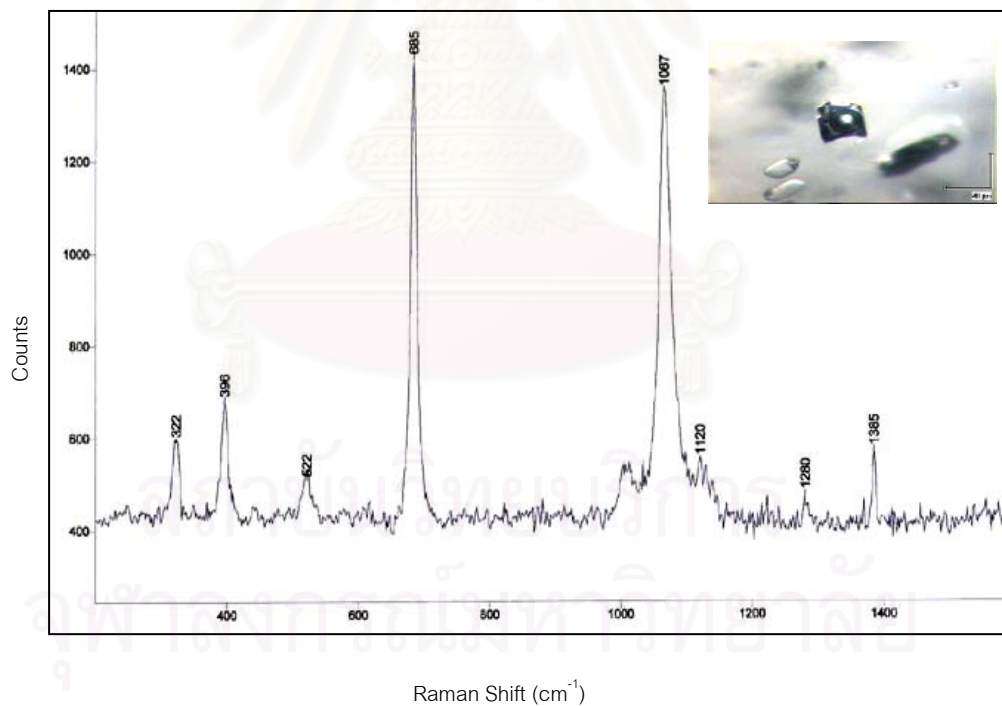


Figure 3.78: Raman spectrum of carbon dioxide within the negative crystal inclusions (ITA8; peaks at  $1280$  and  $1385 \text{ cm}^{-1}$ ).

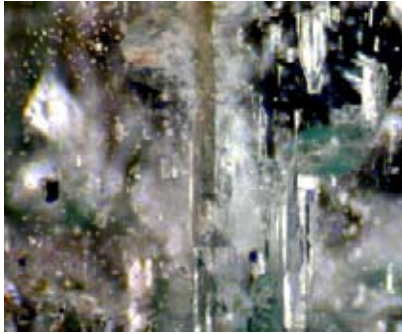


Figure 3.79: Elongated Growth tubes filled with fluids and the associated quartz inclusions (ITA3, 70x).

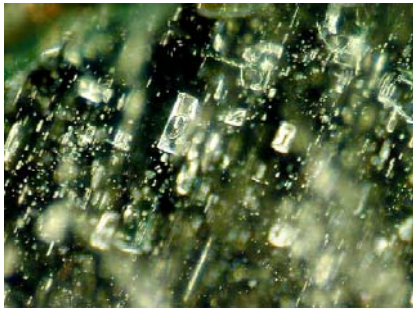


Figure 3.80: Barrel-shaped fluid inclusions (ITA1, 70x).

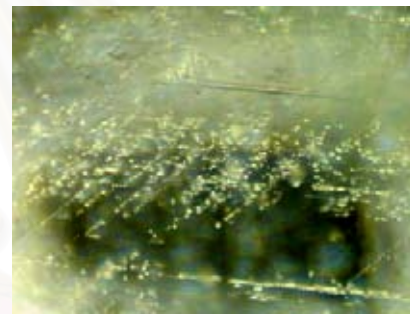
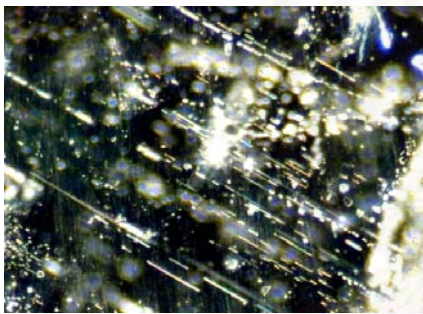


Figure 3.81: On the left; tiny growth tubes filled with fluids (ITA9, 70x) and on the right; some fluid inclusions having 'dust-like' appearance (ITA7, 70x).

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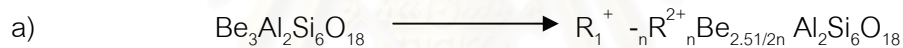
## CHAPTER IV

### CHEMICAL ANALYSIS

#### 4.1 Introduction

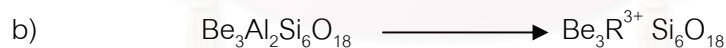
The major components of a natural beryl are  $\text{BeO}$ ,  $\text{Al}_2\text{O}_3$  and  $\text{SiO}_2$  which are normally accompanied by minor alkali elements such as Na, Mg, Cs,... (Bosshart, 1991). Other important chromophoric elements are Cr, V, Fe. In addition, water molecules are commonly found in channels of ring of cyclosilicate structure of emerald (see Figure 3.3)

The ideal formula of a beryl is  $\text{Be}_3\text{Al}_2\text{Si}_6\text{O}_{18}$ , in which alkali, water and other trace elements can substitute into emerald/beryl lattice or structural channels, as follows (Feklichev et al., 1963 in (a) and Schaller et al., 1962 in (b) cited in Sinkankas, 1989):



where:  $\text{R}^+ = \text{Cs}^+, \text{Li}^+, \text{Rb}^+, \text{Na}^+, \text{K}^+$

where:  $\text{R}^{2+} = \text{Ca}^{2+}, \text{Ba}^{2+}, \text{Sr}^{2+}$ ; and  $n = 0, 1$



where:  $\text{R}^{3+} = \text{Al}^{3+}, \text{Fe}^{3+}, \text{Cr}^{3+}, \text{Sc}^{3+}$

The fact that many cations can substitute into the emerald structure, it is expected that emeralds from different geological environments should contain different amounts of trace elements in their structures. Hence trace element contents should be applicable as a chemical fingerprint of emeralds from various geologic occurrences. It is therefore the aim of this study to look into the trace element concentration of emeralds from various locations.



## 4.2 Samples and Procedure

In this study, the chemical analysis of emerald was carried out by a Laser ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICP-MS). This technique is suitable for analysis of trace elements in particular the light elements such as Li, Be, B, Na, Mg, despite the fact that it is a “destructive” technique.

The LA-ICP-MS unit at the Macquarie University in Australia was used in this study (Figure 4.1). Detailed descriptions of LA-ICP-MS instrumental analytical and calibration procedures are similar to those given by Norman et al. (1996). The UV laser ablation microprobe (a New Wave Research 213 nm Nb:YAG) is coupled to an Agilent 7500S ICP-MS. All analyses were done with a pulse rate of 5 Hz and a beam energy of approximately 0.5 mJ per pulse, producing a spatial resolution 50 micrometers. The beam is used to vaporize an extremely small sample of material (in this case, emerald). The ablated sample is carried by a stream of inert gas, usually argon, into a high-temperature field, causing dissociation of molecules and ionization of atoms.

The MS identifies and quantifies elements in terms of mass and charge. Some 41 elements and their relative amounts can be detected even when present in only a few parts per billion. Quantitative results for 16 trace elements (Li, Be, B, Na, Mg, Al, Sr, P, K, Ca, Ti, V, Cr, Mn, Fe and Cs of the emerald samples) were obtained through calibration of relative element sensitivities using the NIST-610 multi-element glass standard and an emerald with 12.5 wt% BeO as internal standards. The BCR2G basaltic glass standard was also used as an external standard. The detection limits vary from analysis to analysis and are typically less than 1 ppm for Li, Be, V and Cs; less than 4 ppm for Mg, Al, Ti, Mn and Cr; less than 13 ppm for B and Na; less than 80 ppm for P and Fe; less than 300 ppm for Si and K and less than 500 ppm for Ca.

In this study a total of 23 emerald samples from 7 different geological occurrences (2-4 representative samples from each locality) were selected for chemical analysis. In each sample at least 5 spots were analyzed across the polished surface normally cut perpendicular to c axis. The detailed positions of the spots analyzed are presented in Figures 4.2 to 4.8.



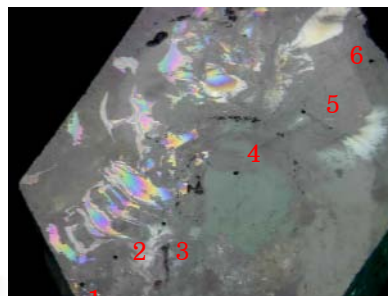
Figure 4.1: Laser ablation Inductively Coupled Plasma Mass Spectrometry Laser Ablation (LA-ICPMS) unit at Macquarie University, Australia used for chemical analysis in this study.

The representative laser points of 23 samples from different geological occurrences in below:

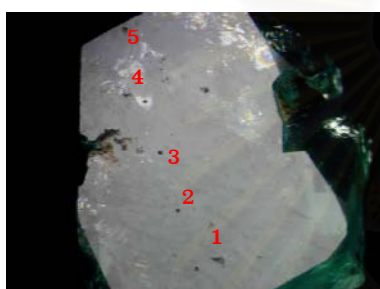
NI 1



NI 11



NI 16



NI 27

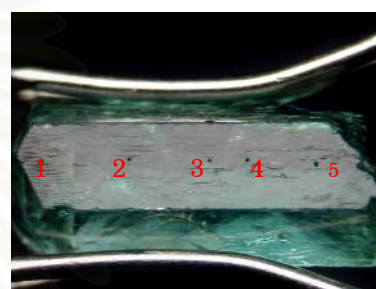
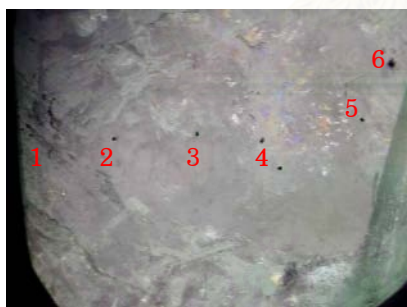
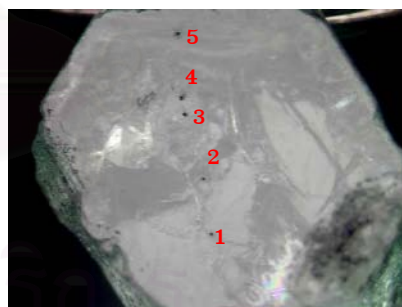


Figure 4.2: Showing the 4 samples of emerald from Kaduna Plateau, Nigeria.

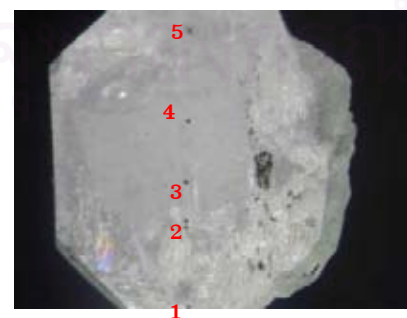
CO 1



CO 11



CO 18

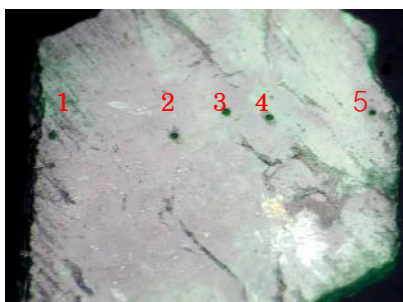


CO 22

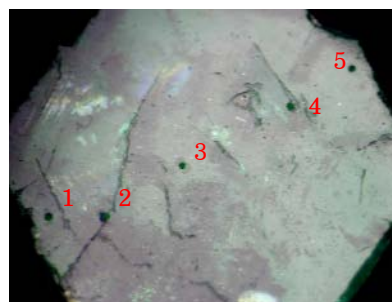


Figure 4.3: Showing the 4 samples of emerald from Cordillera Oriental, Colombia.

STA 1



STA 3



STA 8

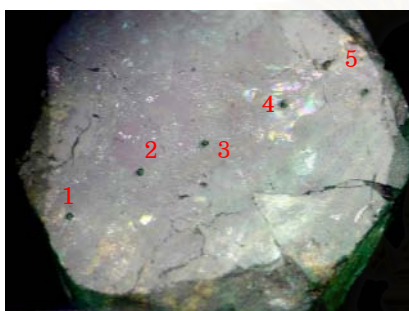
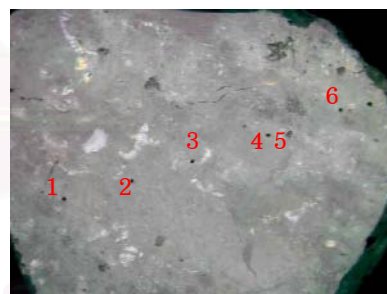


Figure 4.4: Showing the 3 samples of emerald from Santa Terezinha de Goais, Brazil.

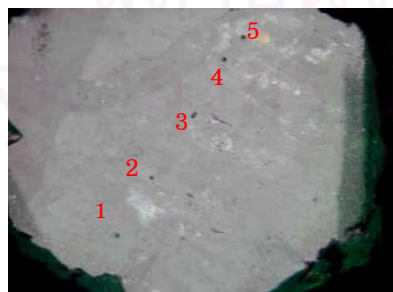
MA 4



MA 7



MA 8



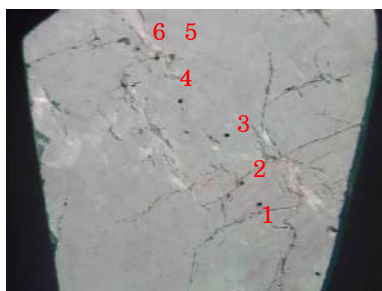
MA 18



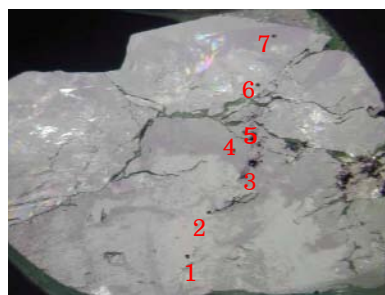
Figure 4.5: Showing the 4 samples of emerald from Mananjary region, Madagascar.



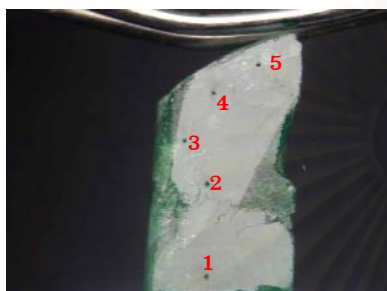
ZA 5



ZA 6



ZA 23

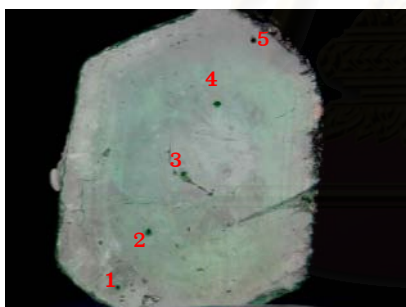


ZA 25



Figure 4.6: Showing the 4 samples of emerald from Ndola Rural district, Zambia.

CAR 4



CAR 7

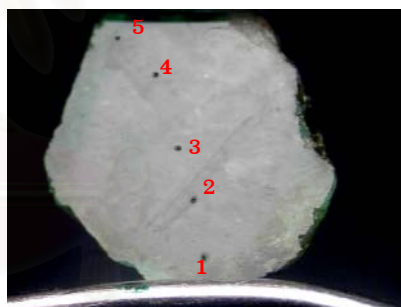
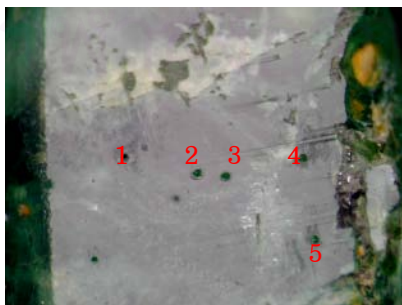


Figure 4.7: Showing the 2 samples of emerald from Carnaiba/Socoto, Bahia State, Brazil.

ITA 2



ITA 10

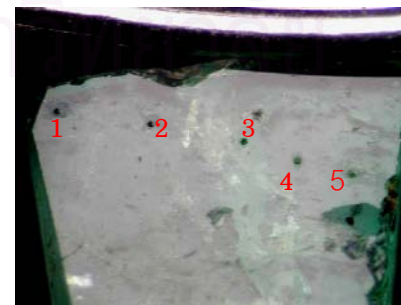


Figure 4.8: Showing the 2 samples of emerald from Itabira, Minas Gerias state in Brazil.

### 4.3 Results

The average concentrations of 16 trace elements; Li, Be, B, Na, Mg, Al, Si, P, K, Ca, Ti, V, Cr, Mn, Fe and Cs of each emerald sample analyzed in this study after recalculation into atom mole ppm are listed in Tables 4.1 to 4.3. The analytical concentrations in ppm by weight of original data, the plot of averages ( $\bar{X}$ ) and standard deviation (STD) are shown in Appendix IV

From the chemical data shown Tables 4.1 to 4.3, the elements of emeralds from this study are mainly as alkalis and iron. Amount of these elements are characteristic, which related to geological environments.

### 4.4 Cross plots

Cross plots of Na versus Fe (or Na vs Fe), Na vs Li, Na vs Mg, Na vs Cs, Na vs Ti, Na vs B, Na vs (Mg+Fe), Li vs K, Li vs Fe, Li vs Cs and Al vs sum of all monovalent+divalent+trivalent ions contents of all the emeralds in this study are shown in Figures 4.9 to 4.19. The majority of the plots include alkali elements because of the fact that they are the major trace elements which have the strongest influence on the optical, physical and chemical properties of emeralds.

A good correlation between Na and Mg is seen in Figures 4.11 and between Na and (Mg+Fe) in Figure 4.15 regardless of localities. The remaining cross plots seem to show no correlation of other pair of elements. With regard to Figures 4.11 and 4.15, the contents of Na and Mg and probably Fe are highest in emeralds from Santa Terezinha, relatively high to moderate in those from Madagascar, Zambia, Carnaiba/Socoto and Itabira, relatively low in those from Colombia and lowest in those from Nigeria.

The emeralds from Zambia and Carnaiba/Socoto contain relatively high contents of Li and Cs (Figure 4.18). The emeralds from Madagascar show exceptionally high K content (Figure 4.16). The Colombian emeralds show consistently low iron content (Figures 4.9 and 4.17). These unusual high or low contents of Li, Cs, K and Fe contents may be used as locality specific criteria for those emeralds.

The emeralds from Brazil, Madagascar and Zambia (Type II deposits) show relatively moderate to high values of Na, Mg, Li and Cs (alkali elements) and Fe (as well as the total foreign elements) as compared with those of emeralds from Nigeria and Colombia (Type I deposits) (Figure 4.19). The highest contents of foreign elements can be found in the Santa Terezinha emerald from Goias State of Brazil (Type IIa) whereas the lowest contents are in Kaduna Plateau of Nigeria (Type Ia) (Figure 4.19).

Table 4.1: Chemical data of 16 trace elements of Nigerian emeralds (NI), and Colombian emeralds (CO) show atom mole ppm 16 major and trace elements.

Cation (atom mole ppm)	Nigeria (Kaduna Plateau)				Colombia (Cordillera Oriental)			
	NI 1	NI 11	NI 16	NI 27	CO 1	CO 11	CO 18	CO 22
Li	174.38	129.03	118.70	110.39	420.19	174.86	86.24	71.56
Be	102944.42	86637.46	86420.81	102404.53	99083.28	95835.41	97121.61	98043.19
B	40.87	35.06	46.93	46.40	42.90	55.15	62.00	55.90
Na	700.57	957.51	1032.32	444.05	3198.94	1893.79	2411.65	2046.69
Mg	318.14	262.99	238.77	217.96	3096.22	1859.98	2538.37	2148.45
Al	61525.47	64903.11	64397.83	61605.85	70778.95	72223.64	71509.80	71244.78
Si	211173.05	223081.16	223696.21	212241.29	201870.96	206546.76	204782.83	204771.24
P	10.51	10.52	11.21	9.71	77.15	13.82	14.32	65.92
K	84.54	39.00	97.90	70.85	27.93	54.92	118.62	113.73
Ca	61.84	51.64	67.47	79.19	126.47	150.56	140.36	113.20
Ti	3.60	7.22	5.48	1.90	2.54	1.10	1.00	0.83
V	145.26	165.05	159.88	108.04	90.64	50.88	49.24	137.27
Cr	105.49	88.90	115.23	121.48	28.75	27.69	27.32	112.07
Mn	2.39	2.97	5.70	1.78	0.49	0.54	0.51	1.23
Fe	1674.49	2591.87	2516.78	1472.44	152.99	54.42	70.24	72.67
Cs	0.06	0.09	24.02	16.02	1.62	1.09	1.00	1.28
<b>Total atom mole %</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.89</b>	<b>37.89</b>	<b>37.90</b>

Table 4.2: Chemical data of 16 trace elements of emeralds from Santa Terezinha (STA; Brazil) and Madagascar (MA) show atom mole ppm 16 major and trace elements.

Cation (atom mole ppm)	Brazil (Santa Terezinha)			Madagascar (Mananjary)			
	STA 1	STA 3	STA 8	MA 4	MA 7	MA 8	MA 18
Li	361.17	382.93	498.33	283.32	216.36	193.46	132.98
Be	95938.26	95213.29	96584.19	97419.18	99566.39	96865.60	99501.36
B	31.51	32.23	27.67	43.57	30.66	44.37	43.45
Na	12395.65	12048.93	12134.87	11737.03	8795.89	9006.50	9809.61
Mg	11396.61	11445.15	11613.44	11621.37	9593.88	9309.79	9707.25
Al	53464.63	53359.60	52854.39	49331.88	51076.70	50975.58	51060.39
Si	201007.10	201268.48	199574.99	204602.02	204334.76	206903.82	204963.37
P	40.96	62.75	44.46	14.12	10.61	9.23	10.48
K	172.04	166.71	193.87	1265.27	2962.34	2738.60	1270.28
Ca	44.28	50.16	140.37	159.65	118.27	129.17	88.51
Ti	1.83	2.81	6.78	2.05	2.51	2.54	2.33
V	134.90	139.76	116.57	20.23	17.35	25.88	19.96
Cr	1358.84	1810.83	1249.17	377.58	237.43	440.39	353.78
Mn	2.39	7.61	7.75	4.27	6.33	3.56	3.99
Fe	2587.58	2947.36	3866.50	2027.88	1985.30	2292.87	1972.50
Cs	62.26	61.40	86.64	35.48	12.17	18.71	16.50
Total atom mole %	37.90	37.90	37.90	37.89	37.90	37.90	37.90





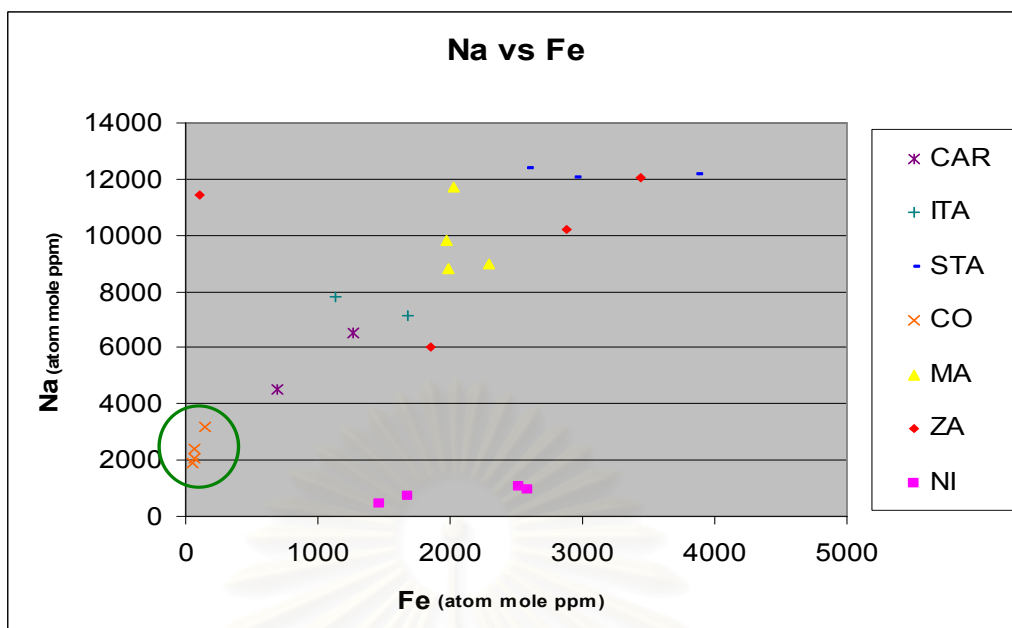


Figure 4.9: Plot of the average Na versus Fe contents of 23 emeralds from different geological occurrences show consistently low iron content in Colombian emerald.

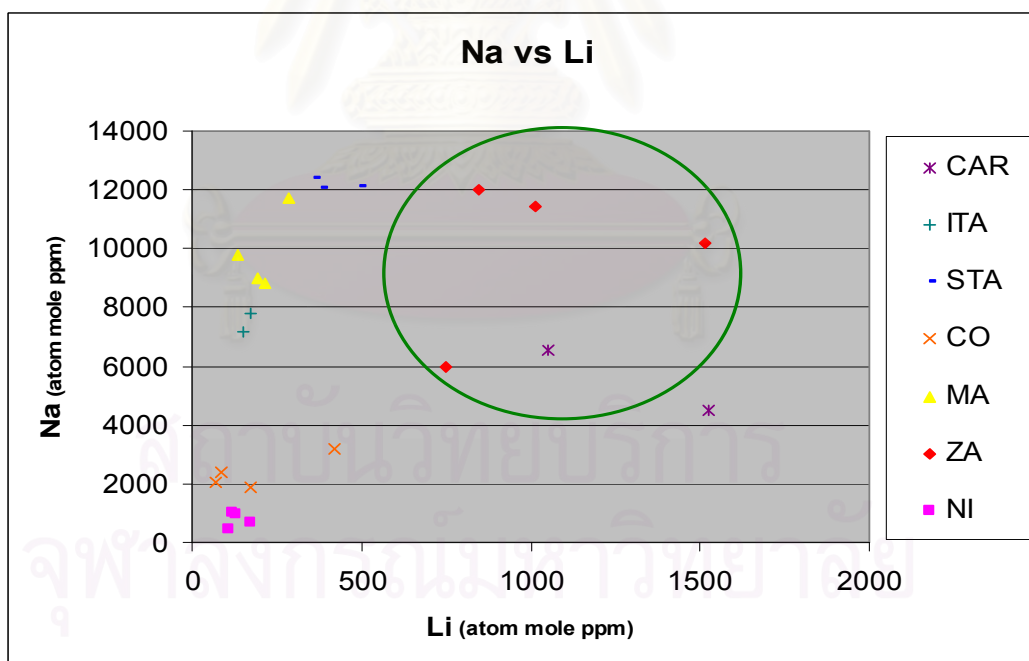


Figure 4.10: Plot of the average Na versus Li contents of 23 emeralds from different geological occurrences show moderate to high lithium content in Zambian emerald.

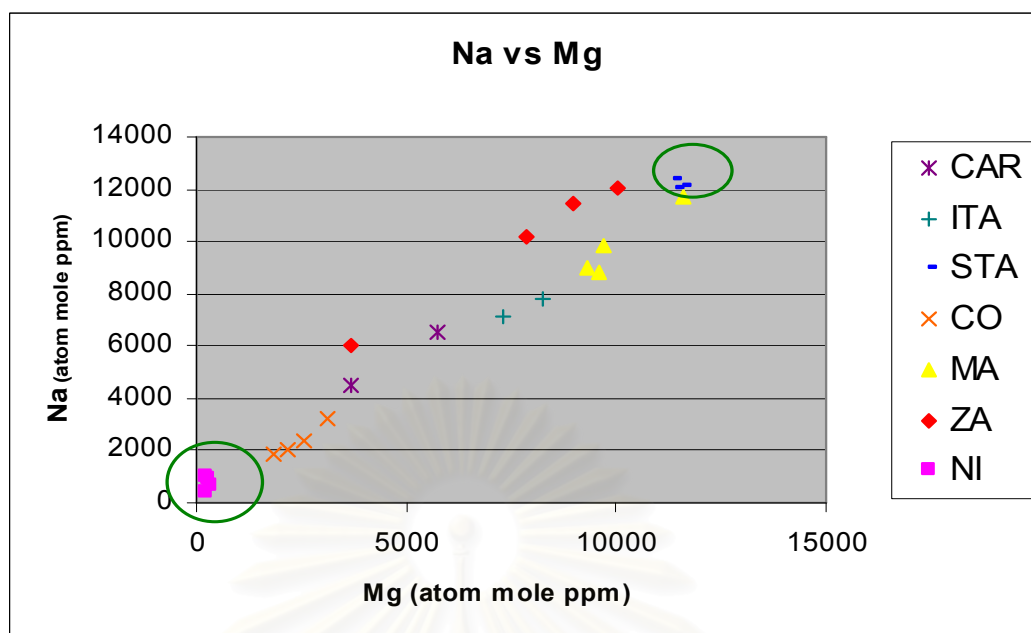


Figure 4.11: Plot of the average Na versus Mg contents of 23 emeralds from different geological occurrences showing the low sodium and magnesium (alkali) content in Nigerian emerald and the highest alkali content in Santa Terezinha from Brazil

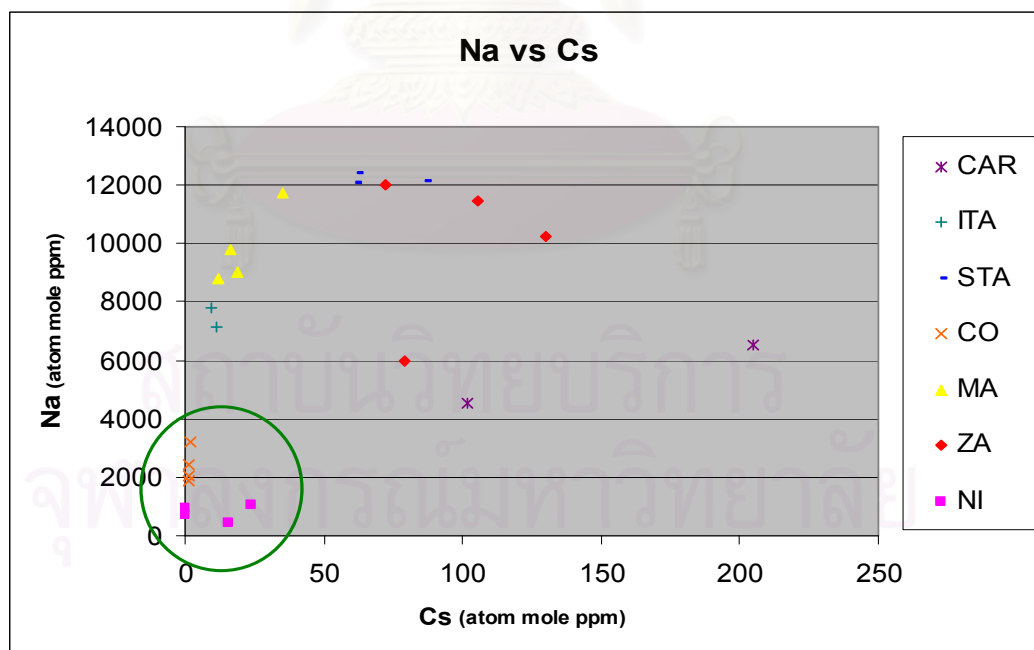


Figure 4.12: Plot of average Na versus Cs contents of 23 emeralds from different geological occurrences show low sodium and cesium content in emerald from Colombia and Nigeria.

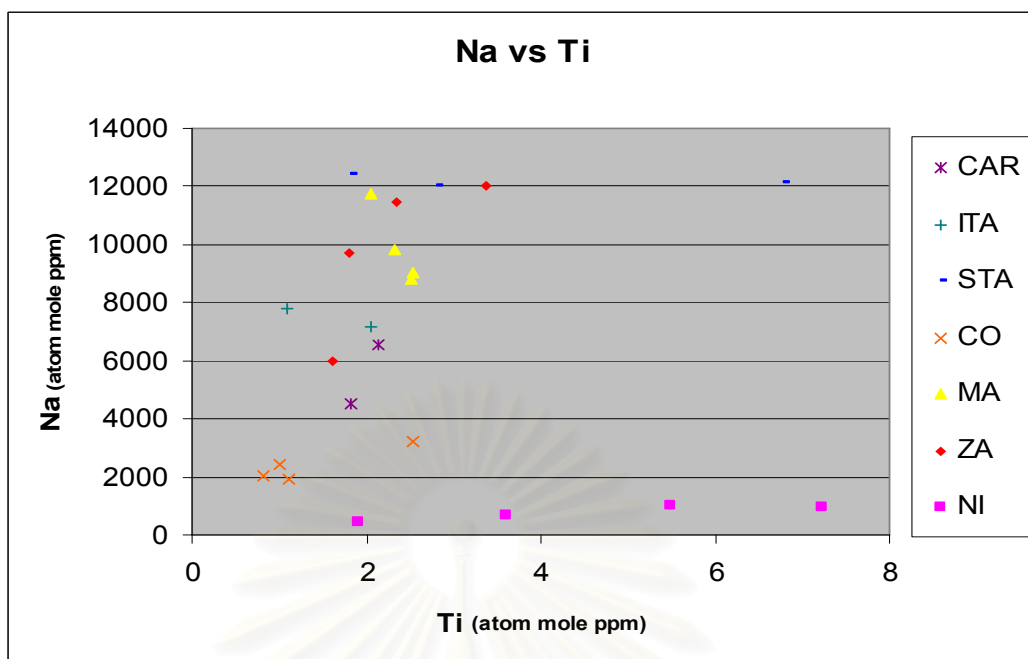


Figure 4.13: Plot of average Na versus Ti contents of 23 emeralds from different geological occurrences show no correlation of other pair of elements.

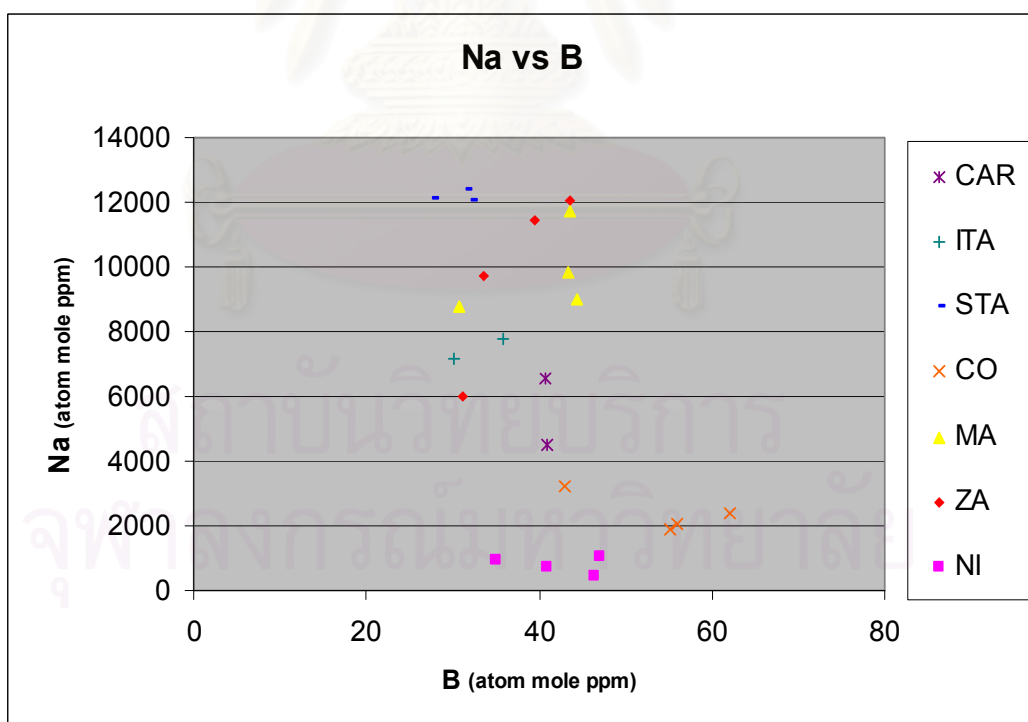


Figure 4.14: Plot of average Na versus B contents of 23 emeralds from different geological occurrences show no correlation of other pair of elements.



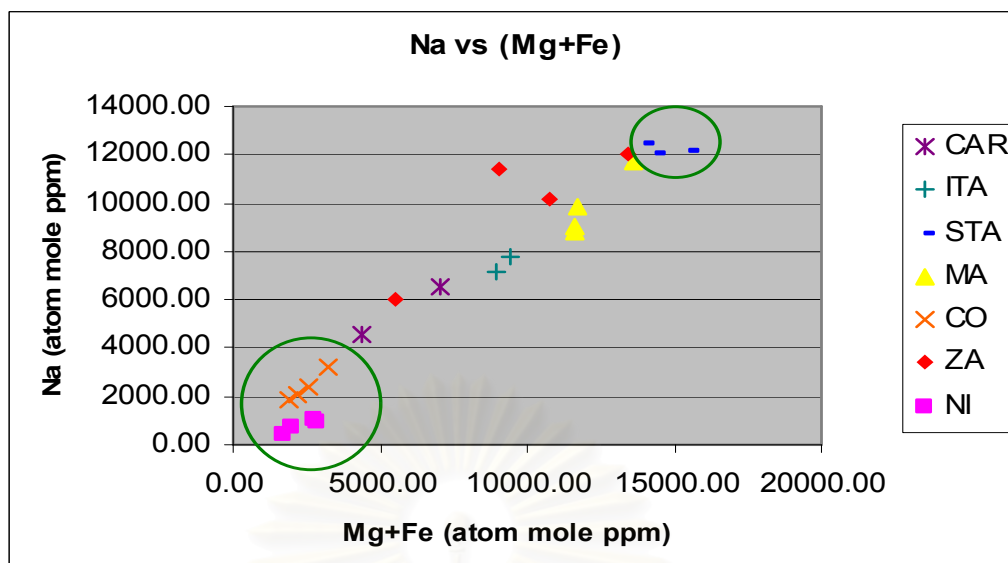


Figure 4.15: Plot of average Na versus (Mg+Fe) contents of 23 emeralds from different geological occurrences show low sodium and (magnesium plus iron) content in emerald from Nigeria and Colombia whereas consistently high these elements content in Santa Terezinha from Brazil.

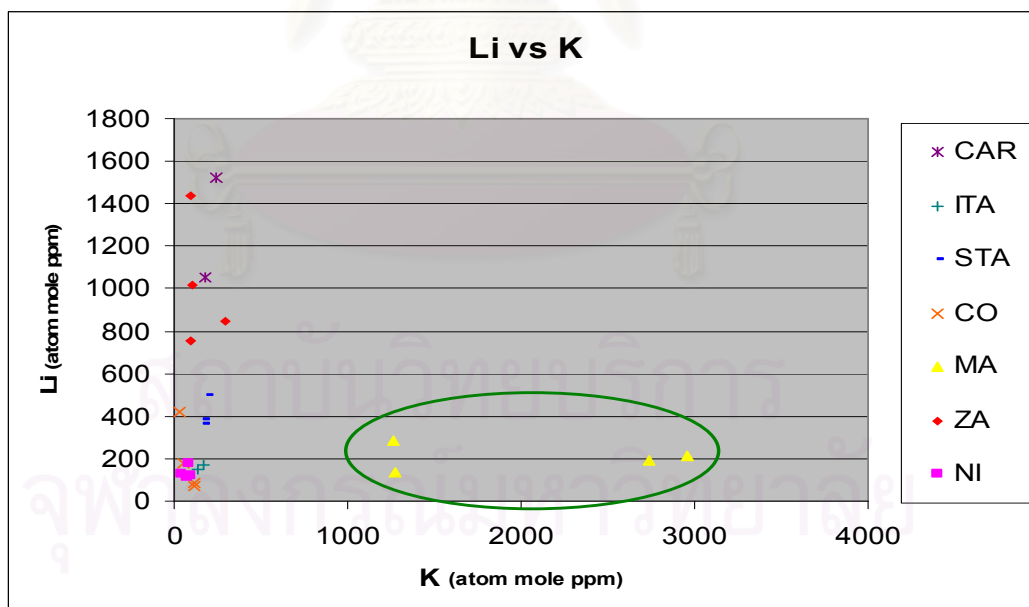


Figure 4.16: Plot of average Li versus K contents of 23 emeralds from different geological occurrences show consistently moderate to high potassium content in emerald from Madagascar.

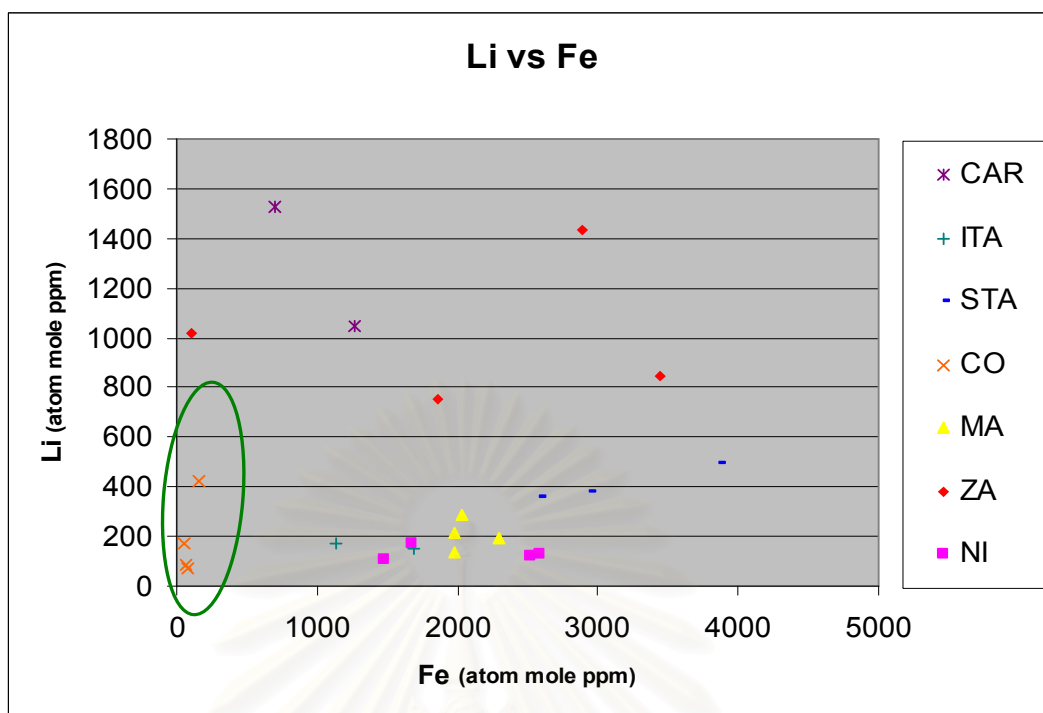


Figure 4.17: Plot of average Li versus Fe contents of 23 emeralds from different geological occurrences show consistently low iron content in Colombian emerald.

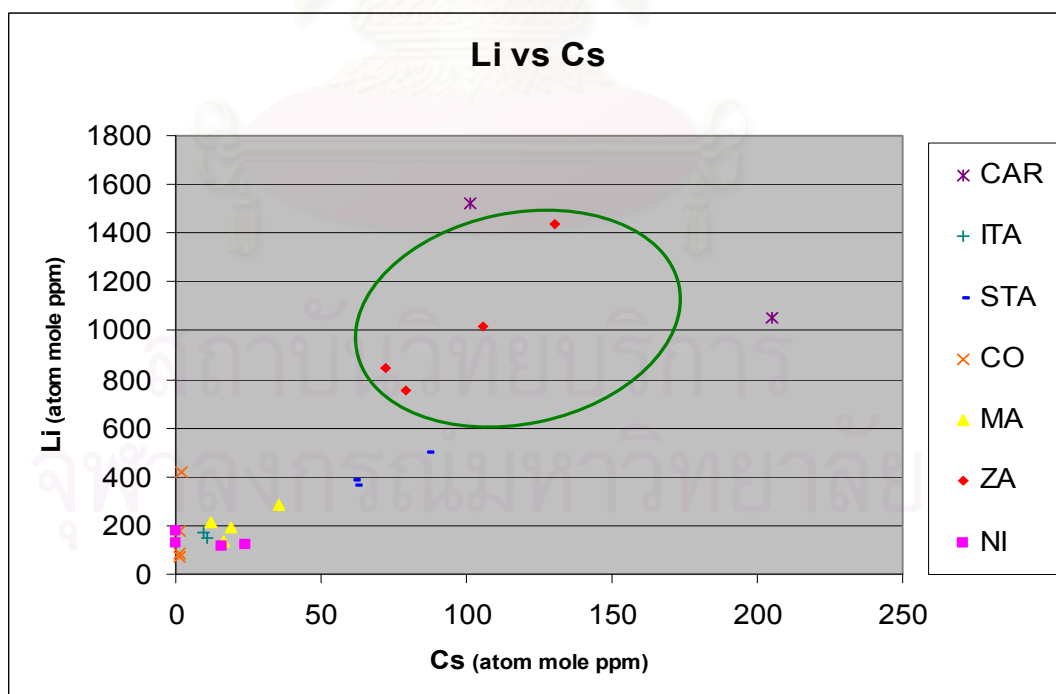


Figure 4.18: Plot of average Li versus Cs contents of 23 emeralds from different geological occurrences show moderate lithium and cesium content in Zambian emerald.

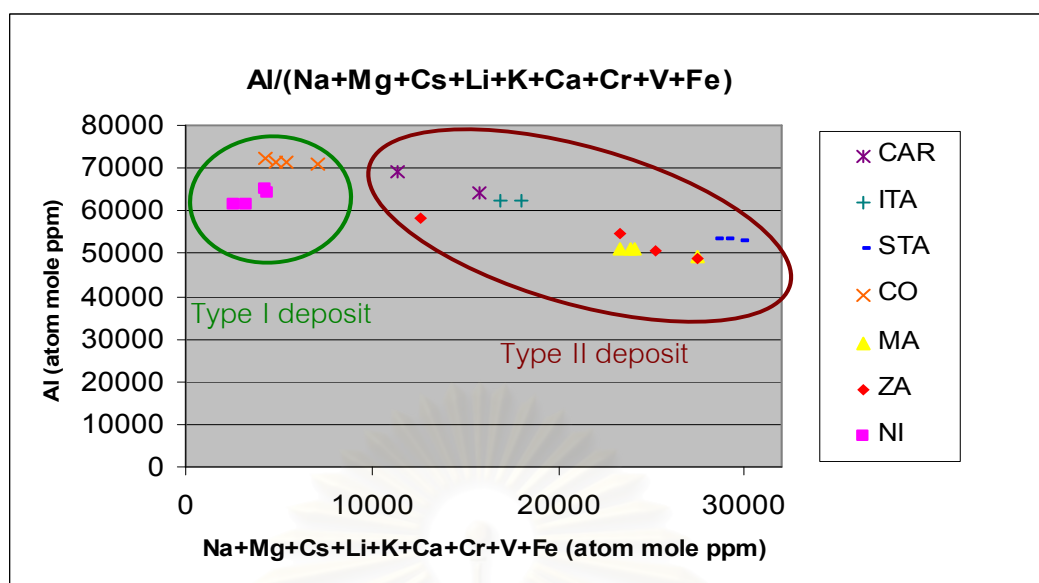


Figure 4.19: Plot of average Al versus sum of all Monovalent+Divalent+Trivalent ions contents from 7 different geological occurrences.

In most emeralds, the contents of sodium and magnesium are higher than the contents of chromophores, except those in Nigerian emeralds. Comparison of significant alkali elements (Na and Mg) and colouring elements (Cr, V and Fe) is shown below. Among those emeralds, the Colombian's contain lowest chromophoric elements because the Colombian emeralds have the lowest iron content (Figures 4.9 and 4.17).

	SUM(Na+Mg) (atom mole ppm)	<	>	SUM(Cr+V+Fe) (atom mole ppm)
Nigeria (NI):	1043.07	<		2316.23
Colombia (CO):	4798.52	>		218.54
Santa Terezinha (STA):	23678.22	>		4190.93
Madagascar (MA):	19895.33	>		2442.79
Zambia (ZA):	17564.25	>		3209.45
Carnaiba/Socoto (CAR):	10236.26	>		775.92
Itabira (ITA):	15239.84	>		1826.75

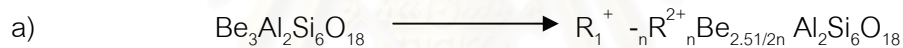
## CHAPTER IV

### CHEMICAL ANALYSIS

#### 4.1 Introduction

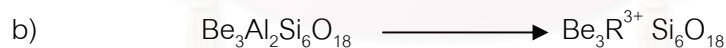
The major components of a natural beryl are  $\text{BeO}$ ,  $\text{Al}_2\text{O}_3$  and  $\text{SiO}_2$  which are normally accompanied by minor alkali elements such as Na, Mg, Cs,... (Bosshart, 1991). Other important chromophoric elements are Cr, V, Fe. In addition, water molecules are commonly found in channels of ring of cyclosilicate structure of emerald (see Figure 3.3)

The ideal formula of a beryl is  $\text{Be}_3\text{Al}_2\text{Si}_6\text{O}_{18}$ , in which alkali, water and other trace elements can substitute into emerald/beryl lattice or structural channels, as follows (Feklichev et al., 1963 in (a) and Schaller et al., 1962 in (b) cited in Sinkankas, 1989):



where:  $\text{R}^+ = \text{Cs}^+, \text{Li}^+, \text{Rb}^+, \text{Na}^+, \text{K}^+$

where:  $\text{R}^{2+} = \text{Ca}^{2+}, \text{Ba}^{2+}, \text{Sr}^{2+}$ ; and  $n = 0, 1$



where:  $\text{R}^{3+} = \text{Al}^{3+}, \text{Fe}^{3+}, \text{Cr}^{3+}, \text{Sc}^{3+}$

The fact that many cations can substitute into the emerald structure, it is expected that emeralds from different geological environments should contain different amounts of trace elements in their structures. Hence trace element contents should be applicable as a chemical fingerprint of emeralds from various geologic occurrences. It is therefore the aim of this study to look into the trace element concentration of emeralds from various locations.



## 4.2 Samples and Procedure

In this study, the chemical analysis of emerald was carried out by a Laser ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICP-MS). This technique is suitable for analysis of trace elements in particular the light elements such as Li, Be, B, Na, Mg, despite the fact that it is a “destructive” technique.

The LA-ICP-MS unit at the Macquarie University in Australia was used in this study (Figure 4.1). Detailed descriptions of LA-ICP-MS instrumentation analytical and calibration procedures are similar to those given by Norman et al. (1996). The UV laser ablation microprobe (a New Wave Research 213 nm Nb:YAG) is coupled to an Agilent 7500S ICP-MS. All analyses were done with a pulse rate of 5 Hz and a beam energy of approximately 0.5 mJ per pulse, producing a spatial resolution 50 micrometers. The beam is used to vaporize an extremely small sample of material (in this case, emerald). The ablated sample is carried by a stream of inert gas, usually argon, into a high-temperature field, causing dissociation of molecules and ionization of atoms.

The MS identifies and quantifies elements in terms of mass and charge. Some 41 elements and their relative amounts can be detected even when present in only a few parts per billion. Quantitative results for 16 trace elements (Li, Be, B, Na, Mg, Al, Sr, P, K, Ca, Ti, V, Cr, Mn, Fe and Cs of the emerald samples) were obtained through calibration of relative element sensitivities using the NIST-610 multi-element glass standard and an emerald with 12.5 wt% BeO as internal standards. The BCR2G basaltic glass standard was also used as an external standard. The detection limits vary from analysis to analysis and are typically less than 1 ppm for Li, Be, V and Cs; less than 4 ppm for Mg, Al, Ti, Mn and Cr; less than 13 ppm for B and Na; less than 80 ppm for P and Fe; less than 300 ppm for Si and K and less than 500 ppm for Ca.

In this study a total of 23 emerald samples from 7 different geological occurrences (2-4 representative samples from each locality) were selected for chemical analysis. In each sample at least 5 spots were analyzed across the polished surface normally cut perpendicular to c axis. The detailed positions of the spots analyzed are presented in Figures 4.2 to 4.8.



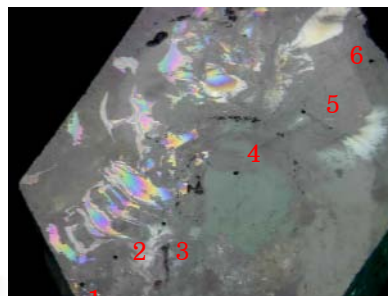
Figure 4.1: Laser ablation Inductively Coupled Plasma Mass Spectrometry Laser Ablation (LA-ICPMS) unit at Macquarie University, Australia used for chemical analysis in this study.

The representative laser points of 23 samples from different geological occurrences in below:

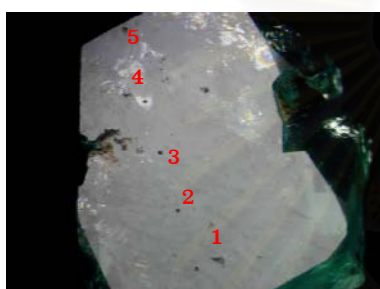
NI 1



NI 11



NI 16



NI 27

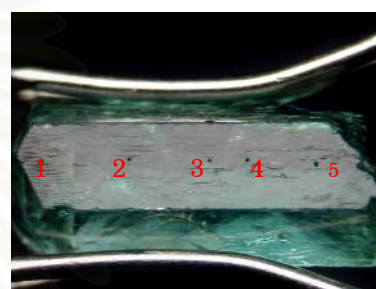
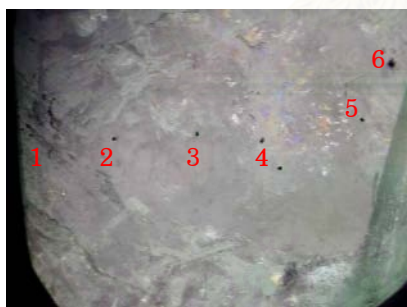
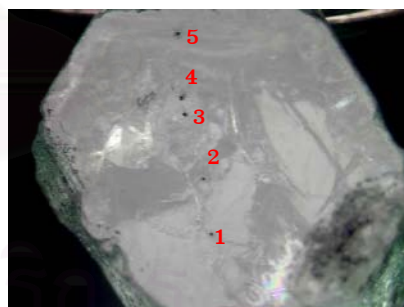


Figure 4.2: Showing the 4 samples of emerald from Kaduna Plateau, Nigeria.

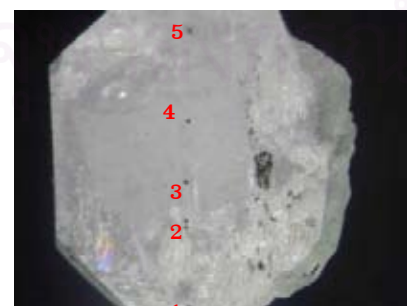
CO 1



CO 11



CO 18



CO 22

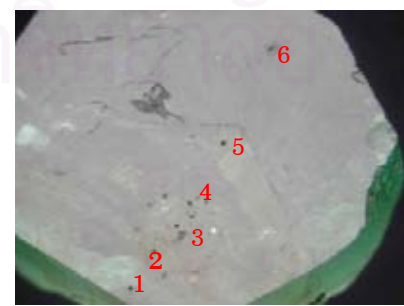
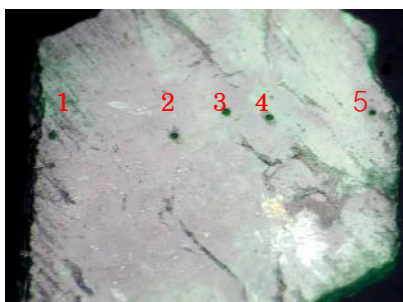
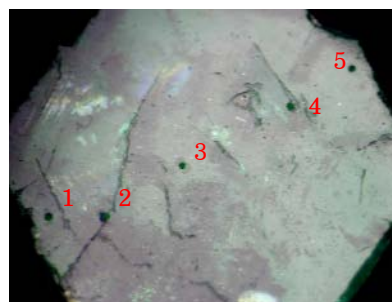


Figure 4.3: Showing the 4 samples of emerald from Cordillera Oriental, Colombia.

STA 1



STA 3



STA 8

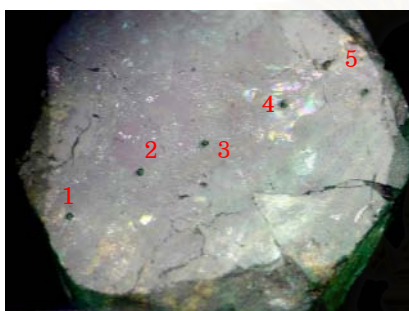
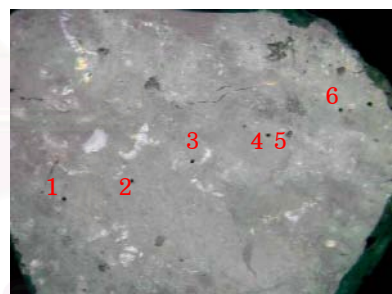


Figure 4.4: Showing the 3 samples of emerald from Santa Terezinha de Goais, Brazil.

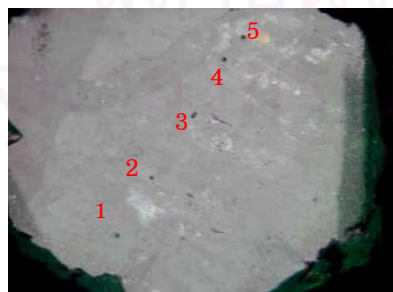
MA 4



MA 7



MA 8



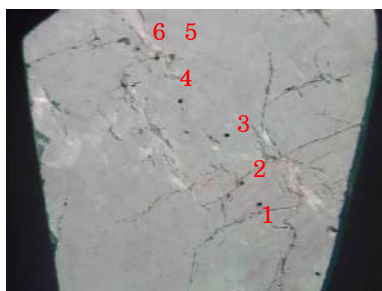
MA 18



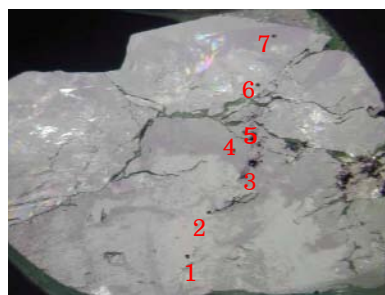
Figure 4.5: Showing the 4 samples of emerald from Mananjary region, Madagascar.



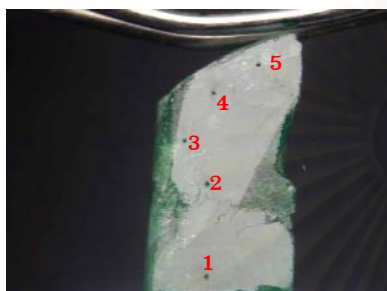
ZA 5



ZA 6



ZA 23

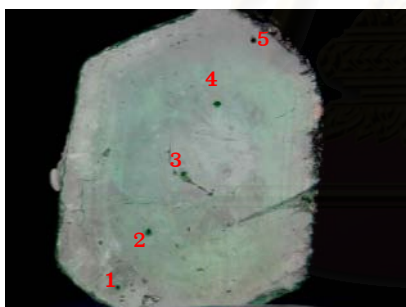


ZA 25



Figure 4.6: Showing the 4 samples of emerald from Ndola Rural district, Zambia.

CAR 4



CAR 7

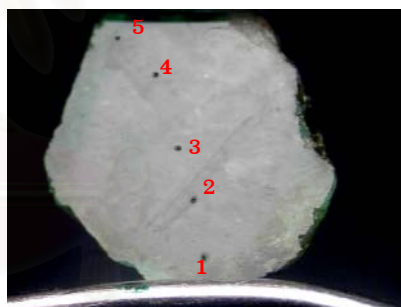
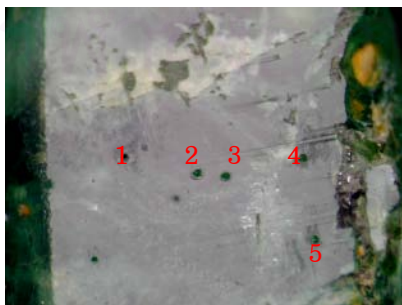


Figure 4.7: Showing the 2 samples of emerald from Carnaiba/Socoto, Bahia State, Brazil.

ITA 2



ITA 10

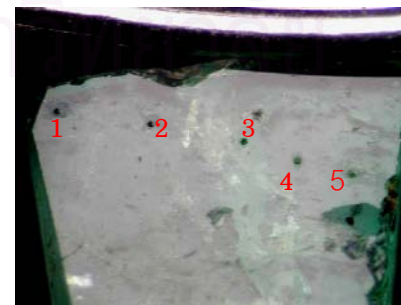


Figure 4.8: Showing the 2 samples of emerald from Itabira, Minas Gerias state in Brazil.



### 4.3 Results

The average concentrations of 16 trace elements; Li, Be, B, Na, Mg, Al, Si, P, K, Ca, Ti, V, Cr, Mn, Fe and Cs of each emerald sample analyzed in this study after recalculation into atom mole ppm are listed in Tables 4.1 to 4.3. The analytical concentrations in ppm by weight of original data, the plot of averages ( $\bar{X}$ ) and standard deviation (STD) are shown in Appendix IV

From the chemical data shown Tables 4.1 to 4.3, the elements of emeralds from this study are mainly as alkalis and iron. Amount of these elements are characteristic, which related to geological environments.

### 4.4 Cross plots

Cross plots of Na versus Fe (or Na vs Fe), Na vs Li, Na vs Mg, Na vs Cs, Na vs Ti, Na vs B, Na vs (Mg+Fe), Li vs K, Li vs Fe, Li vs Cs and Al vs sum of all monovalent+divalent+trivalent ions contents of all the emeralds in this study are shown in Figures 4.9 to 4.19. The majority of the plots include alkali elements because of the fact that they are the major trace elements which have the strongest influence on the optical, physical and chemical properties of emeralds.

A good correlation between Na and Mg is seen in Figures 4.11 and between Na and (Mg+Fe) in Figure 4.15 regardless of localities. The remaining cross plots seem to show no correlation of other pair of elements. With regard to Figures 4.11 and 4.15, the contents of Na and Mg and probably Fe are highest in emeralds from Santa Terezinha, relatively high to moderate in those from Madagascar, Zambia, Carnaiba/Socoto and Itabira, relatively low in those from Colombia and lowest in those from Nigeria.

The emeralds from Zambia and Carnaiba/Socoto contain relatively high contents of Li and Cs (Figure 4.18). The emeralds from Madagascar show exceptionally high K content (Figure 4.16). The Colombian emeralds show consistently low iron content (Figures 4.9 and 4.17). These unusual high or low contents of Li, Cs, K and Fe contents may be used as locality specific criteria for those emeralds.

The emeralds from Brazil, Madagascar and Zambia (Type II deposits) show relatively moderate to high values of Na, Mg, Li and Cs (alkali elements) and Fe (as well as the total foreign elements) as compared with those of emeralds from Nigeria and Colombia (Type I deposits) (Figure 4.19). The highest contents of foreign elements can be found in the Santa Terezinha emerald from Goias State of Brazil (Type IIa) whereas the lowest contents are in Kaduna Plateau of Nigeria (Type Ia) (Figure 4.19).

Table 4.1: Chemical data of 16 trace elements of Nigerian emeralds (NI), and Colombian emeralds (CO) show atom mole ppm 16 major and trace elements.

Cation (atom mole ppm)	Nigeria (Kaduna Plateau)				Colombia (Cordillera Oriental)			
	NI 1	NI 11	NI 16	NI 27	CO 1	CO 11	CO 18	CO 22
Li	174.38	129.03	118.70	110.39	420.19	174.86	86.24	71.56
Be	102944.42	86637.46	86420.81	102404.53	99083.28	95835.41	97121.61	98043.19
B	40.87	35.06	46.93	46.40	42.90	55.15	62.00	55.90
Na	700.57	957.51	1032.32	444.05	3198.94	1893.79	2411.65	2046.69
Mg	318.14	262.99	238.77	217.96	3096.22	1859.98	2538.37	2148.45
Al	61525.47	64903.11	64397.83	61605.85	70778.95	72223.64	71509.80	71244.78
Si	211173.05	223081.16	223696.21	212241.29	201870.96	206546.76	204782.83	204771.24
P	10.51	10.52	11.21	9.71	77.15	13.82	14.32	65.92
K	84.54	39.00	97.90	70.85	27.93	54.92	118.62	113.73
Ca	61.84	51.64	67.47	79.19	126.47	150.56	140.36	113.20
Ti	3.60	7.22	5.48	1.90	2.54	1.10	1.00	0.83
V	145.26	165.05	159.88	108.04	90.64	50.88	49.24	137.27
Cr	105.49	88.90	115.23	121.48	28.75	27.69	27.32	112.07
Mn	2.39	2.97	5.70	1.78	0.49	0.54	0.51	1.23
Fe	1674.49	2591.87	2516.78	1472.44	152.99	54.42	70.24	72.67
Cs	0.06	0.09	24.02	16.02	1.62	1.09	1.00	1.28
<b>Total atom mole %</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.90</b>	<b>37.89</b>	<b>37.89</b>	<b>37.90</b>

Table 4.2: Chemical data of 16 trace elements of emeralds from Santa Terezinha (STA; Brazil) and Madagascar (MA) show atom mole ppm 16 major and trace elements.

Cation (atom mole ppm)	Brazil (Santa Terezinha)			Madagascar (Mananjary)			
	STA 1	STA 3	STA 8	MA 4	MA 7	MA 8	MA 18
Li	361.17	382.93	498.33	283.32	216.36	193.46	132.98
Be	95938.26	95213.29	96584.19	97419.18	99566.39	96865.60	99501.36
B	31.51	32.23	27.67	43.57	30.66	44.37	43.45
Na	12395.65	12048.93	12134.87	11737.03	8795.89	9006.50	9809.61
Mg	11396.61	11445.15	11613.44	11621.37	9593.88	9309.79	9707.25
Al	53464.63	53359.60	52854.39	49331.88	51076.70	50975.58	51060.39
Si	201007.10	201268.48	199574.99	204602.02	204334.76	206903.82	204963.37
P	40.96	62.75	44.46	14.12	10.61	9.23	10.48
K	172.04	166.71	193.87	1265.27	2962.34	2738.60	1270.28
Ca	44.28	50.16	140.37	159.65	118.27	129.17	88.51
Ti	1.83	2.81	6.78	2.05	2.51	2.54	2.33
V	134.90	139.76	116.57	20.23	17.35	25.88	19.96
Cr	1358.84	1810.83	1249.17	377.58	237.43	440.39	353.78
Mn	2.39	7.61	7.75	4.27	6.33	3.56	3.99
Fe	2587.58	2947.36	3866.50	2027.88	1985.30	2292.87	1972.50
Cs	62.26	61.40	86.64	35.48	12.17	18.71	16.50
Total atom mole %	37.90	37.90	37.90	37.89	37.90	37.90	37.90





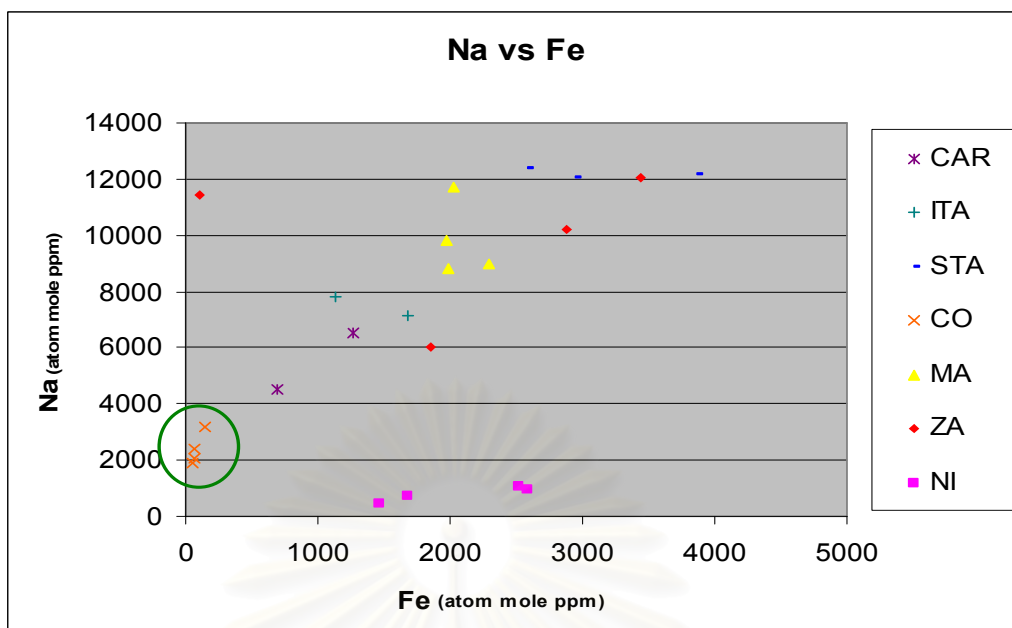


Figure 4.9: Plot of the average Na versus Fe contents of 23 emeralds from different geological occurrences show consistently low iron content in Colombian emerald.

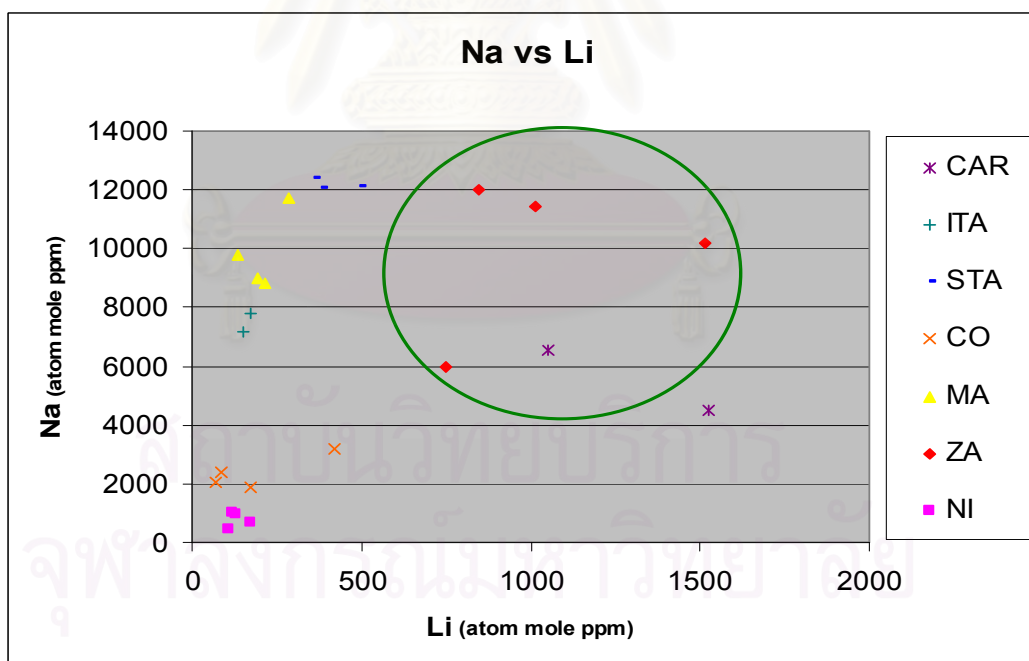


Figure 4.10: Plot of the average Na versus Li contents of 23 emeralds from different geological occurrences show moderate to high lithium content in Zambian emerald.

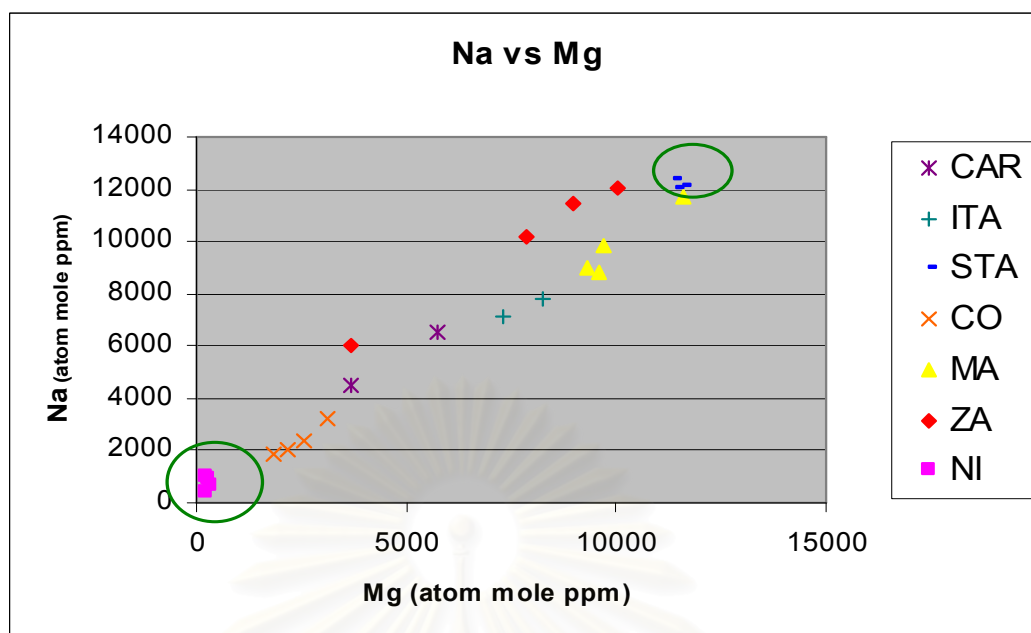


Figure 4.11: Plot of the average Na versus Mg contents of 23 emeralds from different geological occurrences showing the low sodium and magnesium (alkali) content in Nigerian emerald and the highest alkali content in Santa Terezinha from Brazil

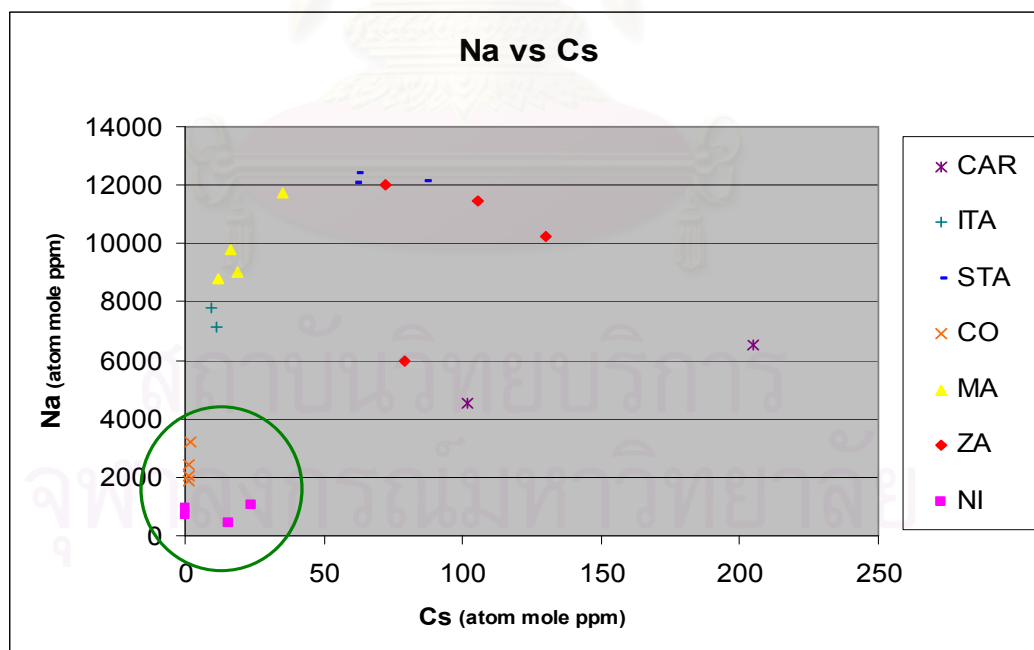


Figure 4.12: Plot of average Na versus Cs contents of 23 emeralds from different geological occurrences show low sodium and cesium content in emerald from Colombia and Nigeria.

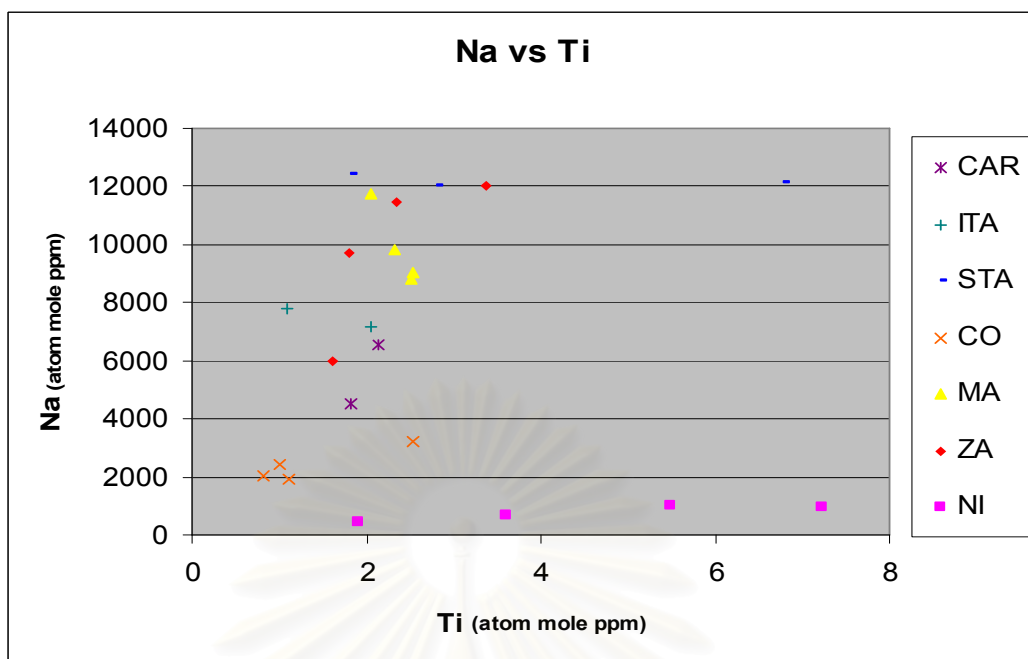


Figure 4.13: Plot of average Na versus Ti contents of 23 emeralds from different geological occurrences show no correlation of other pair of elements.

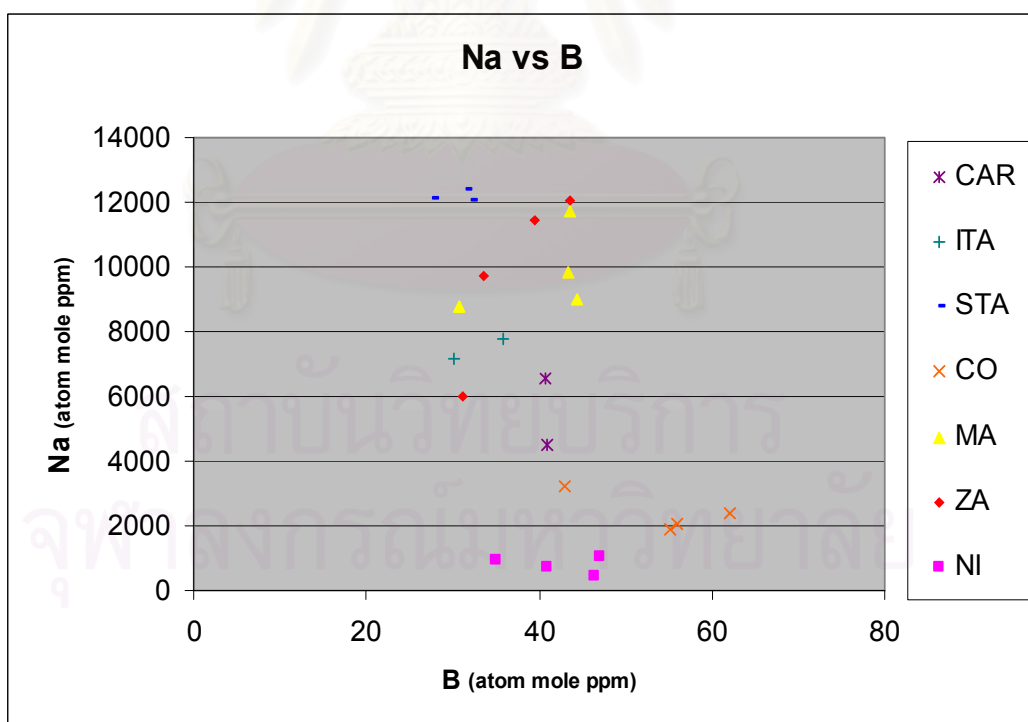


Figure 4.14: Plot of average Na versus B contents of 23 emeralds from different geological occurrences show no correlation of other pair of elements.

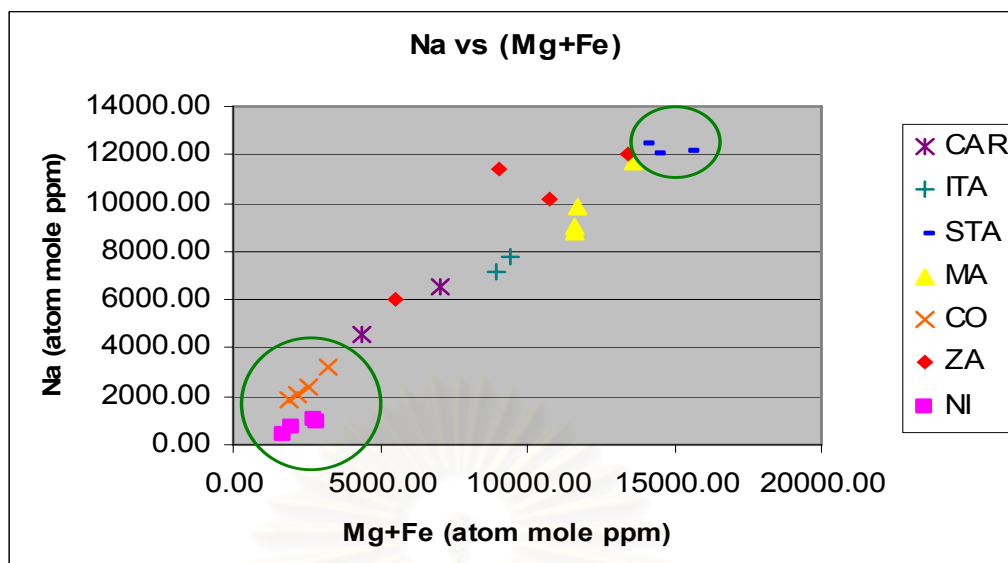


Figure 4.15: Plot of average Na versus (Mg+Fe) contents of 23 emeralds from different geological occurrences show low sodium and (magnesium plus iron) content in emerald from Nigeria and Colombia whereas consistently high these elements content in Santa Terezinha from Brazil.

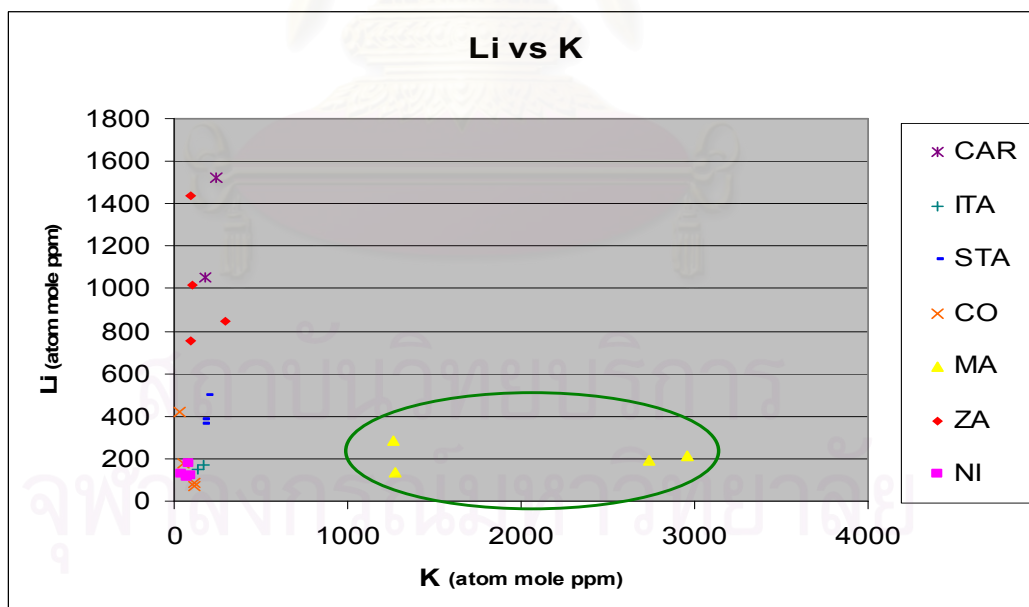


Figure 4.16: Plot of average Li versus K contents of 23 emeralds from different geological occurrences show consistently moderate to high potassium content in emerald from Madagascar.

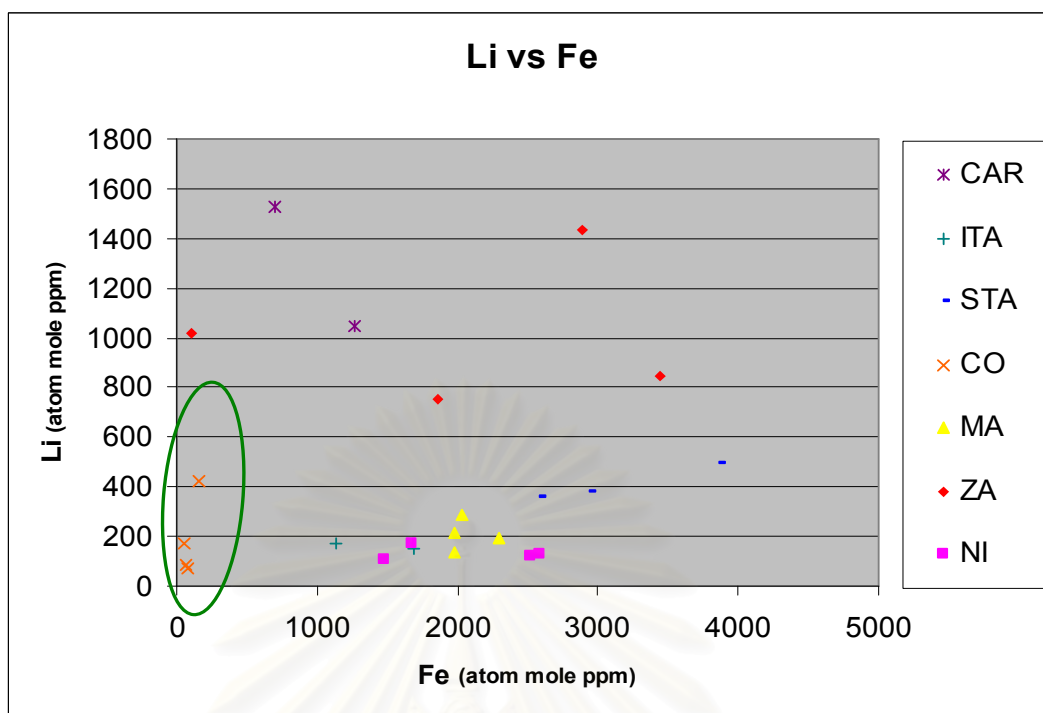


Figure 4.17: Plot of average Li versus Fe contents of 23 emeralds from different geological occurrences show consistently low iron content in Colombian emerald.

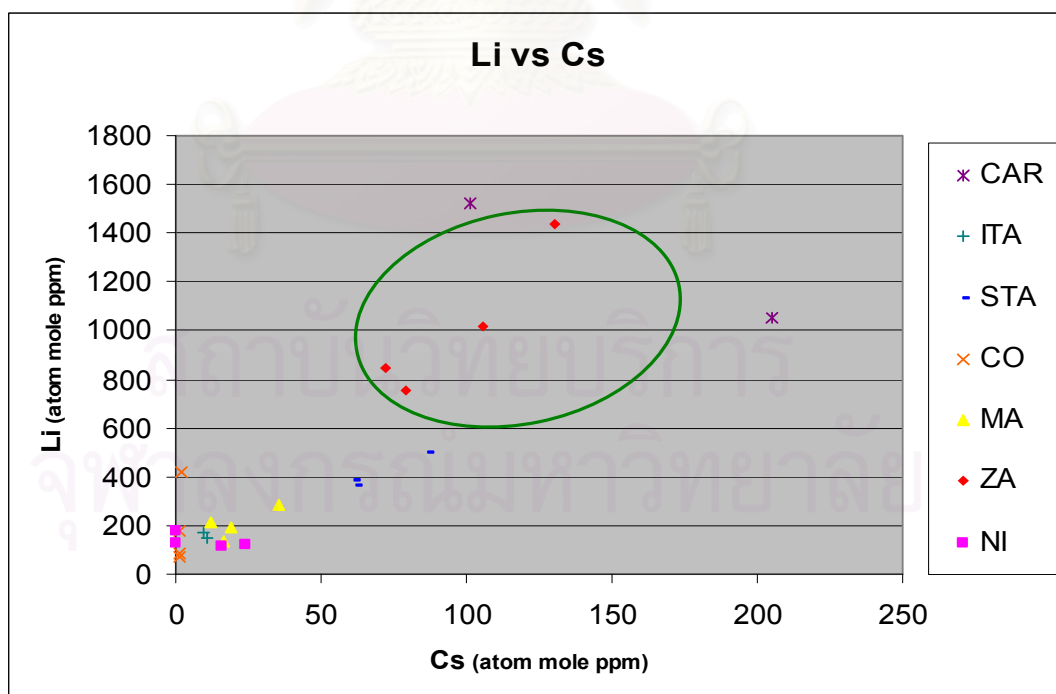


Figure 4.18: Plot of average Li versus Cs contents of 23 emeralds from different geological occurrences show moderate lithium and cesium content in Zambian emerald.



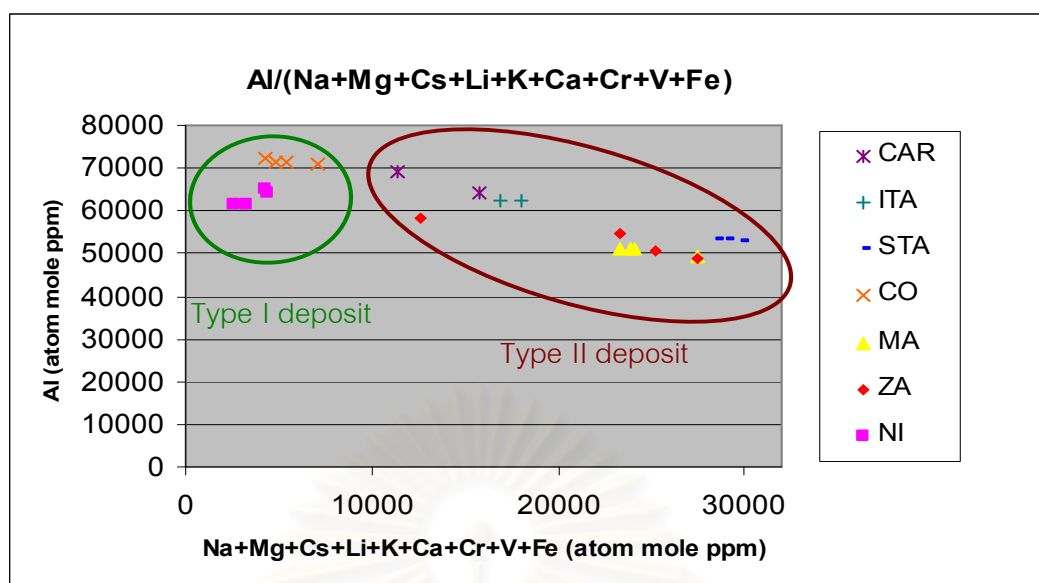


Figure 4.19: Plot of average Al versus sum of all Monovalent+Divalent+Trivalent ions contents from 7 different geological occurrences.

In most emeralds, the contents of sodium and magnesium are higher than the contents of chromophores, except those in Nigerian emeralds. Comparison of significant alkali elements (Na and Mg) and colouring elements (Cr, V and Fe) is shown below. Among those emeralds, the Colombian's contain lowest chromophoric elements because the Colombian emeralds have the lowest iron content (Figures 4.9 and 4.17).

	SUM(Na+Mg) (atom mole ppm)		SUM(Cr+V+Fe) (atom mole ppm)
Nigeria (NI):	1043.07	<	2316.23
Colombia (CO):	4798.52	>	218.54
Santa Terezinha (STA):	23678.22	>	4190.93
Madagascar (MA):	19895.33	>	2442.79
Zambia (ZA):	17564.25	>	3209.45
Carnaiba/Socoto (CAR):	10236.26	>	775.92
Itabira (ITA):	15239.84	>	1826.75

## CHAPTER V

### SPECTROSCOPIC EXAMINATION

In addition to those chemical data, spectroscopic examination was also carried out in this study which was divided into two intervals, namely Ultraviolet-Visible-Near Infrared (UV-Vis-NIR) region and Infrared (IR) region. The absorption spectra in the UV-Vis-NIR region are usually measured by a UV-Vis-NIR spectrophotometer whereas those in the IR region are measured by a Fourier Transform Infrared (FTIR) Spectrophotometer.

#### 5.1 UV-Vis-NIR spectrophotometry

The UV-Vis-NIR spectrophotometer is an equipment used for measurement of the absorption spectra in the UV-Vis-NIR region which can help understand the causes of colour. Typical features of absorption spectra as recorded by the spectrophotometer are assigned to certain trace elements indicating the role of these elements as colouring agents.

##### 5.1.1 Instrument and Analytical condition

In this study, the UV-Vis-NIR absorption spectra of the emerald samples were measured by the spectrophotometer Model HITACHI U-4001 that has the wavelength ranging from 240 to 2500 nm (Figure 5.1) located at the GIT. The spectra were collected with scanning speed of 300 nm/min. and slit opening of 4.00 nm. The spectra were collected at two automatic sampling interval; at the wavelength 240-340 nm of the ultraviolet region (UV), the light sources of deuterium (D2) discharge tube was used ; and at the wavelength 320-1000 nm of the visible and near infrared region, a 50 W iodine tungsten lamp was used. Photomultiplier (for ultraviolet-visible region) and cooled-type PbS cell (for near infrared region) detectors incorporated with automatic wavelength calibration function were used to obtain the spectra with wavelength accuracies of  $\pm 0.2$  nm in ultraviolet-visible region and  $\pm 1.0$  nm in near infrared

region. A standard PC computer with an installed program was used to collect and store the spectra including data processing and analysis. The spectra were recorded with both light polarized parallel (e-ray) and perpendicular (o-ray) to the c-axis by using a Polaroid filter.

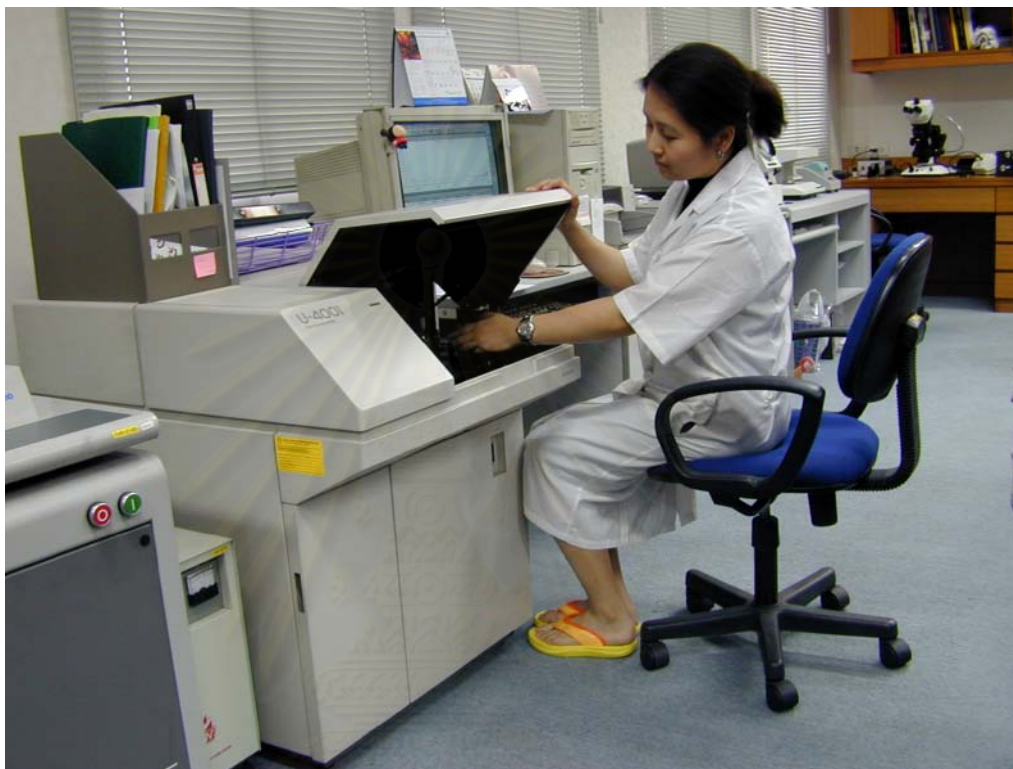


Figure 5.1: A UV-Vis-NIR spectrophotometer Model HITACHI model U-4001 at the GIT.

### 5.1.2 Results

The UV-Vis-NIR absorption spectra of 7 emerald samples from Nigeria, Colombia, Santa Terezinha, Carnaiba/Socoto and Itabira (Brazil), Madagascar and Zambia were measured and representative spectra from each location are shown in Figures 5.2 to 5.8. The spectra of most emeralds except that of the Colombian emerald show typical absorption bands of chromium ( $\text{Cr}^{3+}$ ) and vanadium ( $\text{V}^{3+}$ ) mixed with iron ( $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$ ). The two wide absorption bands of  $\text{Cr}^{3+}$  and  $\text{V}^{3+}$ , the so-called 'emerald component', in 400 – 480 nm and 590 – 640 nm are dominant in those spectra. The

narrow  $\text{Cr}^{3+}$  absorption bands in red to orange region (611, 638, 648, 662, 685 nm) are overlapped by a wide absorption band of the so-called 'aquamarine component' ( $\text{Fe}^{2+}/\text{Fe}^{3+}$  IVCT: inter-valence charge transfer) in 600 – 750 nm and one or two broad band of  $\text{Fe}^{2+}$  in 820 nm (o-ray) and 760 - 920 nm (e-ray). The  $\text{V}^{3+}$  absorption bands (390 – 405 and 570 – 575) are slightly shifted against the  $\text{Cr}^{3+}$  peaks. Other two narrow bands of  $\text{Fe}^{3+}$  are in violet region at about 370 and 426 nm. In contrast the spectra of Colombian emerald show only absorption bands due to  $\text{Cr}^{3+}$  and  $\text{V}^{3+}$  without contribution from iron absorption. The absorption characteristics can be summarized in Table 5.1. Similar results have previously been reported by Schmetzer and Bank (1981), Schwarz and Henn (1992), Schwarz (2002), Petrov and Neumeier (2002) and Bosshart (1982).

Table 5.1: Summary of the absorption bands of emeralds in UV-Vis-NIR ranges found in this study as well as previously reported by Schmetzer and Bank (1981), Bosshart (1982), Schwarz and Henn (1992), Mathew (1998), Schwarz (2002) ,Petrov and Neumeier (2002)

Wave number (nm)	Assignment
about 370, 426	Two narrow bands of $\text{Fe}^{3+}$
611, 638, 648, 662, 685	Narrow bands of $\text{Cr}^{3+}$
400 – 480 and 590 – 640	Two wide bands of emerald component; $\text{Cr}^{3+}$ and $\text{V}^{3+}$
600 – 750	A wide band of aquamarine component; $\text{Fe}^{2+}/\text{Fe}^{3+}$ inter-valence charge transfer
820 (o-ray; in channel site) and 760 – 920 (e-ray; at octahedral Al site)	One or two band of $\text{Fe}^{2+}$
390 – 405 and 570 - 575	$\text{V}^{3+}$ are slightly shifted against the Cr peaks

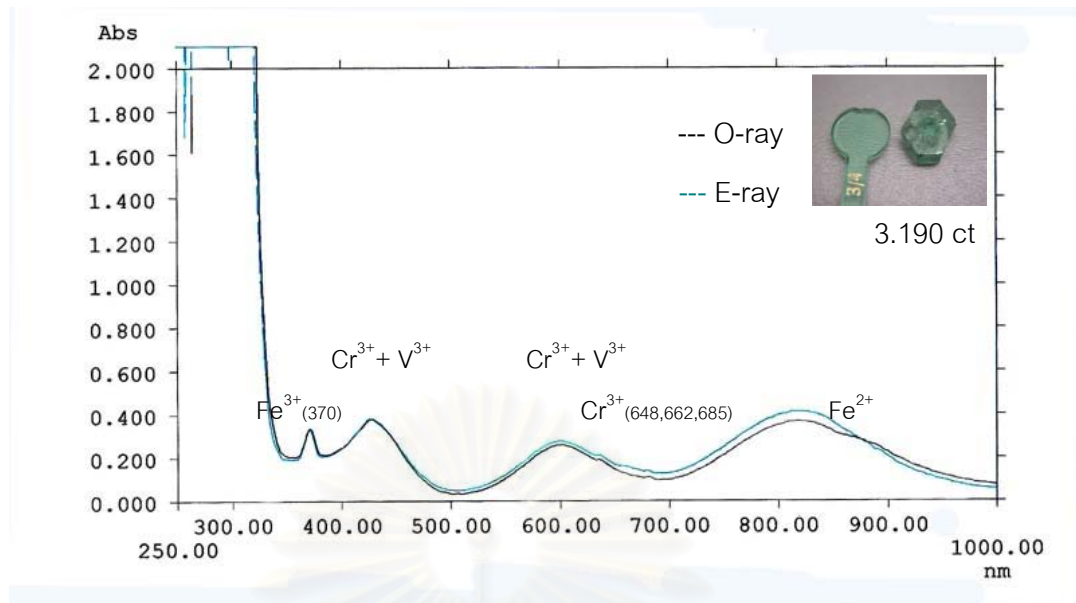


Figure 5.2: UV-Vis-NIR spectra of Nigerian emerald (NI-11) showing chromium-vanadium spectrum, ferric iron ( $\text{Fe}^{3+}$ ) bands in ultraviolet and ferrous iron ( $\text{Fe}^{2+}$ ) band in near-infrared.

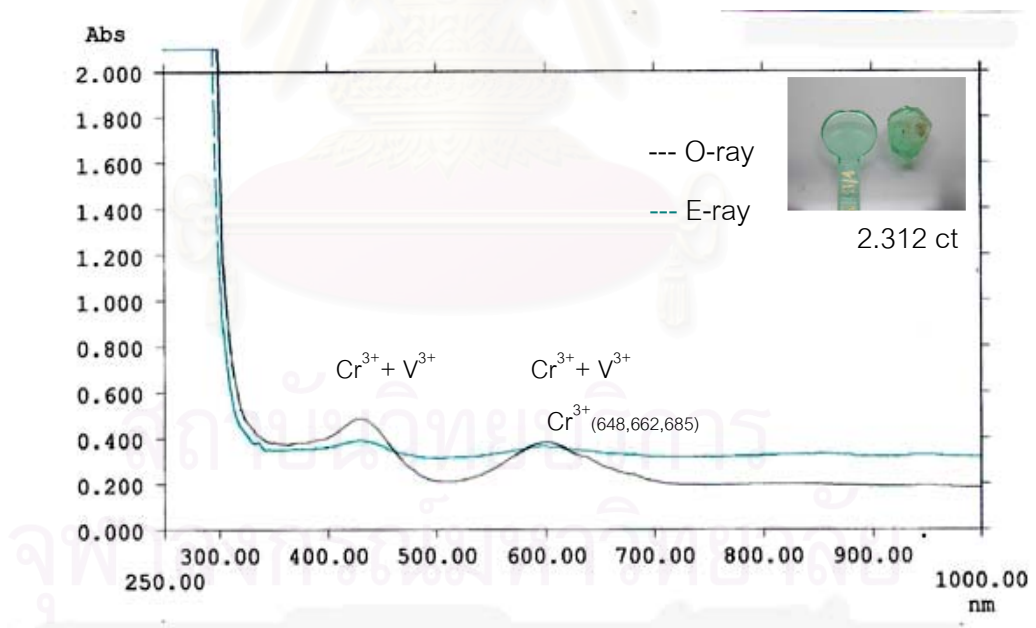


Figure 5.3: UV-Vis-NIR spectra of Colombian emerald (CO-22) showing chromium-vanadium spectrum without iron bands.



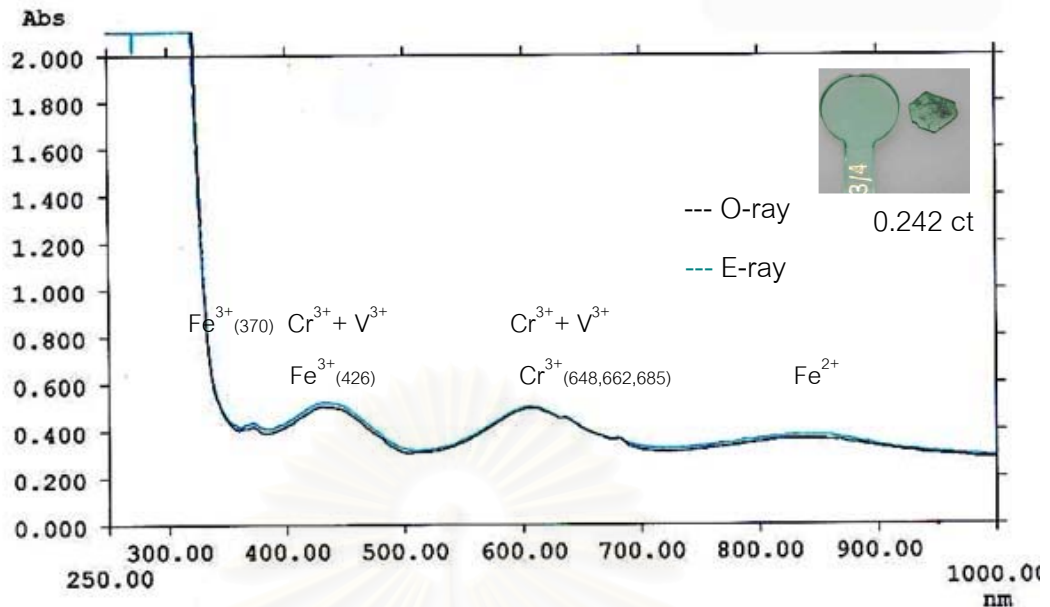


Figure 5.4: UV-Vis-NIR spectra of Santa Terezinha emerald from Brazil (STA-3) showing chromium-vanadium spectrum, ferric iron ( $\text{Fe}^{3+}$ ) bands in ultraviolet and ferrous iron ( $\text{Fe}^{2+}$ ) band in near-infrared.

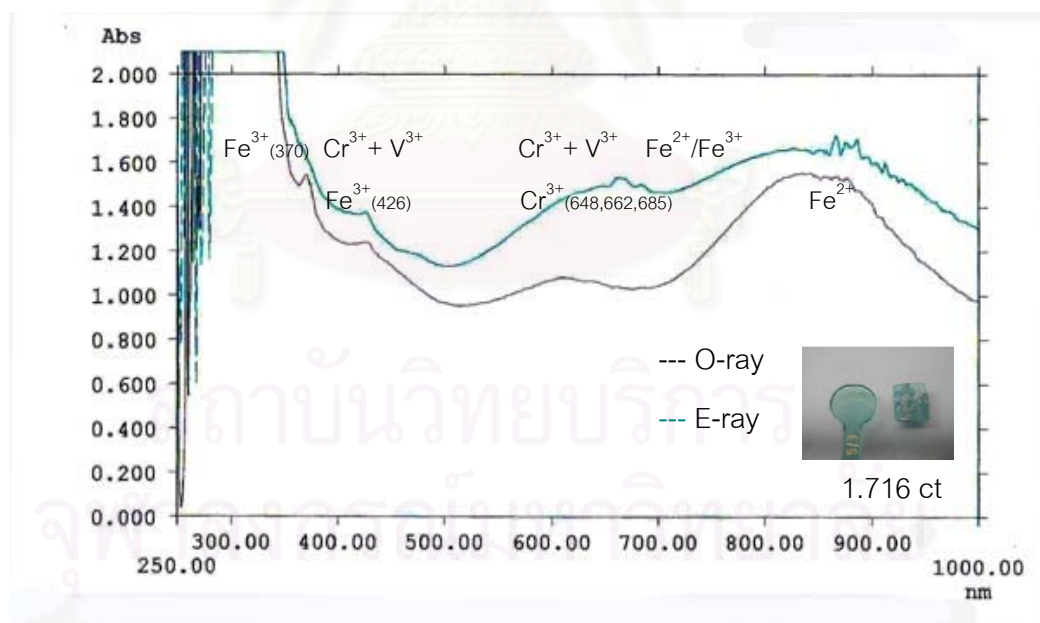


Figure 5.5: UV-Vis-NIR spectra of emerald from Madagascar (MA-28) representing a typical mixed spectrum containing both an emerald component with  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$  bands. The main  $\text{Cr}^{3+}$  absorption bands are overlapped by  $\text{Fe}^{3+}$  absorption.

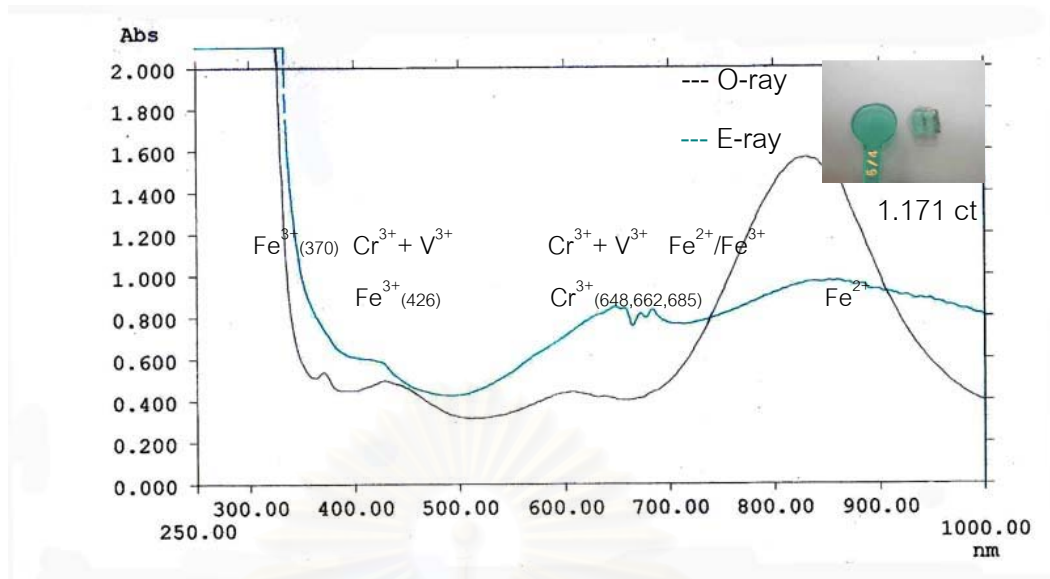


Figure 5.6: UV-Vis-NIR spectra of Zambian emerald (ZAM-29) showing a typical mixed spectrum containing both an emerald component with  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$  bands. The main  $\text{Cr}^{3+}$  absorption bands are overlapped by  $\text{Fe}^{3+}$  absorption.

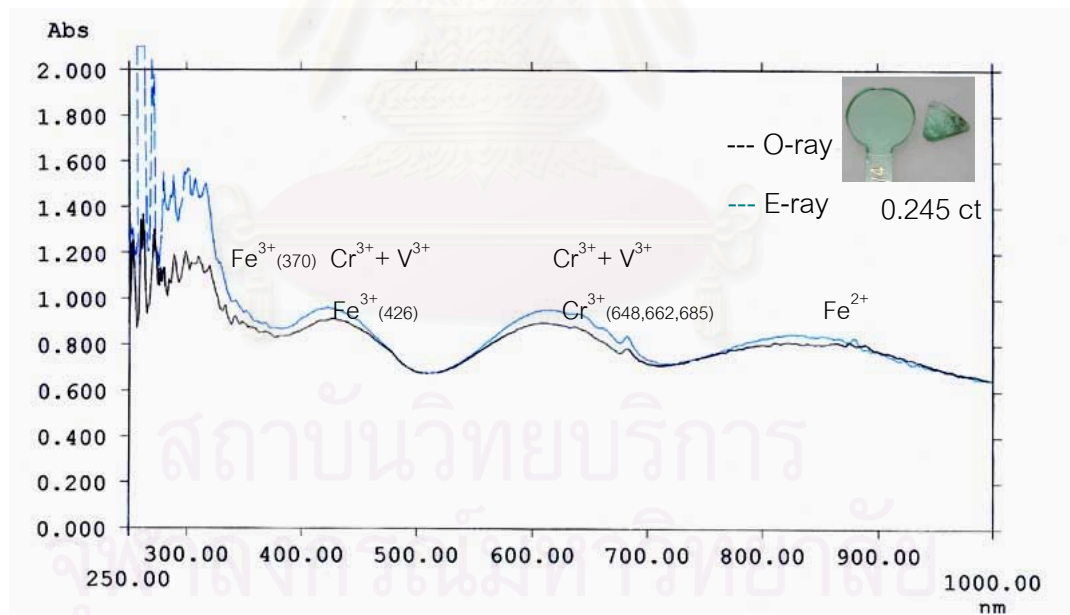


Figure 5.7: UV-Vis-NIR spectra of Carnaiba/Socoto emerald from Brazil (CAR-2) showing chromium-vanadium spectrum, ferric iron ( $\text{Fe}^{3+}$ ) bands in ultraviolet and ferrous iron ( $\text{Fe}^{2+}$ ) band in near-infrared.

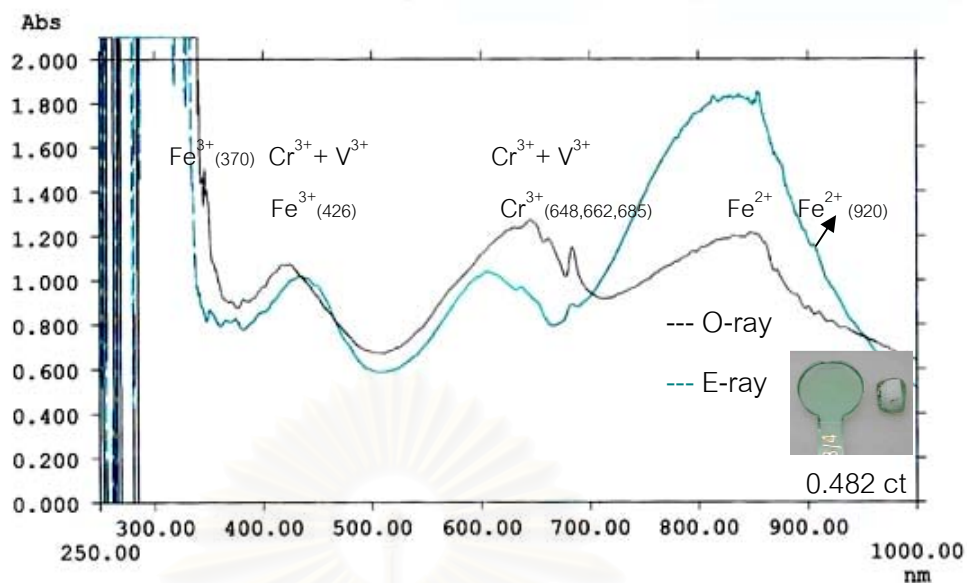


Figure 5.8: UV-Vis-NIR spectra of Itabira emerald from Brazil (ITA-2) showing chromium-vanadium spectrum, ferric iron ( $\text{Fe}^{3+}$ ) bands in ultraviolet and ferrous iron ( $\text{Fe}^{2+}$ ) band in near-infrared.

## 5.2 Fourier Transform InfraRed Spectrophotometry (FTIR)

The infrared region of the electromagnetic spectrum is the energy range just beyond the red end of the visible spectrum. In fact the term infrared is derived from being lower in energy 'infra-' than the red end. The unit by which infrared energy is usually measured is the wave number (number of waves per centimeter), which is expressed in reciprocal centimeters ( $\text{cm}^{-1}$ ). The infrared is thus referred to as the energy range between  $13,333 \text{ cm}^{-1}$  (the edge of the red) and  $33 \text{ cm}^{-1}$  (a limit determined by use and technology). This broad region is divided on the basis of experimental techniques and applications into three parts: near infrared ( $13,500 - 4,000 \text{ cm}^{-1}$ ), mid infrared ( $4,000 - 350 \text{ cm}^{-1}$ ), and far infrared ( $350 - 33 \text{ cm}^{-1}$ ). For most gemological purposes, infrared energy is expressed in  $\text{cm}^{-1}$ ; energies above  $400 \text{ cm}^{-1}$  that is, the mid infrared and the near infrared are of greatest interest gemologically.

Infrared spectroscopy in itself is not new, having become generally available to scientists about 50 years ago. However, technological advances in instrumentation in

the past 10 years have made infrared spectra much more readily and rapidly accessible. Infrared spectroscopy is now widely used in gem laboratories to characterize natural, treated and synthetic stones. It is a powerful tool for gem identification and research. Absorptions of a gem material in the infrared region of the electromagnetic spectrum are due to vibrations in the crystal structure; they can be used to help separate one gem material from another or to detect certain types of treatments. 'Water', either molecular ( $H_2O$ ) or as hydroxyl groups (OH) is combined in various forms in many gemstones or is present as an impurity. These various forms of water have characteristic patterns in the mid infrared and can be good indicators of structure, origin, or treatment

### 5.2.1 Instrument and Analytical condition

The author use Nicolet FTIR Spectrophotometer (Model NEXUS 670 with OMNIC software, see Figure 5.9) at GIT, the FTIR spectra were measured with the transmission mode and display as absorbance mode in the near infrared range at  $4000 - 8000 \text{ cm}^{-1}$  (wavenumber) in this study.

### 5.2.2 Results

The absorption bands of most emeralds appear in near infrared ranges due to the vibration of water molecules in the channel sites of the beryl structure. The assignment of each absorption band to two types of water molecules has previously been studied and are summarized in Tables 5.2 - 5.3 and shown in Figures 5.10 - 5.11, respectively (Schwarz and Henn, 1992; Schmetzer et al, 1997). The type I water exhibit absorption bands at  $5110, 5275, 5450, 5590, 6826, 7138$  and  $7267 \text{ cm}^{-1}$  whereas the type II water exhibits absorption bands at about  $5271$  and  $7077 \text{ cm}^{-1}$ . Especially for Nigerian emeralds show other three small peaks at  $4882, 4804$  and  $4650 \text{ cm}^{-1}$  that may be type I  $H_2O$ .



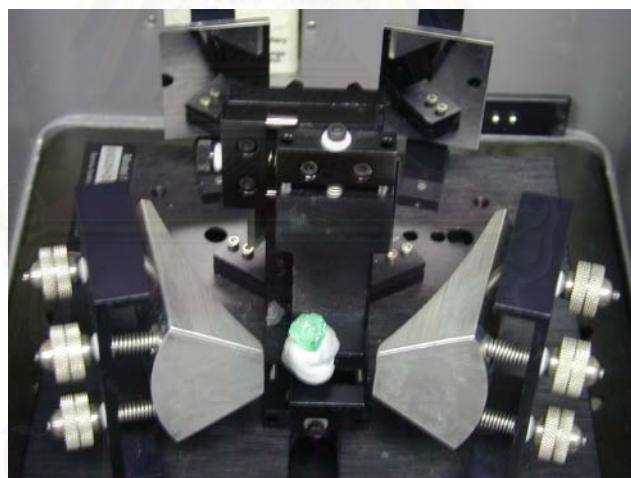


Figure 5.9: Nicolet FTIR Spectrophotometer (Model NEXUS 670) of GIT, the FTIR spectra were measured with the transmission mode in the near infrared.



Table 5.2: H<sub>2</sub>O vibration bands in the 4500 – 7500 cm<sup>-1</sup> range of emerald from Morafeno area, Manajary, Madagascar (Schwarz and Henn, 1992)

Wavenumber (cm <sup>-1</sup> )	Wavelength (nm)	Type H <sub>2</sub> O
5271	1897	II
5590	1789	I
6826	1465	I
7077	1413	II
7138	1401	I
7267	1376	I

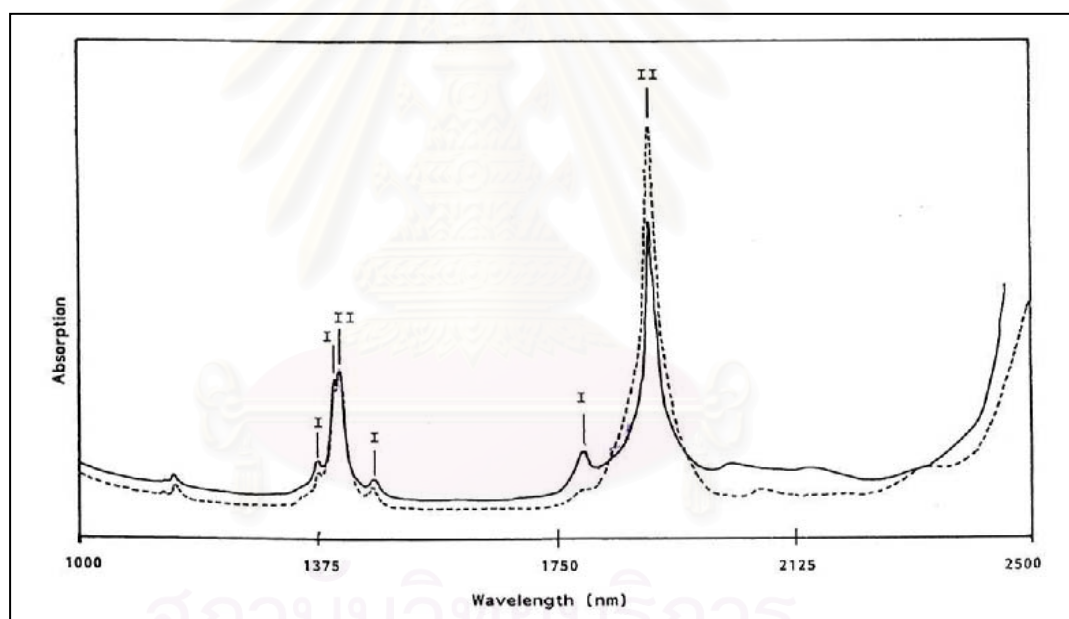


Figure 5.10: A representative FTIR spectra of vibration water molecules of emerald from Morafeno, Madagascar representing the type I H<sub>2</sub>O at 5590, 6826, 7138 and 7267 cm<sup>-1</sup> and the type II H<sub>2</sub>O at 5271 and 7077 cm<sup>-1</sup> (see table 5.2).

Table 5.3: H<sub>2</sub>O vibration bands in the 5000 - 6,000 cm<sup>-1</sup> range of synthetic hydrothermal emerald (Schmetzer et al, 1997) and show spectrum in Figure 5.11

Wavenumber (cm <sup>-1</sup> )	Type H <sub>2</sub> O
5110	I
5275	
5450	

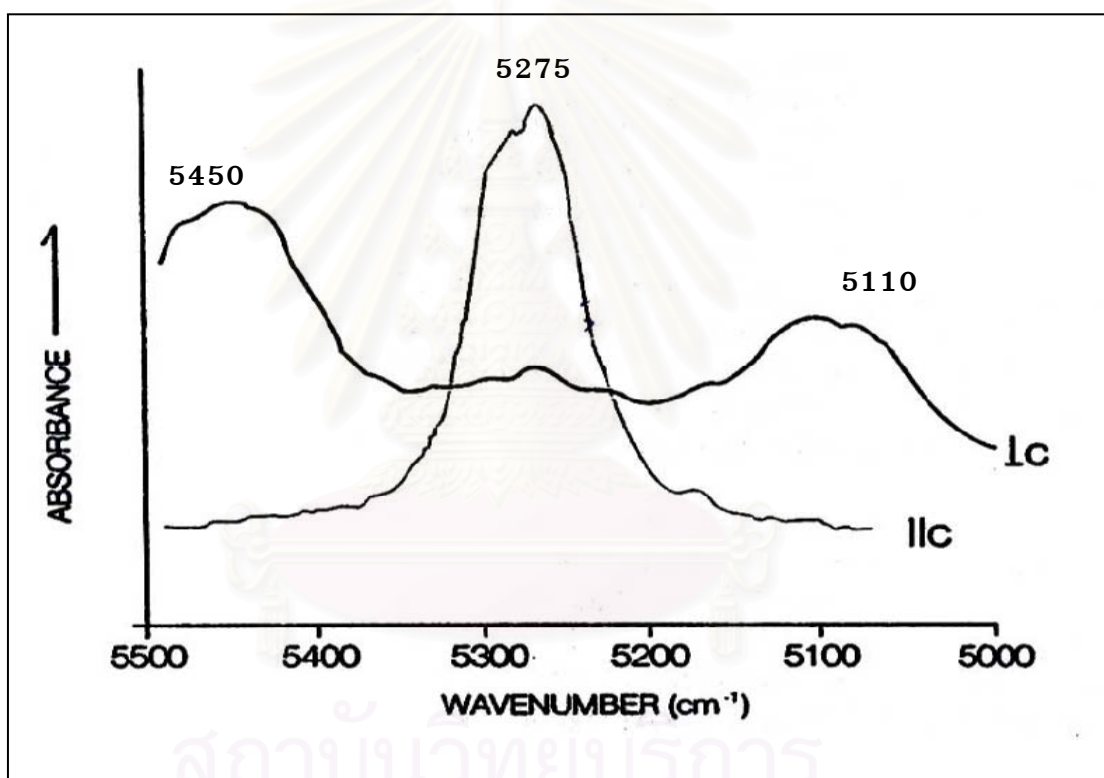


Figure 5.11: A sample of FTIR spectra in range 5000-5500 cm<sup>-1</sup> of vibration water molecules of Chinese hydrothermal synthetic emerald representing type I H<sub>2</sub>O when the beam is polarized parallel to the c-axis, the band at 5275 cm<sup>-1</sup> whereas the beam is polarized parallel to the c-axis, the band at 5110 and 5450 cm<sup>-1</sup>.

The representative FTIR absorption spectra of emeralds from each locality (Brazil, Madagascar, Zambia, Nigeria and Colombia) are shown in Figure 5.12 to 5.18.

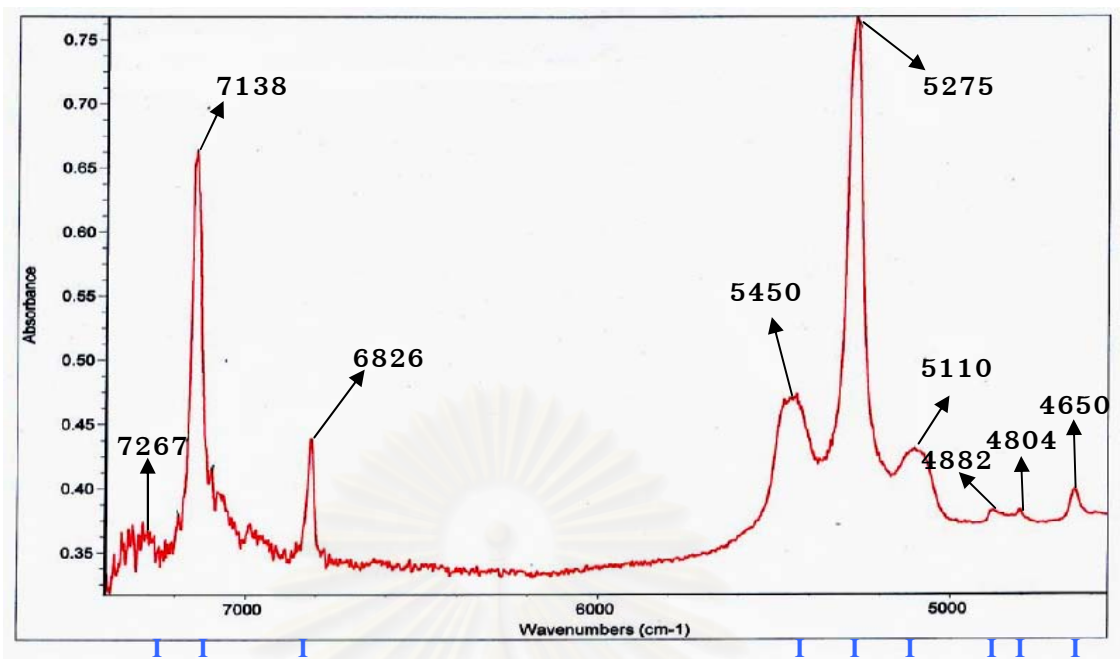


Figure 5.12: FTIR spectra of emerald from Nigeria (NI-24) dominated by type I H<sub>2</sub>O at 5110, 5450, 5275, 6826, 7138 and 7267 cm<sup>-1</sup> and three small peaks at 4882, 4804 and 4650 cm<sup>-1</sup> that may also be of type I H<sub>2</sub>O.

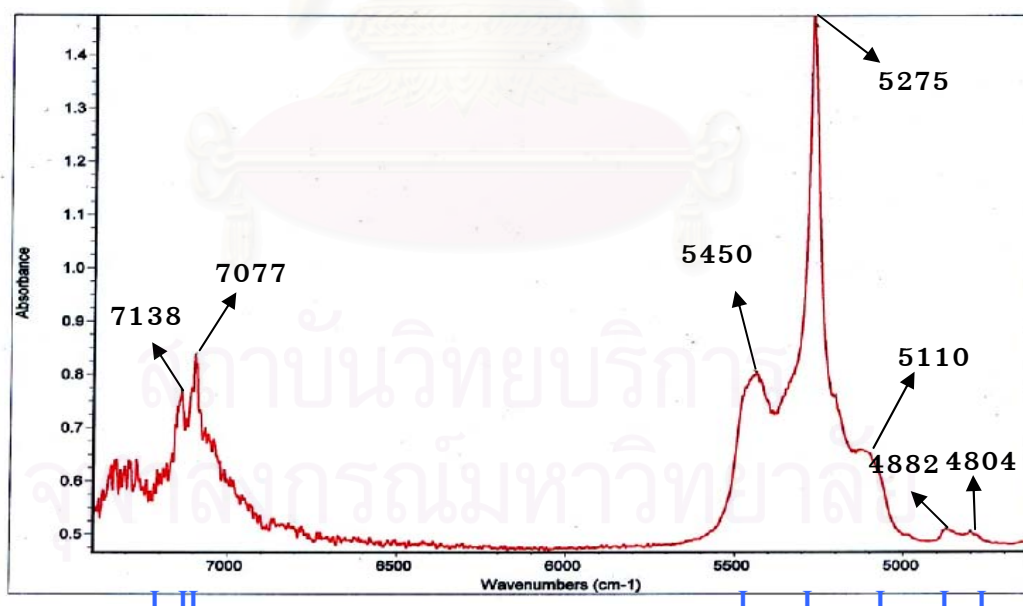


Figure 5.13: FTIR spectra of emerald from Colombia (CO-18) showing type I H<sub>2</sub>O at 5110, 5275, 5450 and 7138 cm<sup>-1</sup> and type II H<sub>2</sub>O at 7077 cm<sup>-1</sup>, that may also be of type I H<sub>2</sub>O at 4882 and 4804 cm<sup>-1</sup>.

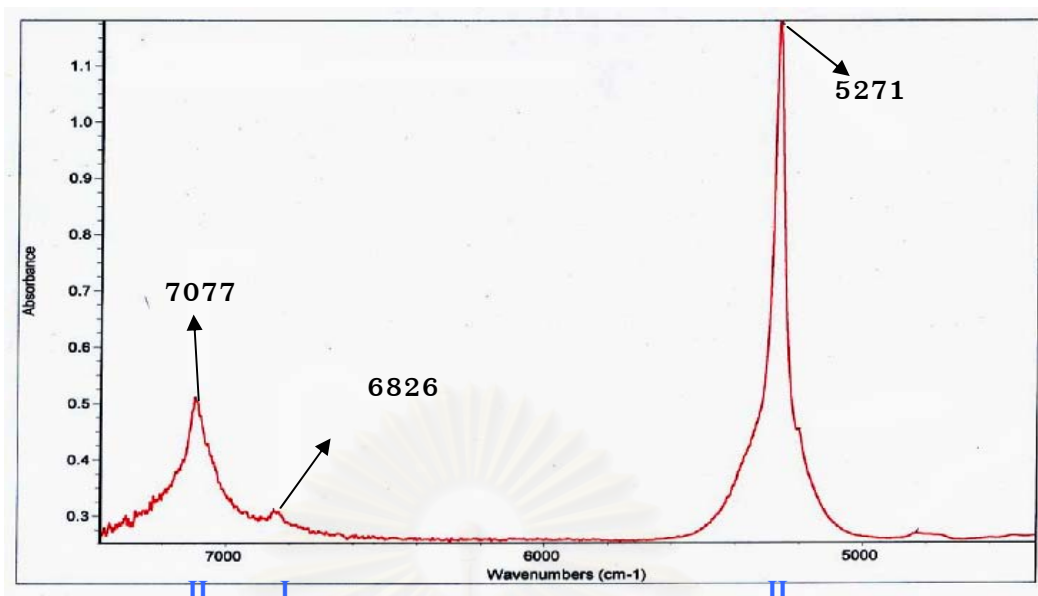


Figure 5.14: FTIR spectra of Santa Terezinha emerald from Goais, Brazil (STA-3) dominated by type II H<sub>2</sub>O at 5271 and 7077 cm<sup>-1</sup> while the type I H<sub>2</sub>O at 6826 cm<sup>-1</sup> is insignificant.

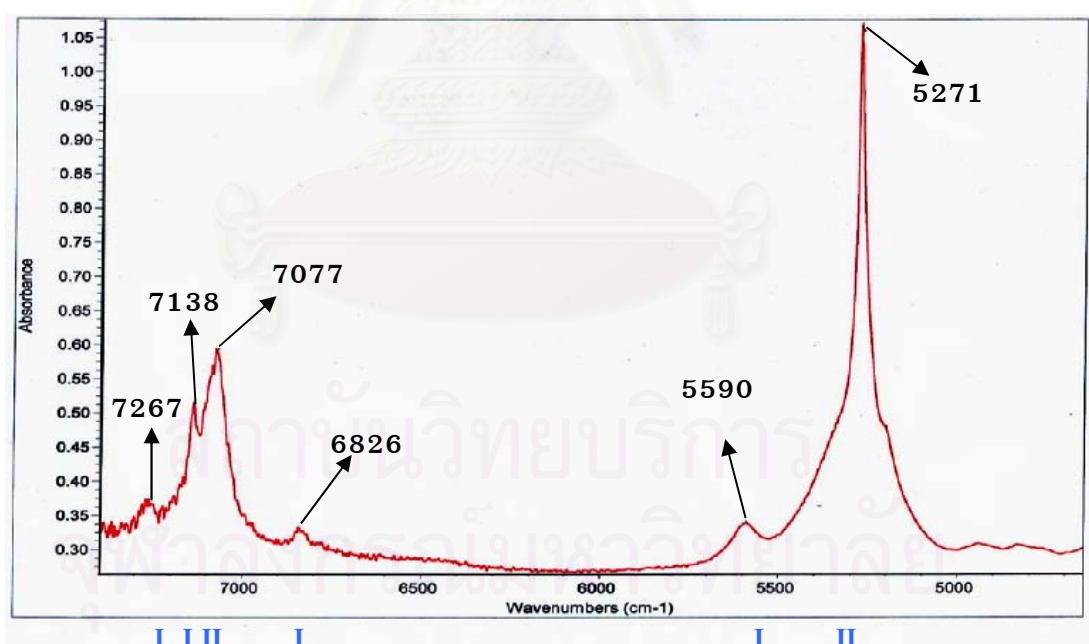


Figure 5.15: FTIR spectra of emerald from Mananjary, Madagascar (MA-21) showing type I H<sub>2</sub>O at 5590, 6826, 7138 and 7267 cm<sup>-1</sup> and type II H<sub>2</sub>O at 5271 and 7077 cm<sup>-1</sup>.

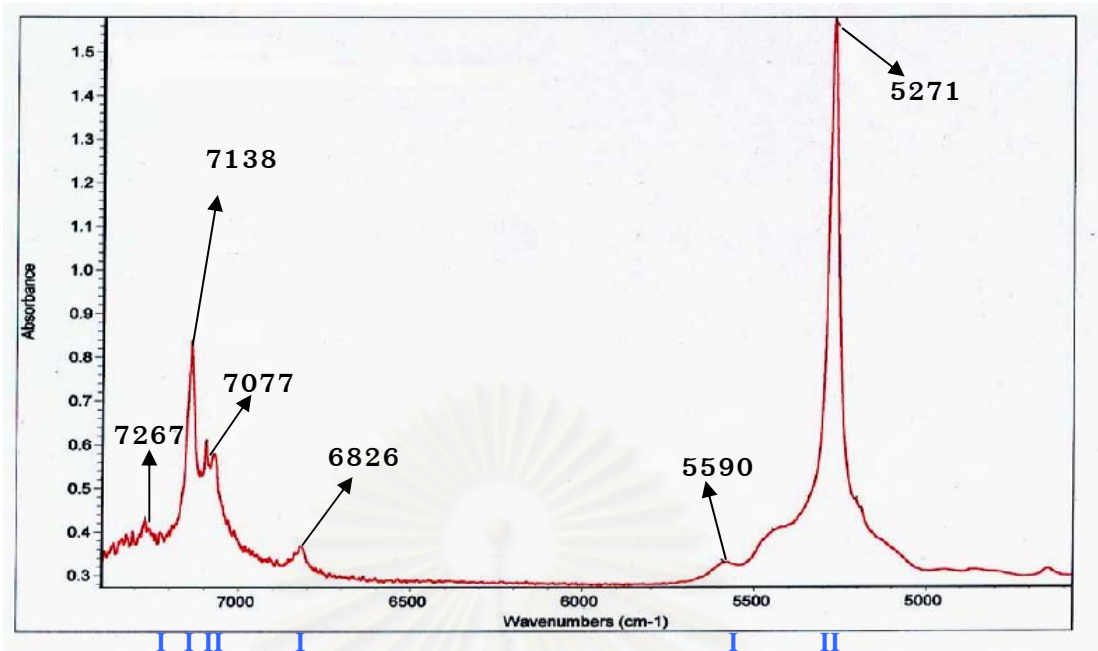


Figure 5.16: FTIR spectra of emerald from Zambia (ZA-3) showing type I H<sub>2</sub>O at 5590, 6826, 7138 and 7267 cm<sup>-1</sup> and type II H<sub>2</sub>O at 5271 and 7077 cm<sup>-1</sup>.

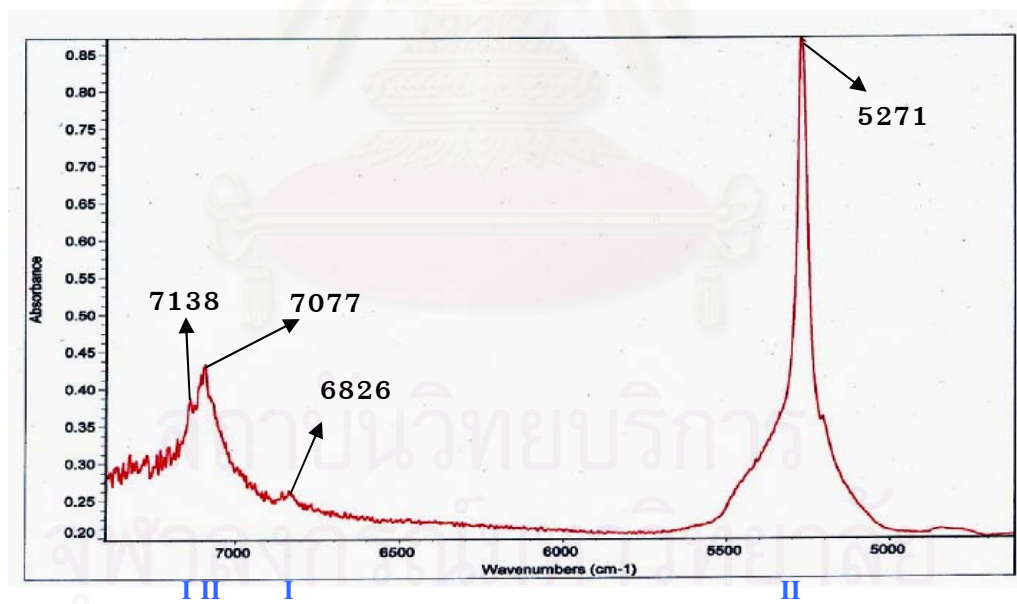


Figure 5.17: FTIR spectra of Canaiba/Socoto emerald from Bahia, Brazil (CAR-3) showing type I H<sub>2</sub>O at 6826 and 7138 cm<sup>-1</sup> and type II H<sub>2</sub>O at the 5271 and 7077 cm<sup>-1</sup>.



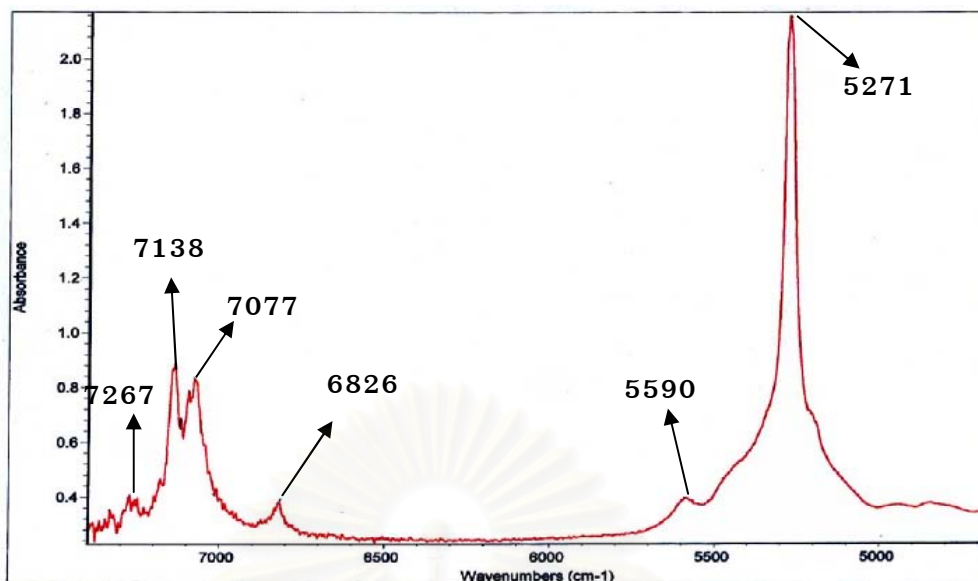


Figure 5.18: FTIR spectra of Itabira emerald from Minas Gerais, Brazil (ITA-7) showing type I H<sub>2</sub>O at 5590, 6826, 7138 and 7267 cm<sup>-1</sup> and type II H<sub>2</sub>O at 5271 and 7077 cm<sup>-1</sup>.

The spectra of almost all of these emeralds (i.e. from Madagascar, Zambia, Carnaiba/Socoto and Itabira of Brazil and Colombia) indicate similarly the presence of both types of water molecules except those of the Nigerian emerald and the Brazilian emerald from Santa Terezinha. The spectrum of Nigerian emerald is dominant by the type I water which is related to low alkali contents whereas that of the Brazilian emerald from Santa Terezinha is dominant by the type II water which is related to high alkali contents.

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## CHAPTER VI

### DISCUSSION AND CONCLUSION

#### 6.1 Discussion

##### 6.1.1 Geological Occurrences:

In this study, emerald deposits have been classified and categorized in two major specific geological occurrences which are summarized and displayed again in Table 6.1.

Table 6.1: Summary of geological occurrences of emeralds from Brazil, Colombia, Madagascar, Zambia and Nigeria

Type I Deposits (Non-schist-related)		Type II Deposits (Schist-related)				
Type Ia (Pegmatite without Schist)	Type Ib (Black Shales with Viens and Breccias)	Type IIa (Schist without Pegmatite)	Type IIb (Pegmatite with Schist); emerald occur between contact zone of country rock and pegmatite veins			
Emerald crystallized in vugs of granitic - rock, which intruded to the country rock.	Host rock is black shales.	Host rock is talc-chlorite-amphibole-magnetite schist.	Host rock is phlogopite-carbonate - talc schist.	Host rock is amphibolite.	Host rock is serpentinite.	Host rock is talc.
Kaduna Plateau, Nigeria	Cordillera Oriental, Colombia	Santa Terezinha (Brazil)	Mananjary, Madagascar	Ndola Rural, Zambia	Carnaiba/Socoto (Brazil)	Itabira (Brazil)

In this chapter the author will try to relate their geologic occurrences (type of deposits) with some physical properties and internal characteristics.

### 6.1.2 Physical and Optical properties:

Schwarz (2004) has classified emerald's refractive indices (e-ray and o-ray) and birefringence into low, medium and high values based on over 1,000 measurements from various localities as follows:

	<i>low</i>	<i>medium</i>	<i>high</i>
<i>e-ray</i>	<1.570	1.570-1.580	>1.580
<i>o-ray</i>	<1.580	1.580-1.590	>1.590
<i>birefringence</i>	<0.006	0.006-0.008	>0.008

From the refractive indices (o-ray) of emeralds measured in this study, the high values are found in emeralds from Santa Terezinha of Brazil and the medium to high values are found in emeralds from Madagascar, Zambia, Carnaiba/Socoto and Itabira of Brazil. These emeralds occur in geologically type II deposit. In contrast the low to medium values are found in emeralds from Colombia and Nigeria which occur in geologically type I deposit (Table 6.2).

The type II emerald deposits also show high specific gravity values whereas the type I emerald deposits give low values. The variation of SG values with the type of emerald deposits are consistent with those found in RI values (Table 6.2). As will be shown later that the variations of both RI and SG are related to the alkaline contents substitute in those emerald structures.

Table 6.2: Refractive indices (RI) of o-ray and specific gravity (SG) of emeralds from Brazil, Colombia, Madagascar, Zambia and Nigeria

Deposit type		Locality	RI (O-ray)	SG
Type I deposits (Non-schist related)	la	Nigeria (NI)	1.569 –1.582	2.601-2.687
	lb	Colombia (CO)	1.571 –1.582	2.612-2.731
Type II Deposits (Schist related)	IIa	Brazil; Santa Terezinha (STA)	1.592– 1.597	2.723-2.812
	IIb	Madagascar (MA)	1.581 – 1.598	2.661-2.785
	IIb	Zambia (ZA)	1.580 – 1.597	2.582-2.782
	IIb	Brazil; Carnaiba/Socoto (CAR)	1.581 –1.591	2.705-2.769
	IIb	Brazil; Itabira (ITA)	1.581 –1.591	2.690-2.727

### 6.1.3 Internal Characteristics:

The common internal features of emeralds found in this study are both solid inclusions and fluid inclusions. Among them mineral inclusions are the most common internal features which are summarized in Table 6.3.

Table 6.3: Summary of mineral and other solid inclusions in emeralds from Brazil, Colombia, Madagascar, Zambia and Nigeria

Mineral inclusions	Type I (Non-schist related)		Type II (Schist related)				
	Nigeria; Kaduna Plateau	Colombia; Cordilla Oriental	Santa Terezinha (Brazil)	Carnaiba/ Socoto (Brazil)	Itabira (Brazil)	Madagas car; Mananjary	Zambia; Ndola Rural
Pyrite		X	X				
Spinel			X				
Talc			X				
Carbonates	X	X	X			X	X
Mica	X		X	X	X	X	X
Albite		X					
Quartz			X		X	X	X
Hematite						X	X
Fluorite	X						
Chlorite				X			
Rutile	X						
Ilmenite	X						
Amphibole						X	X
Organic substance		X					

Inclusion features can provide evidence as to the geochemical environment in which the emerald crystals grew and they can also be used to distinguish emeralds from different locality. As would be expected, micas of probably biotitic or phlogopitic composition are commonly found in emeralds from the schist-related type II deposits of Brazil, Madagascar and Zambia, whereas muscovite may be found in Nigerian emerald

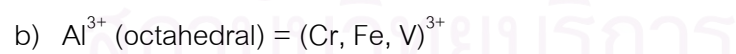
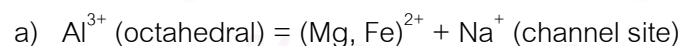


of the non-schist related type I deposits. In contrast Colombian emerald has no mica. Talc, quartz, chlorite, hematite and amphibole inclusions are not found in emeralds from both Nigeria and Colombia (type I; non-schist related). Along with the other characteristic mineral inclusions are dark organic substances, which are distinguished Colombian emerald from the others.

Fluid inclusions are generally found in emeralds from all deposits. However the characteristic 3-phase inclusions of emeralds are only occurred in emeralds from Colombia and Nigeria. Although the 3-phase inclusions from both deposits show jagged outline, but those of Columbia have lower liquid/vapor ratios than and can be distinguished from those of Nigeria (see Figure 3.74 for Nigeria and Figure 3.90 for Colombia).

#### 6.1.4 Trace element mineral chemistry:

Besides the major elements, such as Si, Al, O and Be that are formed an emerald, the other foreign elements, such as Cr, V and Fe (chromophores), sodium, magnesium, cesium (alkalis) etc. can be replaced or built into emerald/beryl lattice or structural channels (see Figure 3.3). Structural and chemical considerations indicated that the following substitutions in emerald could be applicable in chemical formula (Hänni et al., 1987), as a follows:



Divalent (doubly charged) ions substitution by alkalis such as  $\text{Mg}^{2+}$  mainly replace  $\text{Al}^{3+}$  on the lattice sites of structure of natural emerald while the larger sizes such as  $\text{Na}^+$  is built into the ring channels for balancing electric charge compensation. As seen in Tables 4.1 - 4.3 the major trace element substitution in those emerald structures are Na and Mg with subordinate Fe (except for the major Fe substitution in emeralds from Nigeria where there are only minor Na and Mg substitution) while other trace element substitutions are relatively minor. Therefore Na, Mg and probably Fe

substitution should have the largest influence on the emerald's optical and physical properties. As shown in the cross-plots of RI versus Na+Mg+Fe (Figure 6.1) and SG versus Na+Mg+Fe (Figure 6.2), a good correlation can be seen in both diagrams which suggest the dependence of both RI and SG on the Na+Mg+Fe substitution in those emerald structure.

As also shown in Figures 6.1 and 6.2 and mentioned earlier that both RI and SG are relatively high in the Type II emerald deposits and low in the Type I emerald deposits. This data may also suggest that the Type II schist-related deposits could inherently have strong influence from the alkali content during their formation as compared with those of the Type I non-schist-related deposits.

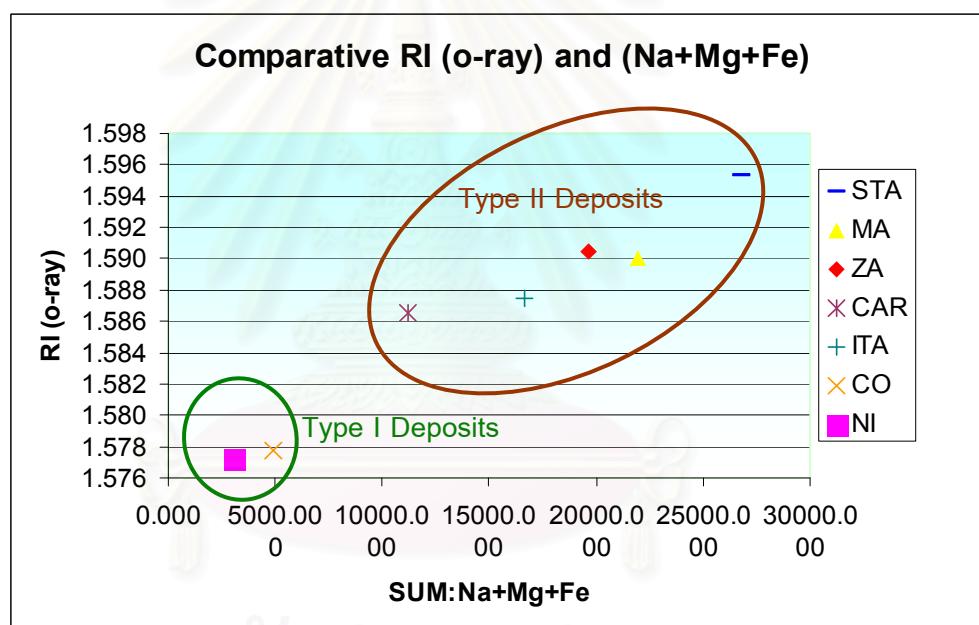


Figure 6.1: Correlation Diagram shows the highest optical values are found in Santa Tererzinha (STA) from Brazil, which also contain the highest values percentages of foreign elements whereas the lowest optical values are found in Nigeria (NI), also lowest foreign elements.

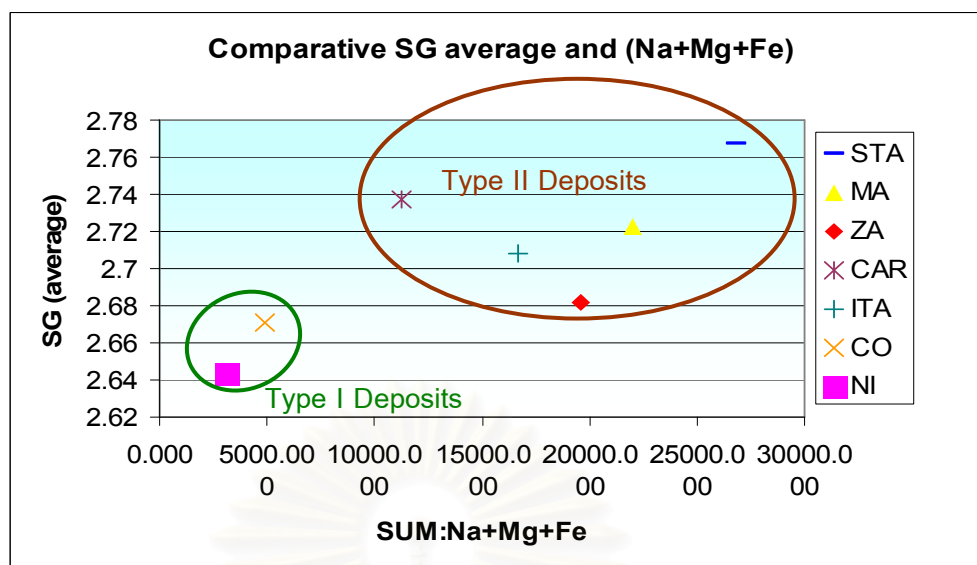
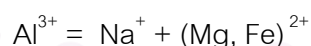


Figure 6.2: Correlation Diagram exhibits the highest specific gravity values are found in Santa Tererzinha (STA) from Brazil, which also contain the highest values percentages of foreign elements whereas the lowest specific gravity values are found in Nigeria (NI), also lowest foreign elements.

The diagram of Figure 6.3 shows that the rising of sodium plus magnesium and probably sodium plus iron plus magnesium in Figure 6.4, content of the emeralds is accompanied by a decreasing of the aluminum concentration. This variation suggest that the following substitution reaction for charge balance is likely to occur;



This can be confirmed by the strong correlation of Na vs Mg plot ( $R^2 = 0.95$ ) in Figure 6.5 and Na vs (Mg+Fe) plot ( $R^2 = 0.93$ ) in Figure 6.6. This assumption may imply that the majority of iron substitution in emeralds in this study could be in the form of divalent iron ( $\text{Fe}^{2+}$ ). The substitution of  $\text{Al}^{3+}$  by  $\text{Fe}^{3+}$  is probably less significant as there is no clear correlation of Al vs Fe plot in Figure 6.7 ( $R^2 = 0.45$ ). A good correlation of Al vs sum of all monovalent+divalent+trivalent ions contents ( $R^2 = 0.74$ ) also suggests that all the cations probably substitute for  $\text{Al}^{3+}$  in octahedral site rather than for  $\text{Be}^{2+}$  in tetrahedral site (Figure 6.8).

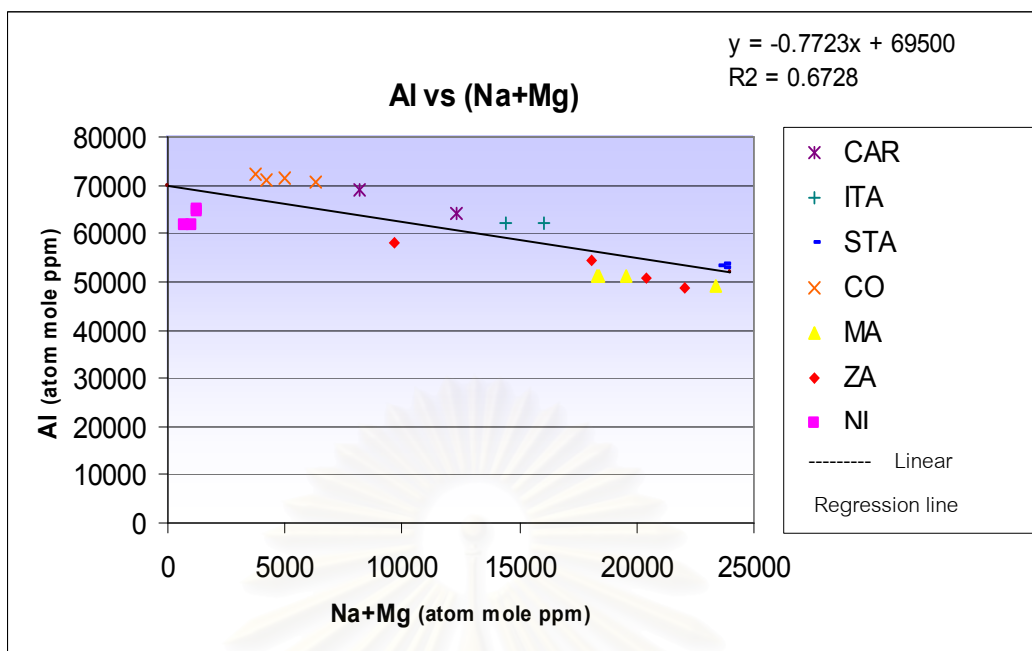


Figure 6.3: Plot of average Al versus Na+Mg contents of 23 emeralds from different geological occurrences.

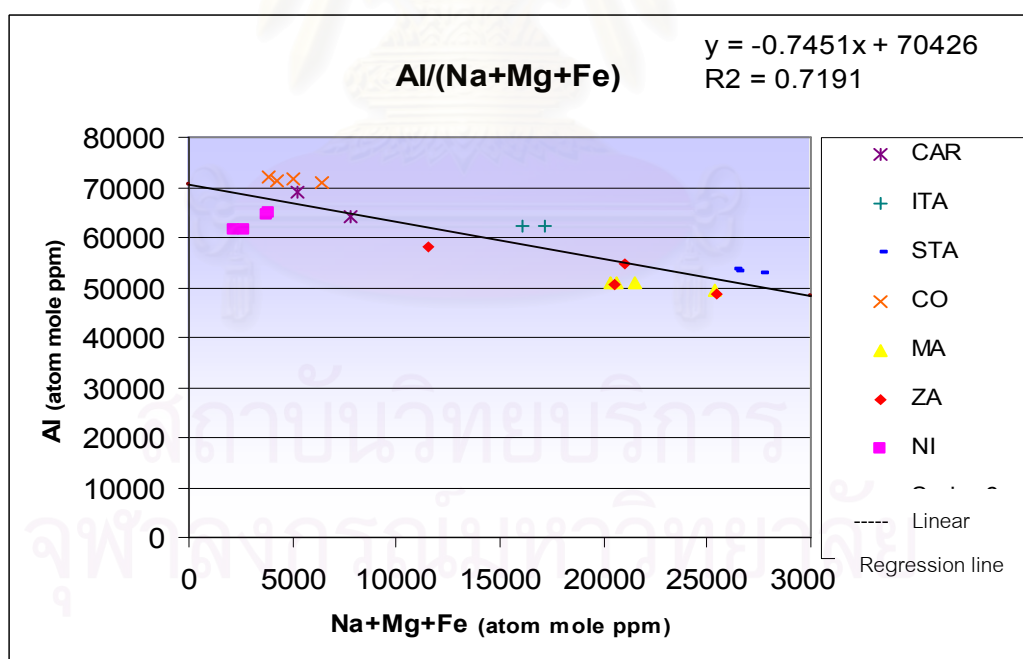


Figure 6.4: Plot of average Al versus Na+Fe+Mg contents of 23 emeralds from different geological occurrences.

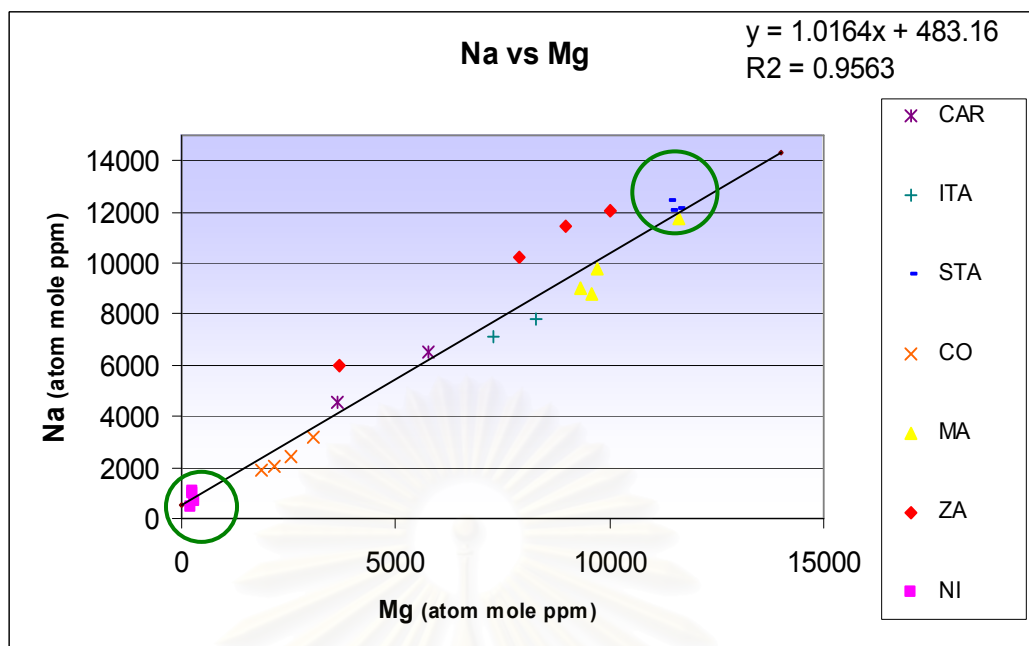


Figure 6.5: Plot of the average Na versus Mg contents of 23 emeralds from different geological occurrences.

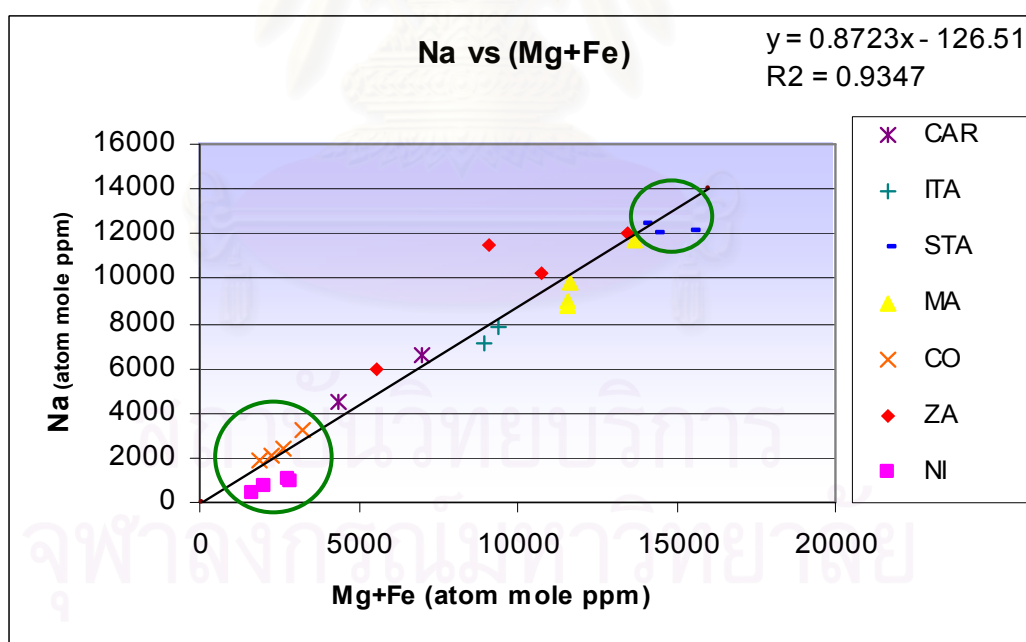


Figure 6.6: Plot of average Na versus (Mg+Fe) contents of 23 emeralds from different geological occurrences.



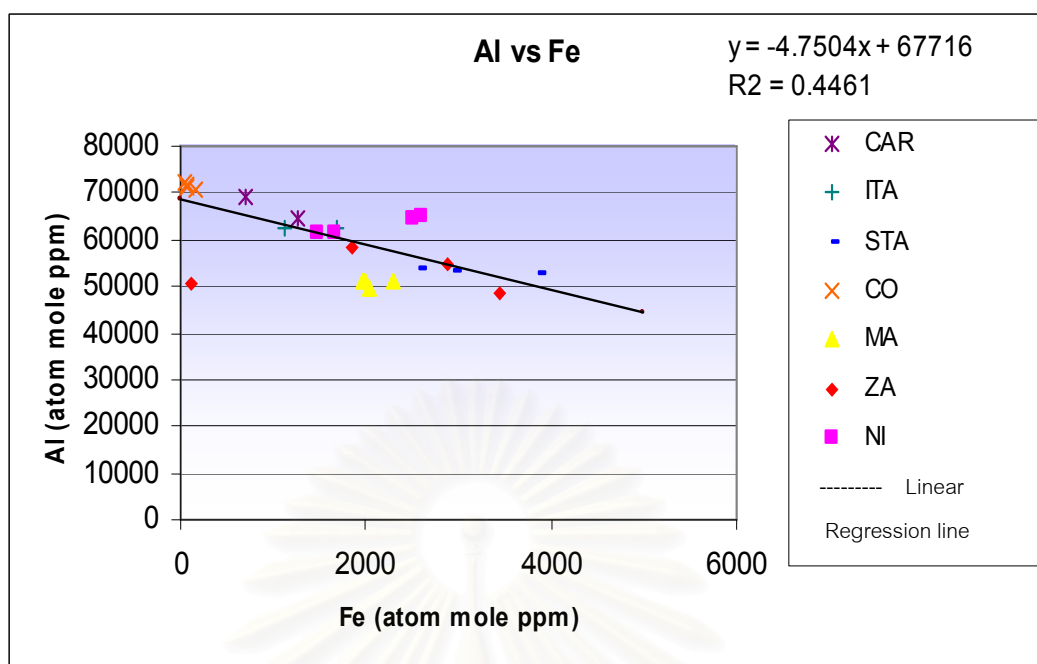


Figure 6.7: Plot of average Fe versus Al contents of 23 emeralds from different geological occurrences.

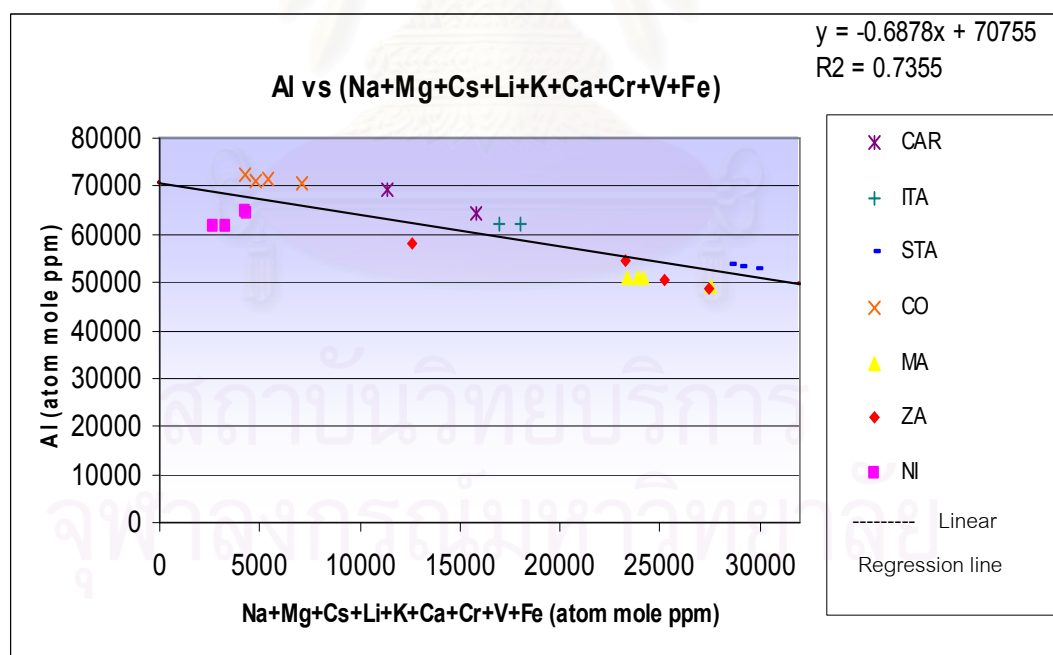


Figure 6.8: Plot of average Al versus sum of all Monovalent+Divalent+Trivalent ions contents from 7 different geological occurrences.

Figure 6.9 shows the relation of Cr and V. The emeralds from the Type II deposits (Santa Terezinha, Carnaiba/Socoto and Itabira, Zambia and Madagascar) contain higher Cr than V contents ( $Cr/V > 1$ ) whereas those of the Type I deposits (Colombia and Nigeria) have lower or approximately equal proportion of Cr and V contents ( $Cr/V < \text{or } \sim 1$ ). This data may imply that Cr might have been the important colouring element in the Type II schist-related deposits while both Cr and V are equally significant colouring elements in the Type I non-schist-related deposits.

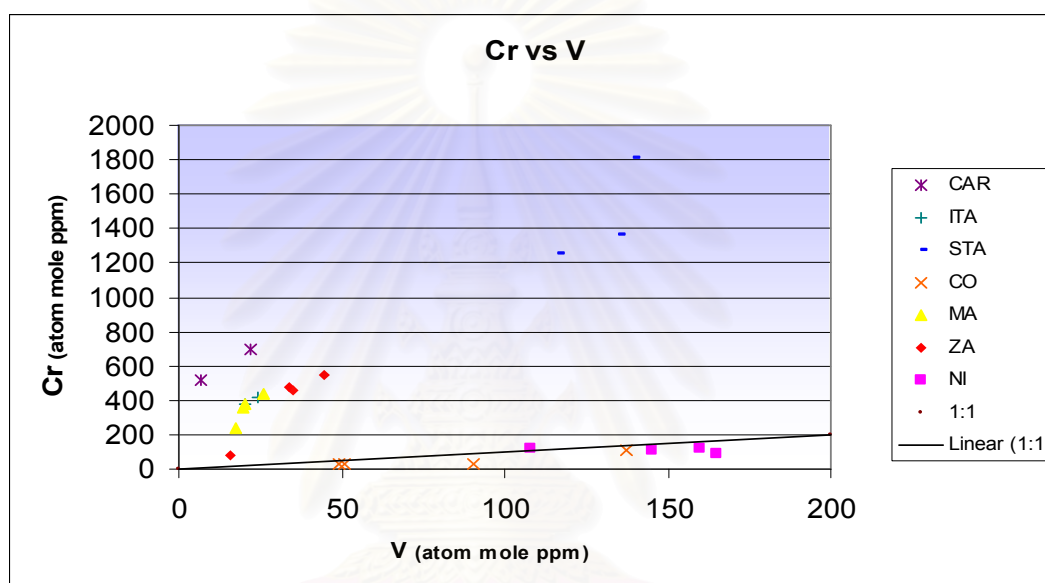


Figure 6.9: Plot of average Cr versus V contents of 23 emeralds from different geological occurrences.

#### 6.1.5 Spectroscopic properties:

Most UV-Vis-NIR absorption spectra of emeralds measured in this study revealed the typical absorption bands of  $Cr^{3+}$  and  $V^{3+}$  combined with  $Fe^{2+}$  and  $Fe^{3+}$ , except those of Colombian emeralds without the iron absorption bands. This observation is consistent with the trace element data which show rather low content of iron in the Colombian emeralds. Furthermore all the spectra show almost identical positions of the major absorption bands for  $Cr^{3+}$  and  $V^{3+}$ . However as shown in the trace element data, most emeralds in this study contain both chromium and vanadium in which the Type II

schist-related deposits are dominated mainly by Cr while the Type I non-schist-related deposits are dominated by both Cr and V. However, the respective portions of the ( $\text{Cr}^{3+}$  and  $\text{V}^{3+}$ ) absorption bands exhibit essential identical spectra which cannot be separated from one another. Consequently, the contribution of these two elements to their respective colour of an emerald is therefore likely to be very similar.

The infrared spectra of almost all of the emeralds examined in this study indicate similarly the presence of both type I and type II water molecules except those of the Nigerian emerald and the Brazilian emerald from Santa Terezinha. The spectrum of Nigerian emerald is dominant by the type I water which is a free water type and indicate low alkali contents. In contrast the spectrum of the Brazilian emerald from Santa Terezinha is dominant by the type II water which is the electric dipole water molecule oriented and bonded to neighboring alkali or other cations in the ring channels. Therefore the emerald dominated by type II water indicates high alkali contents. These results are consistent with the trace element data which show the lowest alkali content in the Nigerian emeralds and the highest content in the Brazilian emeralds from Santa Terezinha. The spectra of emeralds from Madagascar, Zambia, Carnaiba/Socoto and Itabira and Colombia show a mixture of type I and type II water which are related to intermediate alkali contents.

## 6.2 Conclusions

It is rather difficult to separate between emeralds from similar geologic environments based on gemological data. However, if specific characteristics of geological occurrences are present distinctly, it is possible to determine the origin of an emerald.

1. In this study, emerald deposits have been classified and categorized in two major specific geological occurrences, non-schist-related type I deposits and schist-related type II deposits.
2. Inclusion features can be used to distinguish emeralds from different locality. As would be expected from mineral inclusions, biotite or phlogopite are commonly found in

emeralds from the schist-related type II deposits but they not found in non-schist related type I deposits. Emeralds from Colombia and Nigeria contain different jagged outline in 3-phase inclusions in which those from Colombia have lower liquid/vapor ratios.

3. It is shown in this study that some foreign elements such as Na, Mg and Fe have a strong influence on RI and SG of those emeralds in which the type II schist-related deposits show relatively higher values (the highest values are found in emerald from Santa Terezinha) while the type I non-schist-related deposits have somewhat lower values (the lowest values are found in emeralds from Nigeria).

4. It is likely that the majority of iron substitution in emeralds in this study could be in the form of divalent iron ( $\text{Fe}^{2+}$ ). And all cations probably substitute for  $\text{Al}^{3+}$  in octahedral site rather than for  $\text{Be}^{2+}$  in tetrahedral site.

5. It was found that Cr might be the important colouring element in the type II schist-related deposits while both Cr and V have equally significant colouring elements in the type I non-schist-related deposits.

6. The emeralds from Zambia and Carnaiba/Socoto contain relatively high contents of Li and Cs, while the emeralds from Madagascar show exceptionally high K content. In the Colombian emeralds, both trace element data and UV-Vis-NIR spectra show consistently very low content of iron.

7. Nigerian emerald is dominated by the type I water and indicates low alkali contents. In contrast the Brazilian emerald from Santa Terezinha is dominated by the type II water which indicates high alkali contents.

8. These locality specific characteristics should be very useful for origin determination.

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APPENDICES

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## APPENDIX I

Figure I.1: GIA GemSet colour specimens; showing the range of emerald colouration in this study

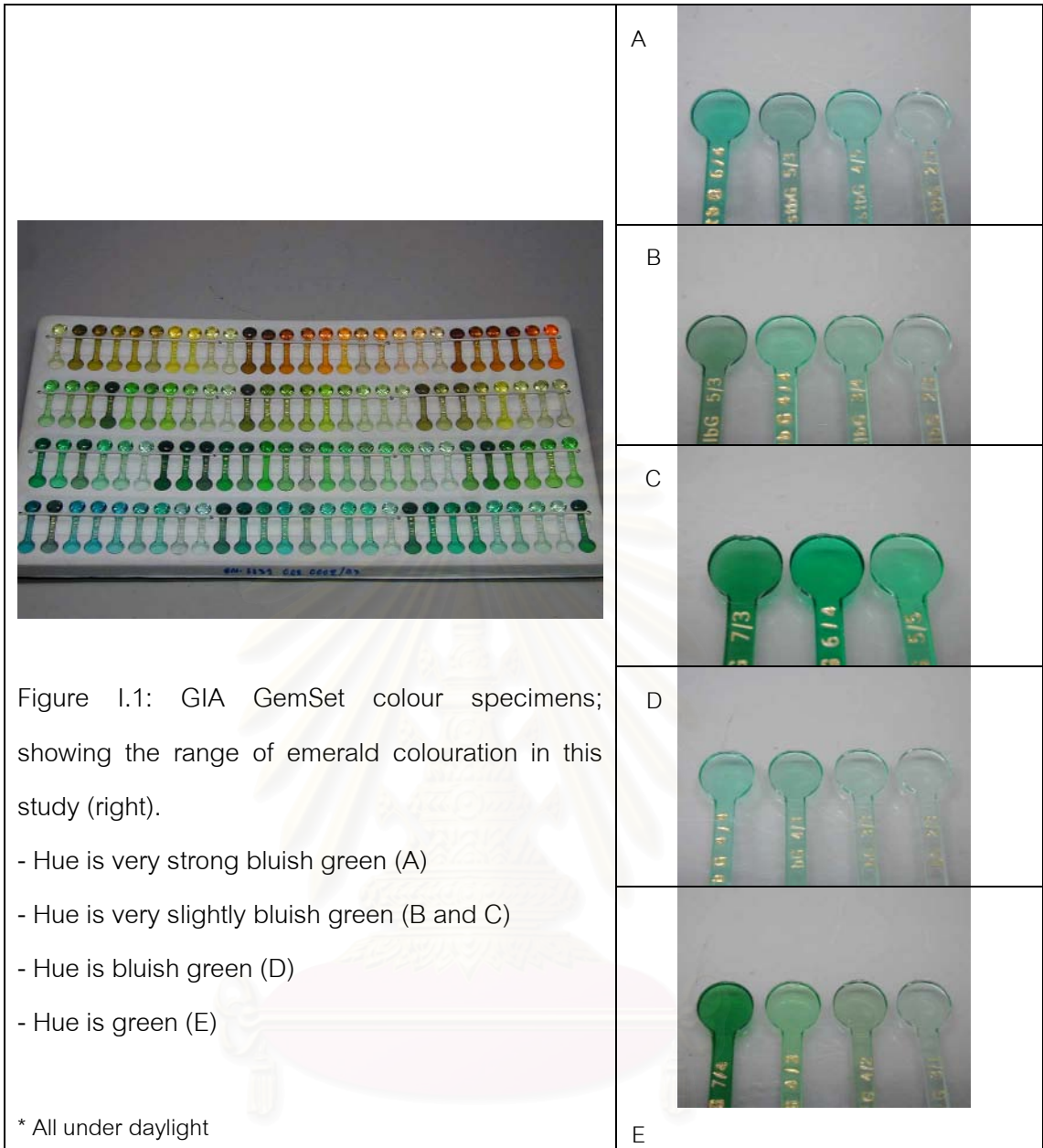
Table I.1: Detail of hue, tone and saturation terms for gem colors

Table I.2: Colour codes of emeralds in this study from different areas based on the GIA Gemset

Table I.3: Showing the gemological properties of 149 samples of emerald from Brazil, Colombia, Madagascar, Zambian and Nigeria

Table I.4: Major Raman Shift peaks of the mineral (William and other, 1997).

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Table I.1: List of hue, tone and saturation terms for describing gem colors, with their abbreviations.

Hue			
P	Purple	slyG	Slightly yellowish Green
Rp	reddish Purple	G	Green
RP/PR	red-purple or purple-red	VslbG	Very slightly bluish Green
slpR	slightly purplish Red	BG	Bluish Green
stpR	strongly purplish Red	VstbG	Very strongly bluish Green
R	Red	GB/BG	GREEN-BLUE or BLUE-GREEN
oR	Orangy red	VstgB	Very strongly greenish Blue
RO/OR	Red-orange-Orange-red	gB	Blue
rO	reddish Orange	vslgB	violetish Blue
O	Orange	B	bluish Violet
yO	yellowish Orange	vB	Violet
Oy	Orangy Yellow	bV	Bluish Purple
Y	Yellow	V	Pink
gY	greenish Yellow	bP	Brown
YG/GY	Yellow-Green or Green Yellow	Pk	
styG	strongly yellowish Green	Br	
yG	yellowish Green		

TONE		SATURATION	
0	colorless or white	1	grayish (brownish)
1	extremely light	2	slightly grayish (brownish)
2	very light	3	very slight grayish (brownish)
3	light	4	moderately strong
4	medium light	5	strong
5	medium	6	vivid
6	medium dark		
7	dark		
8	very dark		
9	extremely dark		
10	black		

Table 1.2: Colour codes of emeralds in this study from different areas based on the GIA Gemset

Origin/ Sample		GIA Gemset colour Code (Tone, Saturation, Hue)	
Nigeria (Kaduna Platua)	NI-14	vs1bG2/3	very light, very slightly greyish, very slightly bluish Green
	NI-8,9,11,26	vs1bG3/4	light, moderately strong, very slightly bluish Green
	NI-2,3,10,16,25	vs1bG4/4	medium light, moderately strong, very slightly bluish Green
	NI-13,17,30	bG2/3	very light, very slightly greyish, bluish Green
	NI-1,7,19,20,24	bG3/3	light, very slightly greyish, bluish Green
	NI-4,5,6,12,18,21,22,27,28, 29	bG4/3	medium light, very slightly greyish, bluish Green
	NI-15,23	G3/1	light, grayish, Green
	NI-14	vs1bG2/3	very light, very slightly greyish, very slightly bluish Green
	NI-8,9,11,26	vs1bG3/4	light, moderately strong, very slightly bluish Green

Origin	Sample	GIA Gemset colour	
		Code	(Tone, Saturation, Hue)
	NI-2,3,10,16,25	vslbG4/4	medium light, moderately strong, very slightly bluish Green
Colombia (Cordilla Oreintal- Coscués, Muzo)	CO-1,5,7,9,14,16,21,24	vslbG2/3	very light, very slightly grayish, very slightly bluish Green
	CO-6,22,29	vslbG3/4	light, moderate strong, very slightly bluish Green
	CO-10,15	vslbG4/4	medium light, moderate strong, very slightly bluish Green
	CO-23,26	vslbG5/3	medium, very slightly grayish, very slightly bluish Green
	CO-18,27,28	bG2/3	very light, very slightly greyish, bluish Green
	CO-2,12,13,17,19,25,30	bG3/3	light, very slightly greyish, bluish Green
	CO-8	bG4/4	medium light, moderately strong, bluish Green

Origin	Sample	GIA Gemset colour (Tone, Saturation, Hue)	
		Code	
	CO-3,4,11,20	G3/1	light, greyish, Green
Brazil: (Santa Terezinha in Goais State)	STA-2,3,4,5,8	vslbG3/4	light, moderately strong, very slightly bluish Green
	STA-7,9	vslbG4/4	medium light, moderately strong, very slightly bluish Green
	STA-1,6	vslbG7/3	dark, very slightly grayish, very slightly bluish Green
	STA-10	bG4/3	medium light, very slightly grayish, bluish Green
Madagascar (Manajary region)	MA-25	vstbG4/5	medium light, strong, very strongly bluish Green
	MA-24,28	vstbG5/3	medium, very slightly grayish, very strongly bluish Green
	MA-3,10	vslbG3/4	light, moderate strong, very slightly bluish Green
	MA-1,4,7,8,11,13,17,18,19, 22	vslbG4/4	medium light, moderate strong, very slightly bluish Green
	MA-2,5,6,26,29,30	vslbG5/5	medium, strong, very slightly bluish Green

Origin	Sample	GIA Gemset colour (Tone, Saturation, Hue)	
		Code	
	MA-14	bG3/3	light, very slightly greyish, bluish Green
	MA-9,12,15,16,20,21,23	bG4/3	medium light, very slightly grayish, bluish Green
	MA-15,27	bG4/4	medium light, moderate strong, bluish Green
Zambia (Ndola Rural district, Kafubu area	ZA-1	vstbG2/3	very light, very slightly grayish, very strongly bluish Green
	ZA-29	vstbG6/4	medium dark, moderate strong, very strongly bluish Green
	ZA-21	vslbG2/3	very light, very slightly greyish, very slightly bluish Green
	ZA-2,3,4,7,8,10,13,22	vslbG3/4	light, moderately strong, very slightly bluish Green
	ZA-9,17,30	vslbG4/4	medium light, moderately strong, very slightly bluish Green
	ZA-14	vslbG5/3	medium, very slightly greyish, very slightly bluish Green
	ZA-5,15,16,19,20,23	vslbG5/5	medium, strong, very slightly bluish Green



Origin	Sample	GIA Gemset colour (Tone, Saturation, Hue)	
		Code	
	ZA-24	vs1bG6/4	medium dark, moderately strong, very slightly bluish Green
	ZA-28	bG3/3	light, very slightly greyish, bluish Green
	ZA-6,11,12,18,26,27	bG4/3	medium light, very slightly greyish, bluish Green
	ZA-25	bG4/4	medium light, moderately strong, bluish Green
Brazil: (Carnaiba/ Socoto in Bahia State)	CAR-10	vs1bG2/3	very light, very slightly grayish, very slightly bluish Green
	CAR-2	vs1bG3/4	light, Moderately strong, very slightly bluish Green
	CAR-3,4	vs1bG4/4	medium light, moderately strong, very slightly bluish Green
	CAR-7	bG2/3	very light, very slightly grayish, bluish Green
	CAR-8	bG3/3	light, very slightly grayish, bluish Green

Origin	Sample	GIA Gemset colour (Tone, Saturation, Hue)	
		Code	
	CAR-1	G4/3	medium light, very slightly grayish, Green
	CAR-5,9	G3/1	light, grayish, Green
	CAR-6	G4/2	medium light, slightly grayish, Green
Brazil: (Itabira in Minas Gerais State)	ITA-2,3,5,8	vslbG3/4	light, moderately strong, very slightly bluish Green
	ITA-1,6	vslbG4/4	medium light, moderately strong, very slightly bluish Green
	ITA-7,9,10	bG4/3	medium light, very slightly grayish, bluish Green

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Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria

*Brazil (Carnaiba/Socoto)*

Sample no. (ct)	RI		Birefringence	SG Approximately	Fluorescence	
	ne	no			LWUV	SWUV
CAR1 (0.208)	1.581	1.590	0.009	2.77	inert	inert
CAR2 (0.245)	1.580	1.591	0.011	2.74	inert	inert
CAR3 (0.229)	1.580	1.589	0.009	2.76	inert	inert
CAR4 (0.303)	1.581	1.591	0.010	2.74	inert	inert
CAR5 (0.378)	1.572	1.581	0.009	2.70	inert	inert
CAR6 (0.405)	1.571	1.585	0.014	2.73	inert	inert
CAR7 (0.189)	1.571	1.582	0.011	2.73	inert	inert
CAR8 (0.308)	1.572	1.585	0.013	2.71	inert	inert
CAR9 (0.589)	1.575	1.584	0.005	2.72	inert	inert
CAR10 (0.264)	1.572	1.581	0.005	2.73	inert	inert

*Brazil (Itabira)*

ITA1 (0.342)	1.576	1.582	0.006	2.69	inert	inert
ITA2 (0.482)	1.579	1.586	0.007	2.70	inert	inert
ITA3 (0.852)	1.579	1.585	0.006	2.71	inert	inert
ITA5 (0.265)	1.570	1.581	0.010	2.69	inert	inert
ITA6 (0.178)	1.580	1.591	0.010	2.70	inert	inert
ITA7 (1.178)	1.580	1.589	0.009	2.70	inert	inert
ITA8 (0.713)	1.580	1.589	0.009	2.69	inert	inert
ITA9 (0.211)	1.580	1.589	0.009	2.72	inert	inert
ITA10 (0.292)	1.580	1.589	0.009	2.72	inert	inert

*Brazil (Santa Terezinha)*

STA1 (0.330)	1.583	1.596	0.013	2.72	inert	inert
STA2 (0.245)	1.586	1.598	0.012	2.80	inert	inert
STA3 (0.242)	1.585	1.595	0.010	2.74	inert	inert
STA4 (0.257)	1.579	1.593	0.014	2.73	inert	inert
STA5 (0.192)	1.586	1.597	0.011	2.81	inert	inert
STA6 (0.590)	1.582	1.595	0.013	2.76	inert	inert
STA7 (0.419)	1.585	1.592	0.012	2.76	inert	inert
STA8 (1.025)	1.583	1.595	0.012	2.72	inert	inert

Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria (*Brazil (Santa Terezinha)*)

Sample no. (ct)	RI		Birefringence	SG Approximately	Fluorescence	
	ne	no			LWUV	SWUV
STA9 (0.593)	1.586	1.596	0.010	2.77	inert	inert
STA10 (2.702)	1.582	1.594	0.012	2.76	inert	inert

*Colombia*

CO1 (1.286)	1.569	1.581	0.012	2.70	inert	inert
CO2 (0.059)	1.571	1.580	0.009	2.62	inert	inert
CO3 (0.458)	1.571	1.580	0.009	2.71	inert	inert
CO4 (0.524)	1.570	1.578	0.008	2.73	inert	inert
CO5 (0.959)	1.571	1.581	0.010	2.70	inert	inert
CO6 (0.826)	1.566	1.576	0.008	2.66	inert	inert
CO7 (1.714)	1.569	1.580	0.011	2.68	inert	inert
CO8 (1.190)	1.572	1.580	0.008	2.70	inert	inert
CO9 (0.498)	1.570	1.579	0.009	2.69	inert	inert
CO10 (1.519)	1.561	1.571	0.010	2.64	inert	inert
CO11 (1.806)	1.562	1.575	0.013	2.66	inert	inert
CO12 (2.104)	1.570	1.582	0.012	2.70	inert	inert
CO13 (1.832)	1.569	1.580	0.018	2.64	inert	inert
CO14 (1.207)	1.570	1.581	0.011	2.67	inert	inert
CO15 (2.070)	1.570	1.579	0.009	2.69	inert	inert
CO16 (1.117)	1.569	1.576	0.007	2.69	inert	inert
CO17 (1.368)	1.571	1.582	0.011	2.67	inert	inert
CO18 (1.764)	1.569	1.576	0.007	2.64	inert	inert
CO19 (2.110)	1.570	1.579	0.009	2.67	inert	inert
CO20 (2.427)	1.568	1.578	0.010	2.70	inert	inert
CO21 (0.863)	1.570	1.580	0.010	2.69	inert	inert
CO22 (2.312)	1.570	1.579	0.009	2.68	inert	inert
CO23 (4.423)	1.570	1.579	0.009	2.61	inert	inert
CO24 (3.208)	1.571	1.580	0.009	2.65	inert	inert
CO25 (1.012)	1.571	1.579	0.008	2.69	inert	inert
CO26 (1.352)	1.570	1.579	0.009	2.65	inert	inert
CO27 (0.574)	1.569	1.580	0.018	2.70	inert	inert

Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria (continued)

*Colombia*

Sample no.	RI		Birefringence	SG	Fluorescence	
	ne	no			LWUV	SWUV
CO28 (1.769)	1.567	1.575	0.008	2.69	inert	inert
CO29 (0.709)	1.570	1.579	0.009	2.71	inert	inert
CO30 (0.709)	1.570	1.580	0.010	2.70	inert	inert

*Madagascar*

MA1 (1.316)	1.572	1.581	0.009	2.74	inert	inert
MA2 (0.865)	1.584	1.594	0.010	2.71	inert	inert
MA3 (1.663)	1.581	1.590	0.009	2.74	inert	inert
MA4 (0.666)	1.580	1.591	0.011	2.69	inert	inert
MA5 (0.818)	1.584	1.593	0.009	2.75	inert	inert
MA6 (0.892)	1.581	1.591	0.010	2.71	inert	inert
MA7 (2.131)	1.580	1.590	0.010	2.70	inert	inert
MA8 (1.271)	1.580	1.589	0.009	2.71	inert	inert
MA9 (0.309)	1.582	1.591	0.009	2.73	inert	inert
MA10 (0.279)	1.581	1.580	0.009	2.66	inert	inert
MA11 (0.359)	1.570	1.589	0.010	2.66	inert	inert
MA12 (0.481)	1.576	1.591	0.013	2.72	inert	inert
MA13 (0.314)	1.582	1.590	0.008	2.75	inert	inert
MA14 (0.264)	1.582	1.591	0.009	2.77	inert	inert
MA15 (1.316)	1.580	1.590	0.010	2.72	inert	inert
MA16 (0.182)	1.579	1.591	0.012	2.78	inert	inert
MA17 (0.334)	1.581	1.591	0.010	2.66	inert	inert
MA18 (0.447)	1.580	1.590	0.010	2.71	inert	inert
MA19 (1.531)	1.582	1.598	0.016	2.71	inert	inert
MA20 (1.301)	1.582	1.592	0.010	2.72	inert	inert
MA21 (0.431)	1.580	1.591	0.011	2.68	inert	inert
MA22 (0.303)	1.581	1.590	0.009	2.74	inert	inert
MA23 (0.338)	1.584	1.593	0.009	2.66	inert	inert
MA24 (0.418)	1.575	1.598	0.023	2.74	inert	inert
MA25 (0.466)	1.580	1.590	0.010	2.73	inert	inert



Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria (continued)

*Madagascar*

Sample no.	RI		Birefringence	SG	Fluorescence	
	ne	no			LWUV	SWUV
MA26 (0.758)	1.579	1.592	0.013	2.73	inert	inert
MA27 (0.443)	1.584	1.596	0.012	2.71	inert	inert
MA28 (1.716)	1.580	1.590	0.010	2.70	inert	inert
MA29 (1.236)	1.580	1.590	0.010	2.71	inert	inert
MA30 (1.763)	1.584	1.595	0.011	2.72	inert	inert

*Zambia*

ZA1 (0.364)	1.580	1.591	0.011	2.74	inert	inert
ZA2 (7.753)	1.575	1.586	0.011	2.70	inert	inert
ZA3 (1.651)	1.578	1.584	0.006	2.72	inert	inert
ZA4 (1.899)	1.578	1.591	0.013	2.72	inert	inert
ZA5 (2.986)	1.581	1.591	0.010	2.72	inert	inert
ZA6 (3.988)	1.579	1.589	0.010	2.71	inert	inert
ZA7 (4.081)	1.580	1.589	0.009	2.71	inert	inert
ZA8 (8.351)	1.579	1.589	0.010	2.72	inert	inert
ZA9 (1.027)	1.578	-	-	2.73	inert	inert
ZA10 (0.847)	1.582	1.590	0.008	2.78	inert	inert
ZA11 (0.580)	1.578	1.586	0.008	2.68	inert	inert
ZA12 (0.365)	1.581	1.590	0.009	2.58	inert	inert
ZA13 (0.779)	1.581	1.590	0.009	2.69	inert	inert
ZA14 (0.384)	1.573	1.588	0.015	2.73	inert	inert
ZA15 (0.628)	1.574	1.589	0.015	2.70	inert	inert
ZA16 (0.551)	1.582	1.597	0.015	2.71	inert	inert
ZA17 (0.654)	1.584	1.592	0.008	2.74	inert	inert
ZA18 (0.557)	1.586	1.596	0.010	2.73	inert	inert
ZA19 (0.451)	1.578	1.598	0.016	2.70	inert	inert
ZA20 (2.447)	1.582	1.592	0.010	2.75	inert	inert
ZA21 (0.696)	1.575	1.588	0.013	2.70	inert	inert
ZA22 (1.091)	1.582	1.591	0.009	2.70	inert	inert
ZA23 (1.038)	1.579	1.591	0.012	2.72	inert	inert

Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria (continued)

*Zambia*

Sample no.	RI		Birefringence	SG	Fluorescence	
	ne	no			LWUV	SWUV
ZA24 (1.579)	1.580	1.589	0.009	2.74	inert	inert
ZA25 (0.589)	1.579	1.591	0.012	2.73	inert	inert
ZA26 (0.969)	1.580	1.590	0.010	2.73	inert	inert
ZA27 (0.823)	1.580	1.589	0.009	2.74	inert	inert
ZA28 (2.890)	1.578	1.590	0.012	2.71	inert	inert
ZA29 (1.171)	1.578	1.590	0.012	2.70	inert	inert
ZA30 (0.353)	1.580	1.590	0.010	2.72	inert	inert

*Nigeria*

NI1 (1.022)	1.562	1.578	0.016	2.61	inert	inert
NI2 (1.057)	1.561	1.575	0.014	2.65	inert	inert
NI3 (1.446)	-	1.572	-	2.64	inert	inert
NI4 (1.531)	1.562	1.570	0.008	2.64	inert	inert
NI5 (0.828)	1.560	1.573	0.013	2.66	inert	inert
NI6 (2.660)	1.560	1.571	0.011	2.66	inert	inert
NI7 (0.678)	1.565	1.572	0.007	2.66	inert	inert
NI8 (1.154)	1.561	1.570	0.009	2.62	inert	inert
NI9 (0.698)	1.568	1.575	0.007	2.63	inert	inert
NI10 (0.874)	1.576	1.582	0.016	2.66	inert	inert
NI11 (3.190)	1.569	1.579	0.010	2.61	inert	inert
NI12 (2.416)	1.568	1.575	0.007	2.67	inert	inert
NI13 (0.189)	1.561	1.569	0.008	2.60	inert	inert
NI14 (0.312)	1.568	1.576	0.008	2.64	inert	inert
NI15 (2.064)	1.569	1.578	0.009	2.63	inert	inert
NI16 (4.110)	1.570	1.579	0.009	2.64	inert	inert
NI17 (0.805)	1.561	1.571	0.001	2.65	inert	inert
NI18 (2.316)	1.563	1.571	0.008	2.60	inert	inert
NI19 (0.472)	1.567	1.580	0.013	2.61	inert	inert
NI20 (0.606)	1.568	1.577	0.009	2.66	inert	inert
NI21 (0.602)	1.569	1.582	0.013	2.66	inert	inert

Table I.3: List the gemological properties of emeralds from Brazil, Colombia, Madagascar, Zambian and Nigeria (continued)

*Nigeria*

Sample no.	RI		Birefringence	SG	Fluorescence	
	ne	no			LWUV	SWUV
NI22 (1.850)	1.565	1.573	0.008	2.65	inert	inert
NI23 (0.612)	1.562	1.570	0.008	2.65	inert	inert
NI24 (0.848)	1.560	1.570	0.010	2.66	inert	inert
NI25 (0.877)	1.562	1.571	0.009	2.65	inert	inert
NI26 (1.108)	1.563	1.571	0.008	2.60	inert	inert
NI27 (1.640)	1.563	1.573	0.010	2.68	inert	inert
NI28 (1.068)	1.567	1.577	0.010	2.67	inert	inert
NI29 (4.083)	1.565	1.574	0.009	2.67	inert	inert
NI30 (0.813)	1.569	1.578	0.009	2.68	inert	inert

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Table I.4: Major Raman Shift peaks of the mineral inclusions in this study (William et al., 1997)

Minerals	Major Raman Shift peaks at 200 – 2000 cm <sup>-1</sup>
Actinolite	383, 665, 669, 1066
Albite	279, 472, 510, 1126
Apatite	433, 434, 593, 965, 966, 1056
Barytocalcite	227, 228, 264, 703, 1087
Beryl	323, 324, 398, 399, 402, 451, 526, 529, 532, 686, 687, 690, 691, 1012, 1013, 1019, 1060, 1070, 1072, 1075, 1240
Biotite	204, 359, 362, 406, 413, 543, 549, 552, 556, 557, 563, 714, 715, 752, 756, 782, 787, 789, 802, 805, 913, 1070, 1073, 1082, 1287
Calcite	278, 283, 285, 714, 716, 1083, 1087, 1088, 1434, 1437, 1746
Chlorite	203, 352, 361, 549, 676, 678, 679, 1081
Chromite	222, 538, 549, 573, 583, 586, 658, 671, 700, 703, 752, 1580
Dolomite	300, 301, 302, 724, 725, 727, 1097, 1099, 1100, 1440, 1442, 1444, 1758, 1760
Emerald	315, 319, 322, 323, 393, 394, 395, 397, 398, 681, 683, 1033, 1066, 1067, 1068, 1069, 1071
Feldspar	212, 285, 286, 293, 455, 459, 476, 478, 481, 510, 514, 766, 817, 1101
Fluorite	320, 837, 840, 841, 913, 920, 926, 929, 946, 947, 950, 1031, 1125, 1154, 1253, 1293, 1294, 1297, 1358, 1360, 1364, 1391, 1395, 1398, 1449, 1466, 1505, 1576, 1580
Hematite (Fe <sub>2</sub> O <sub>3</sub> )	202, 227, 230, 248, 250, 260, 292, 295, 298, 303, 350, 392, 412, 413, 416, 417, 499, 500, 504, 615, 617, 666, 675, 804, 1078, 1157, 1321, 1324, 1337, 1342
Magnetite	313, 555, 668, 680, 729, 738, 739,
Phlogopite	282, 331, 565, 685, 746, 785, 1091
Pyrite	347, 352, 381, 388, 433, 443
Pyroxene	230, 321, 323, 386, 387, 506, 526, 559, 664, 666, 858, 1012, 1014

Table I.4: William et al. (1997), Major Raman Shift peaks of the mineral inclusions in this study  
(continued).

Spinel	402, 664, 718, 762, 763, 803, 930
Rutile	237, 238, 240, 242, 441, 448, 449, 610, 612, 613, 800
Sulpher / Sulphide	217, 245, 437, 472
Talc	289, 362, 434, 468, 678, 793
Quartz	204, 207, 208, 214, 264, 265, 354, 356, 391, 398, 412, 462, 465, 466, 467, 505, 693, 697, 799, 813, 1159, 1163, 1164



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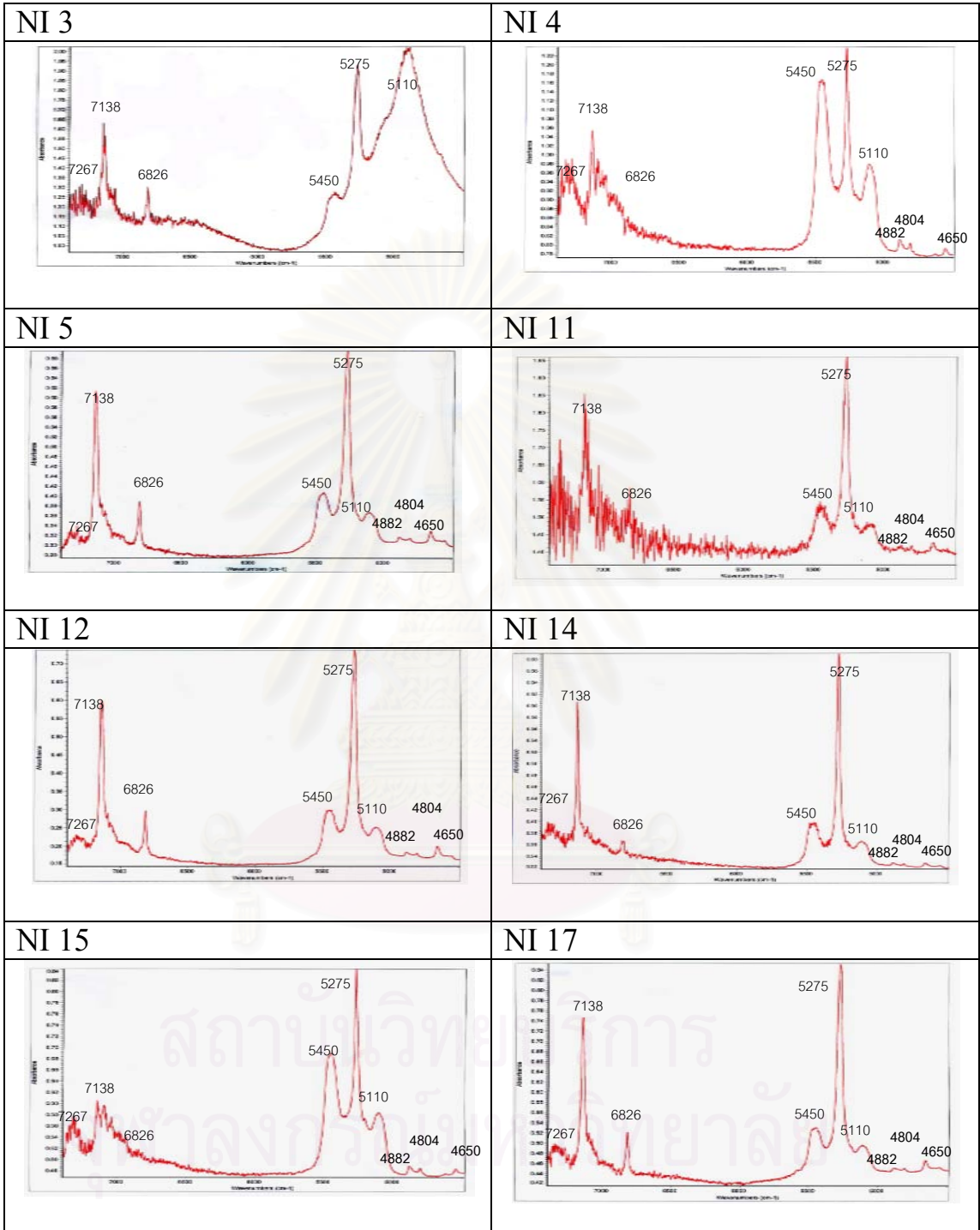


## APPENDIX II

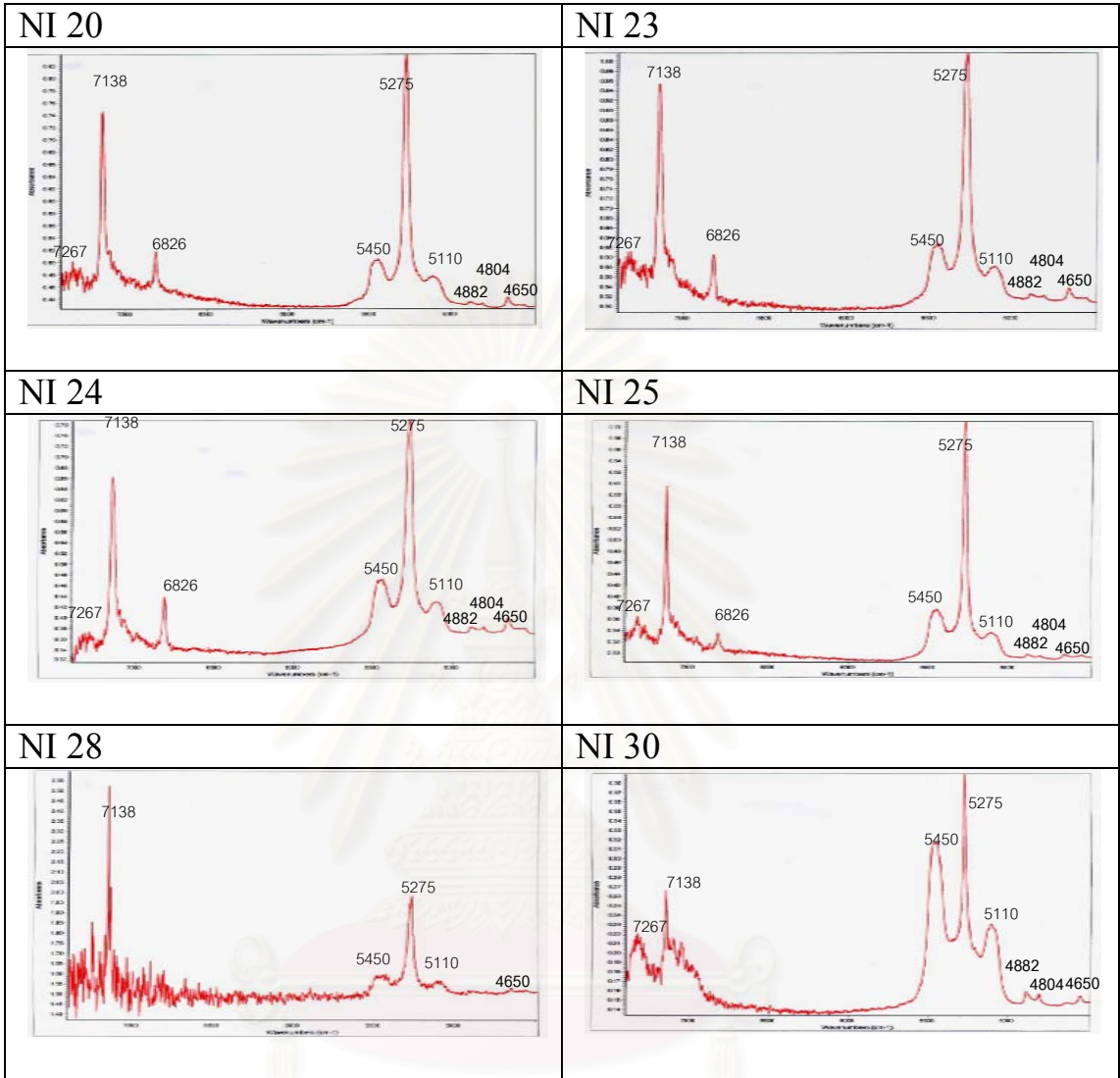
FTIR spectra of emerald samples from Nigeria, Colombia, Brazil, Madagascar and Zambia;

- 15 samples of Kaduna Plateau emeralds from Nigeria: NI3, NI4, NI5, NI11, NI12, NI14, NI15, NI17, NI10, NI20, NI23, NI24, NI25, NI28, NI30
- 12 samples of Cordillera Oriental emeralds from Colombia: CO1, CO2, CO4, CO6, CO8, CO11, CO13, CO18, CO20, CO22, CO26, CO27
- 4 samples of Santa Terezinha from Goais State, Brazil: SAN3, SAN7, SAN8, SAN10
- 14 samples of Manajary emeralds from Madagascar: MA1, MA3, MA9, MA15, MA17, MA18, MA21, MA23, MA24, MA26, MA27, MA28, MA30
- 14 samples of emeralds from Ndola Rural district: ZA3, ZA4, ZA5, ZA11, ZA12, ZA14, ZA15, ZA17, ZA20, ZA23, ZA24, ZA25, ZA28, ZA30
- 5 samples of Carnaiba/Socoto emeralds from Bahia State, Brazil: CAR2, CAR3, CAR4, CAR8, CAR9
- 5 samples of Itabira emeralds from Minas Gerias State, Brazil: ITA1, ITA2, ITA3, ITA7, ITA9

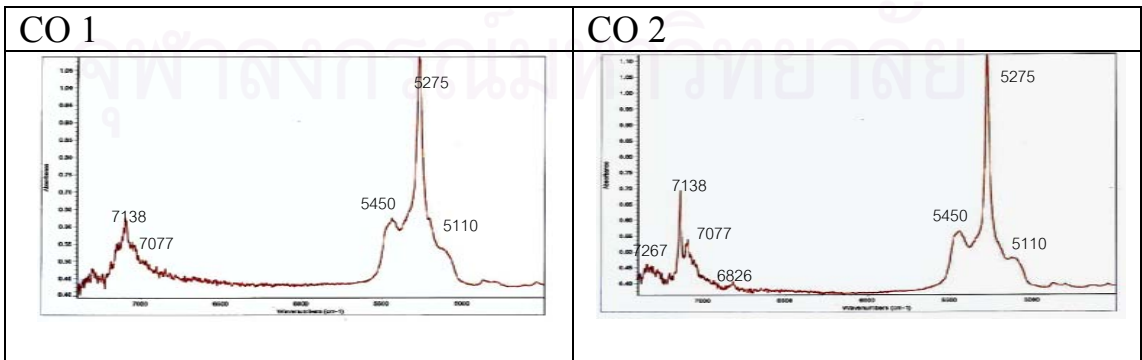
FTIR spectra of emerald samples from Nigeria: *Pure Type I H<sub>2</sub>O (Ia)*



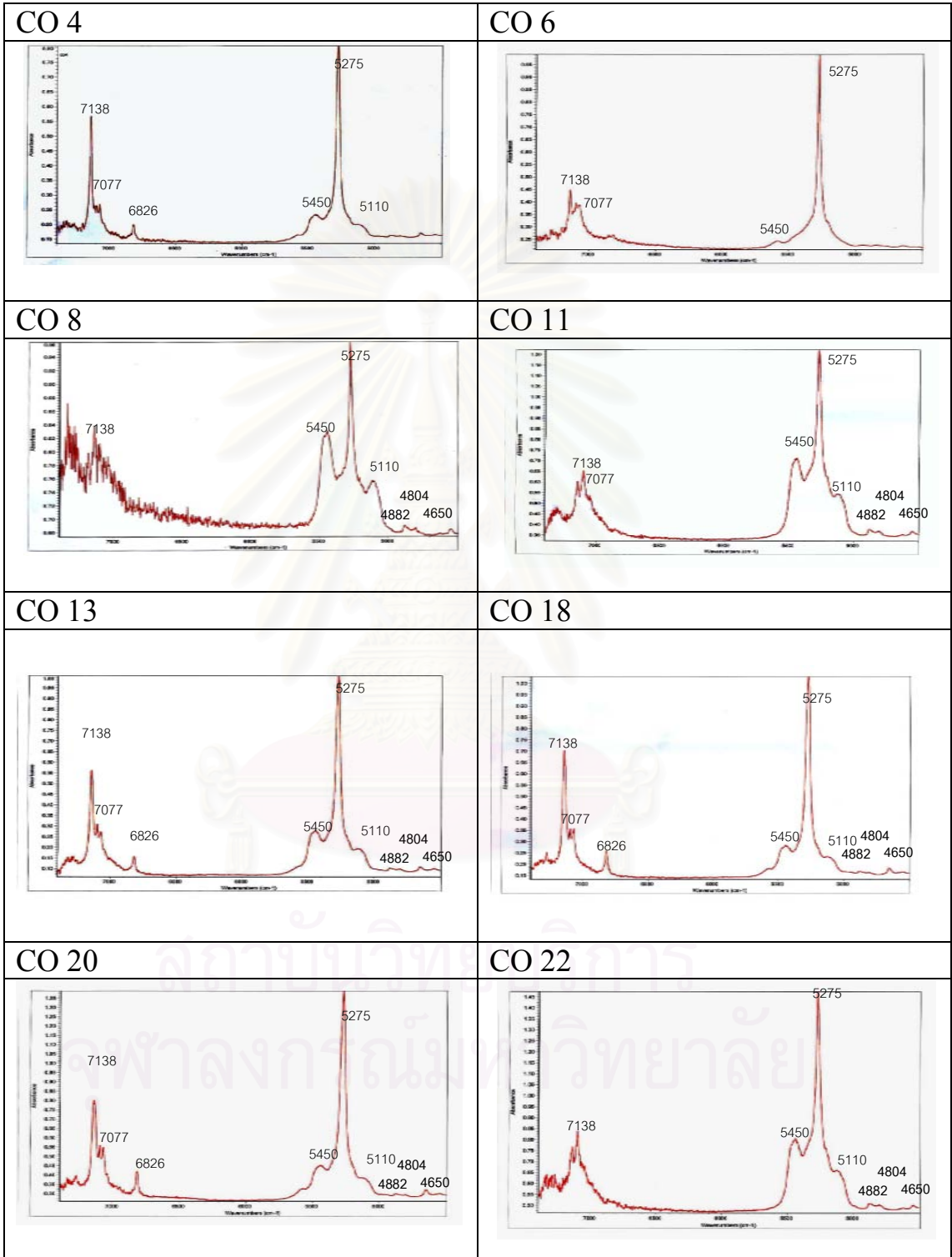
FTIR spectra of emerald samples from Nigeria: *Pure Type I H<sub>2</sub>O (Ia), continue*



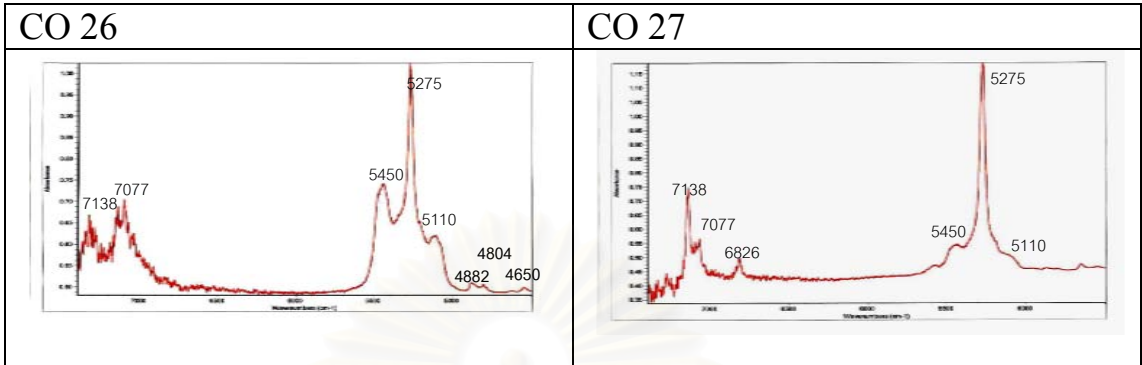
FTIR spectra of emerald samples from Colombia: *Type I H<sub>2</sub>O dominant + II H<sub>2</sub>O (Ib)*



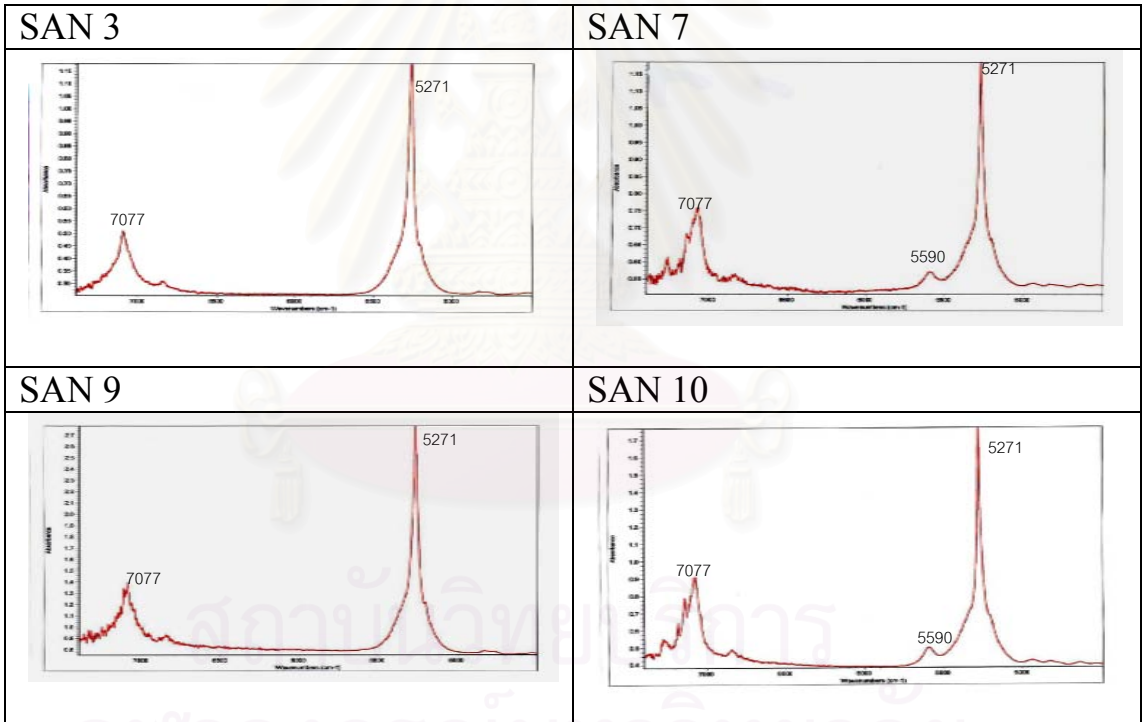
FTIR spectra of emerald samples from Colombia: *Type I H<sub>2</sub>O dominant + II H<sub>2</sub>O (lb)*,  
 continue



FTIR spectra of emerald samples from Colombia: *Type I H<sub>2</sub>O dominant + II H<sub>2</sub>O (Ib)*,  
 continue

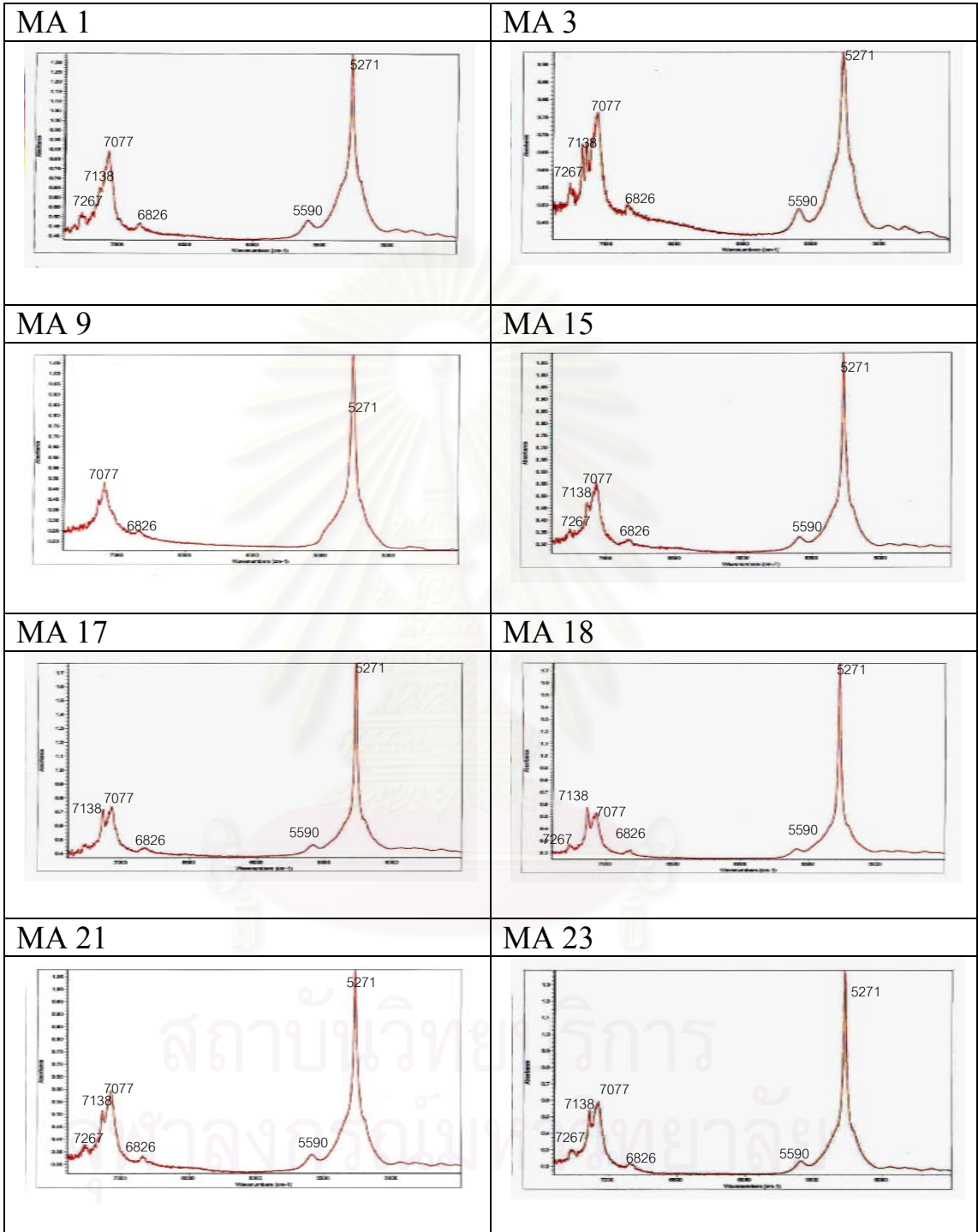


FTIR spectra of emerald samples from Santa Terezinha, Brazil: *Type II H<sub>2</sub>O dominant (IIa)*

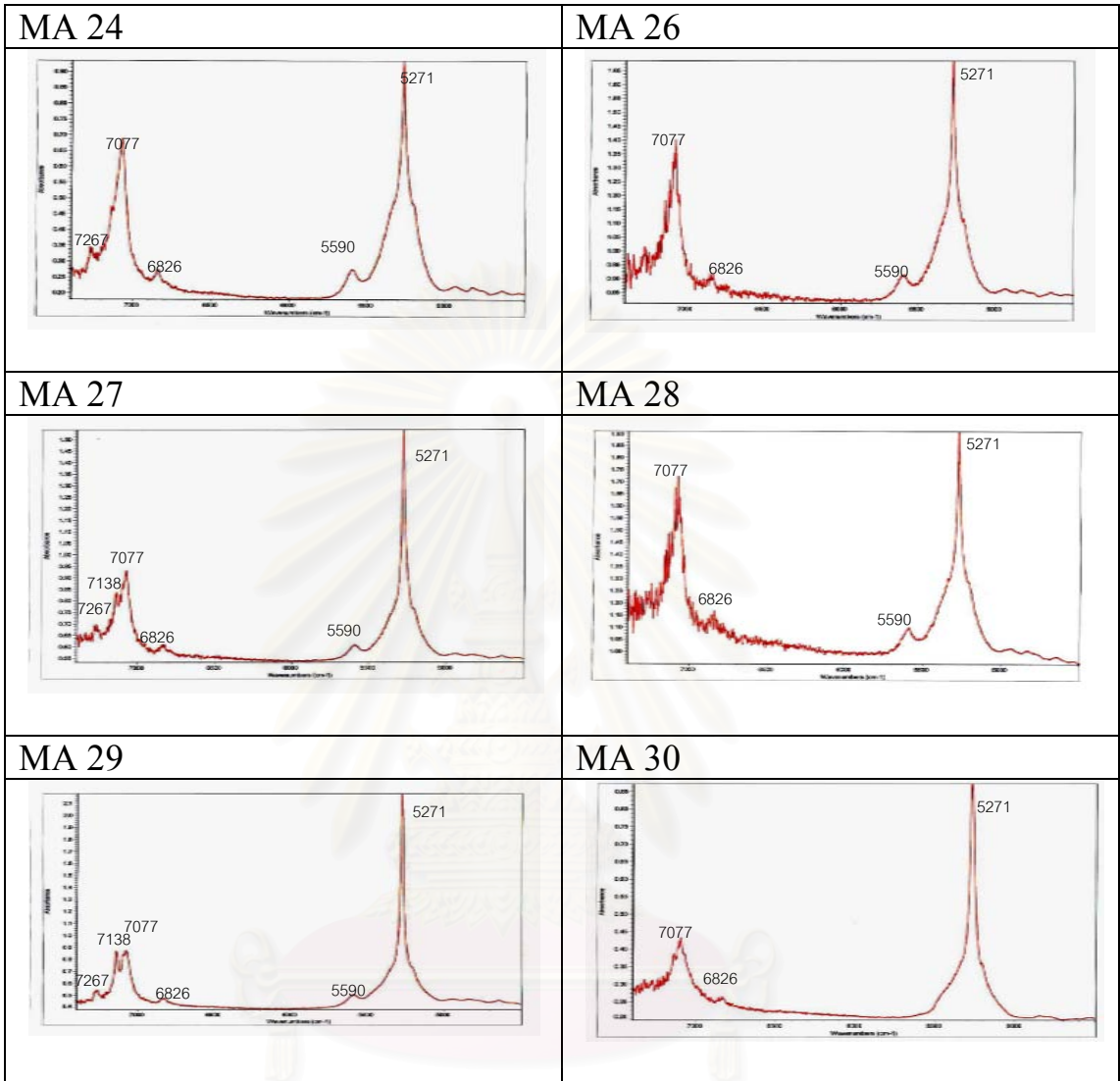




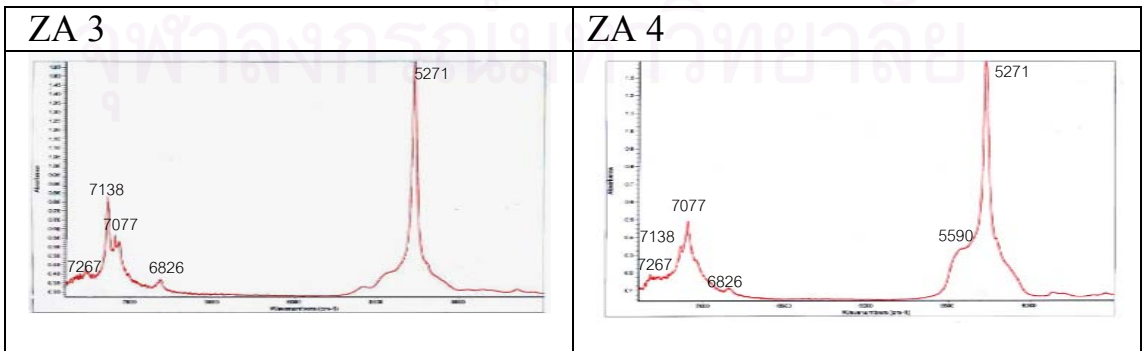
FTIR spectra of emerald samples from Madagascar: Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)



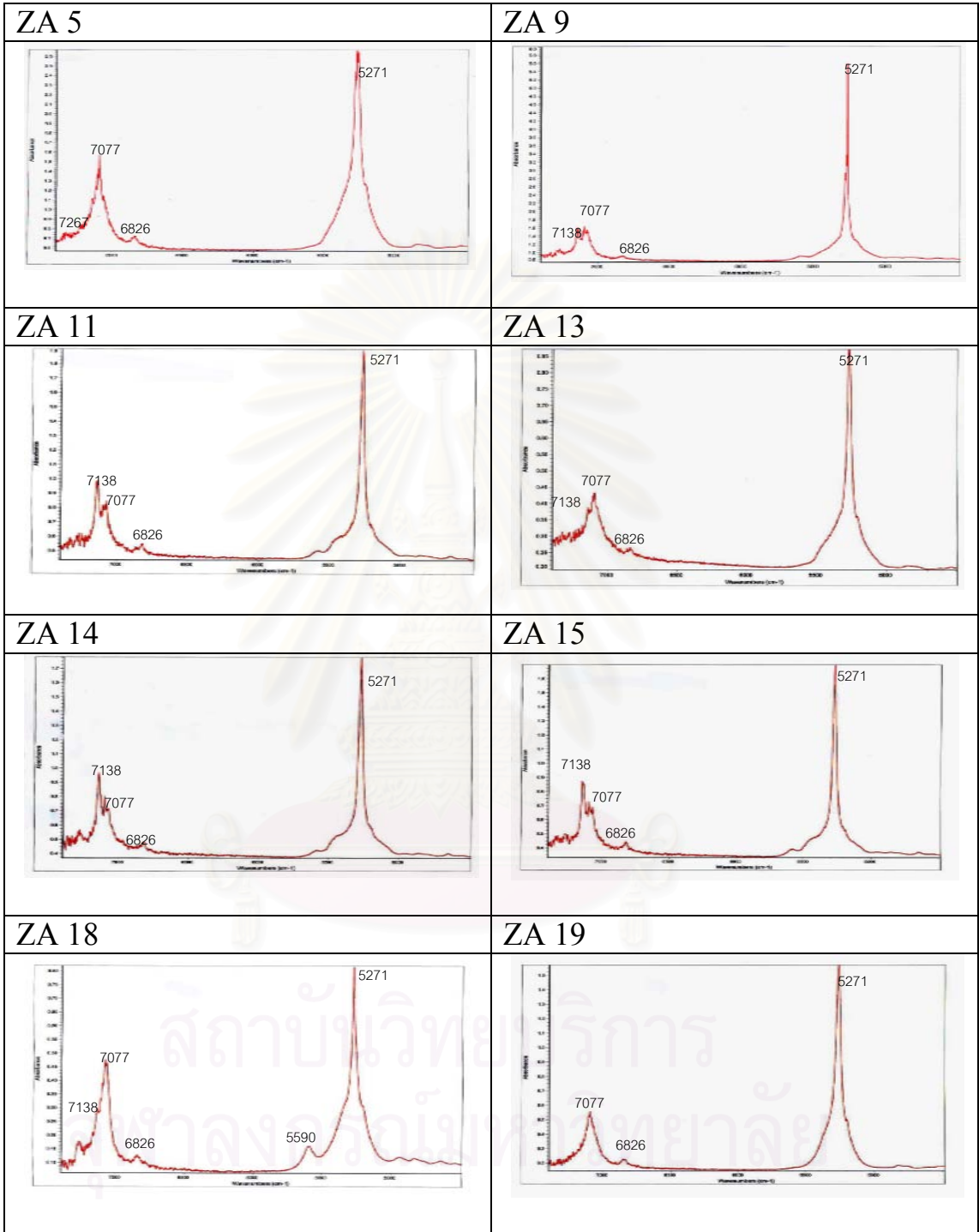
FTIR spectra of emerald samples from Madagascar: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)*, continue



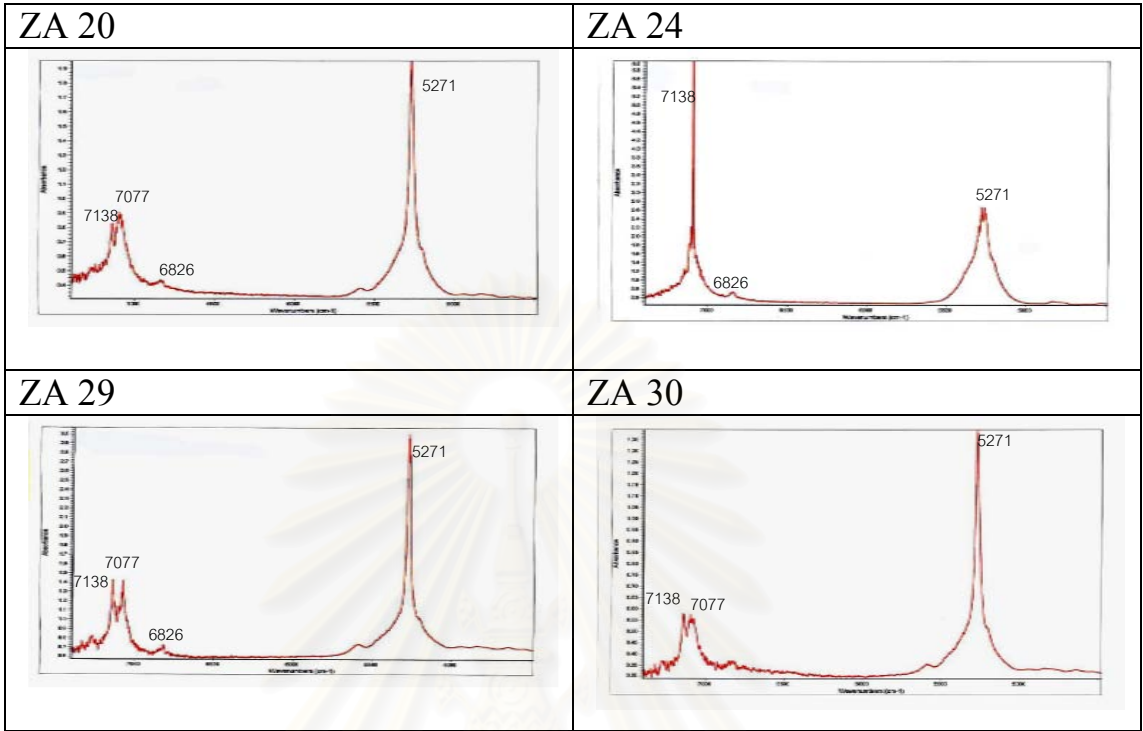
FTIR spectra of emerald samples from Zambia: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)*



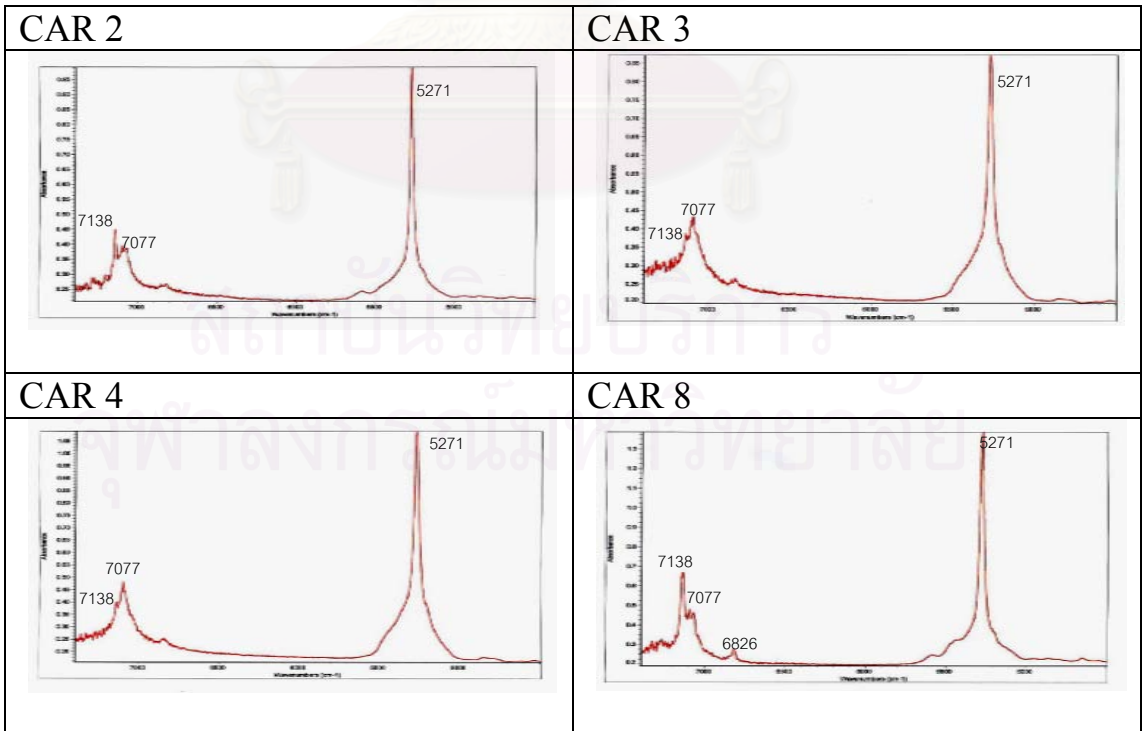
FTIR spectra of emerald samples from Zambia: Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb), continue



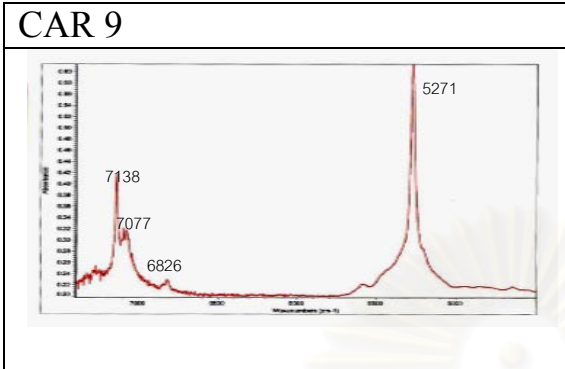
FTIR spectra of emerald samples from Zambia: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb), continue*



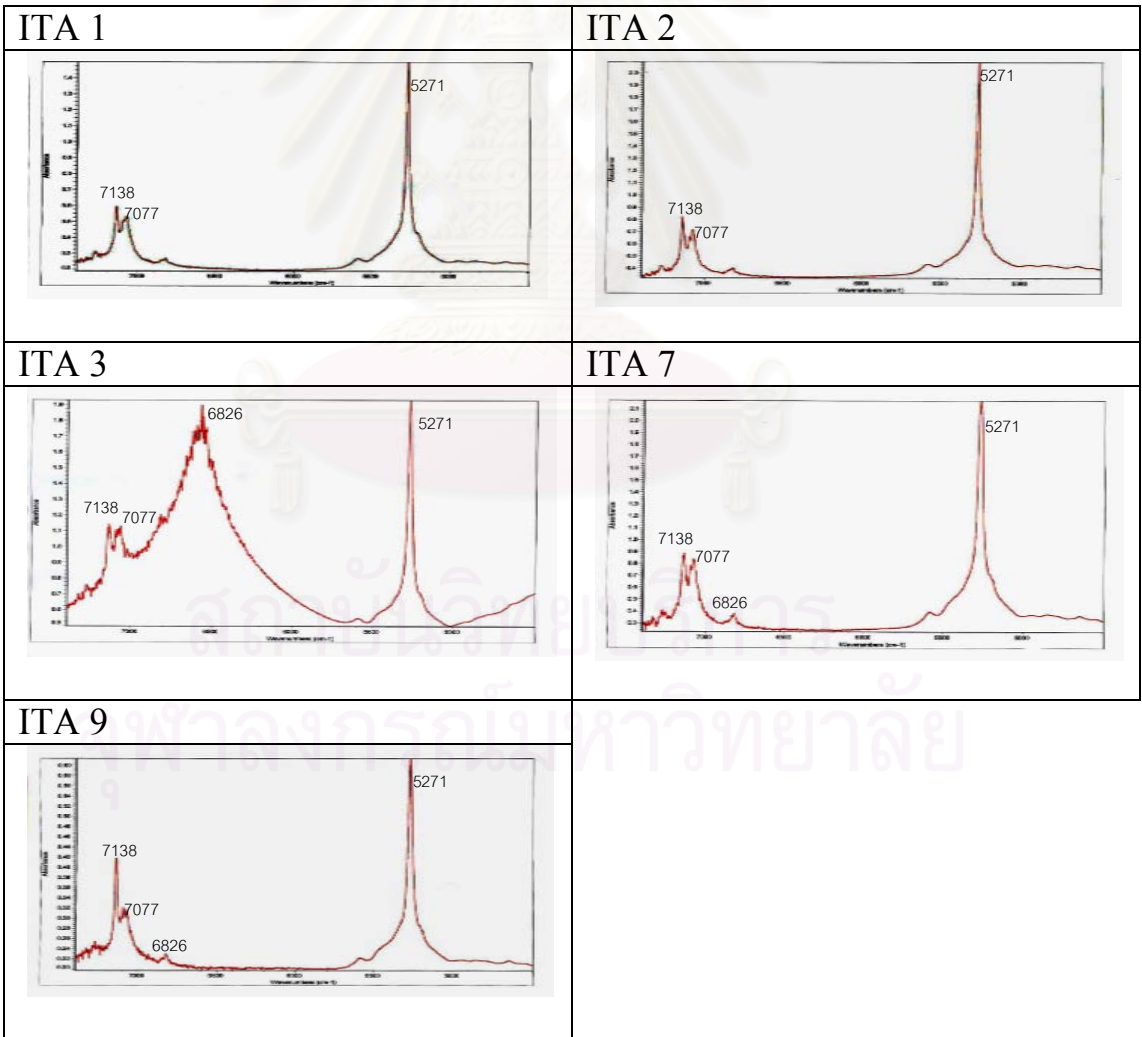
FTIR spectra of emerald samples from Carnaiba/Socoto, Brazil: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)*



FTIR spectra of emerald samples from Carnaiba/Socoto, Brazil: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)*,  
 continue



FTIR spectra of emerald samples from Itabira, Brazil: *Type I H<sub>2</sub>O + II H<sub>2</sub>O (IIb)*



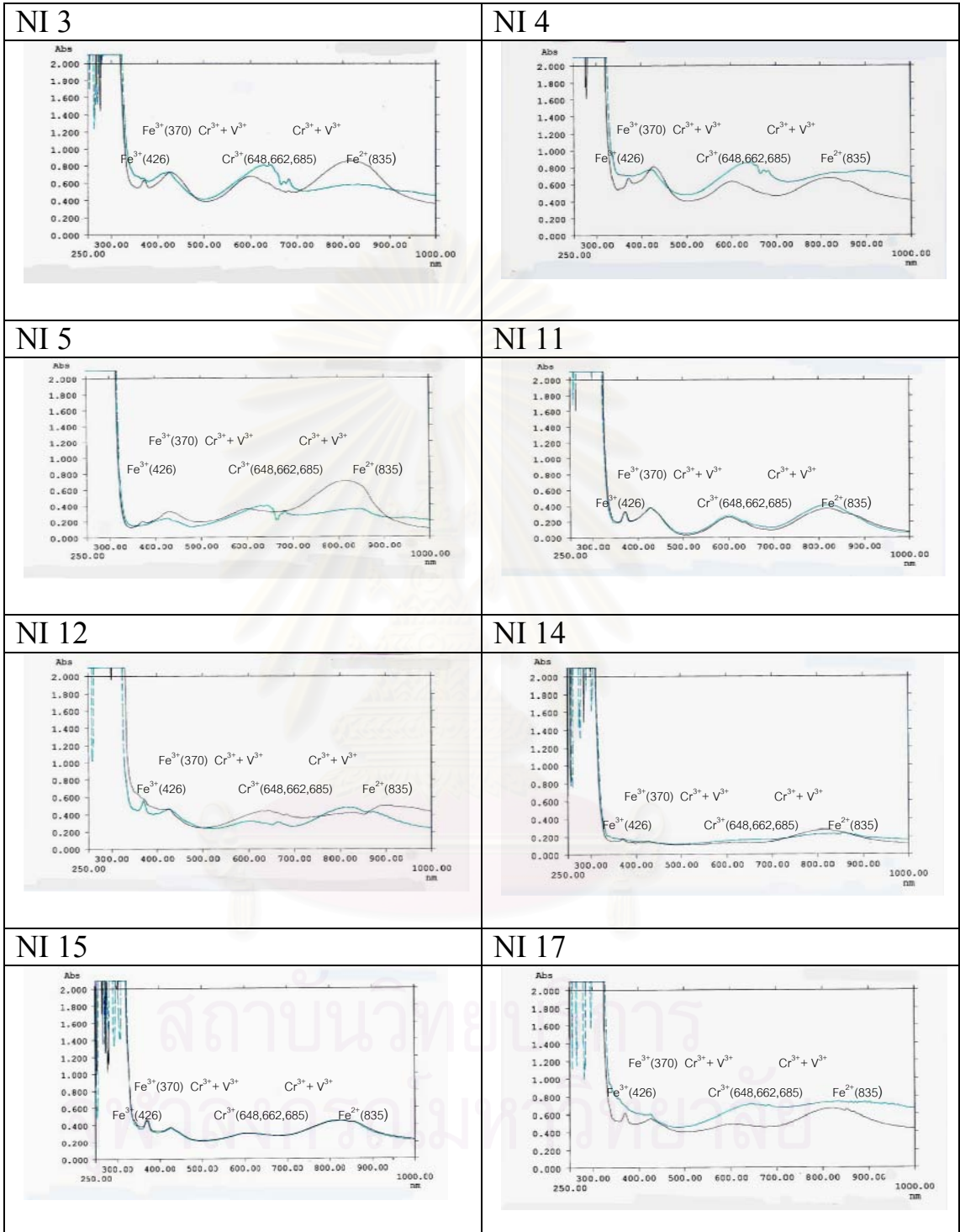


### APPENDIX III

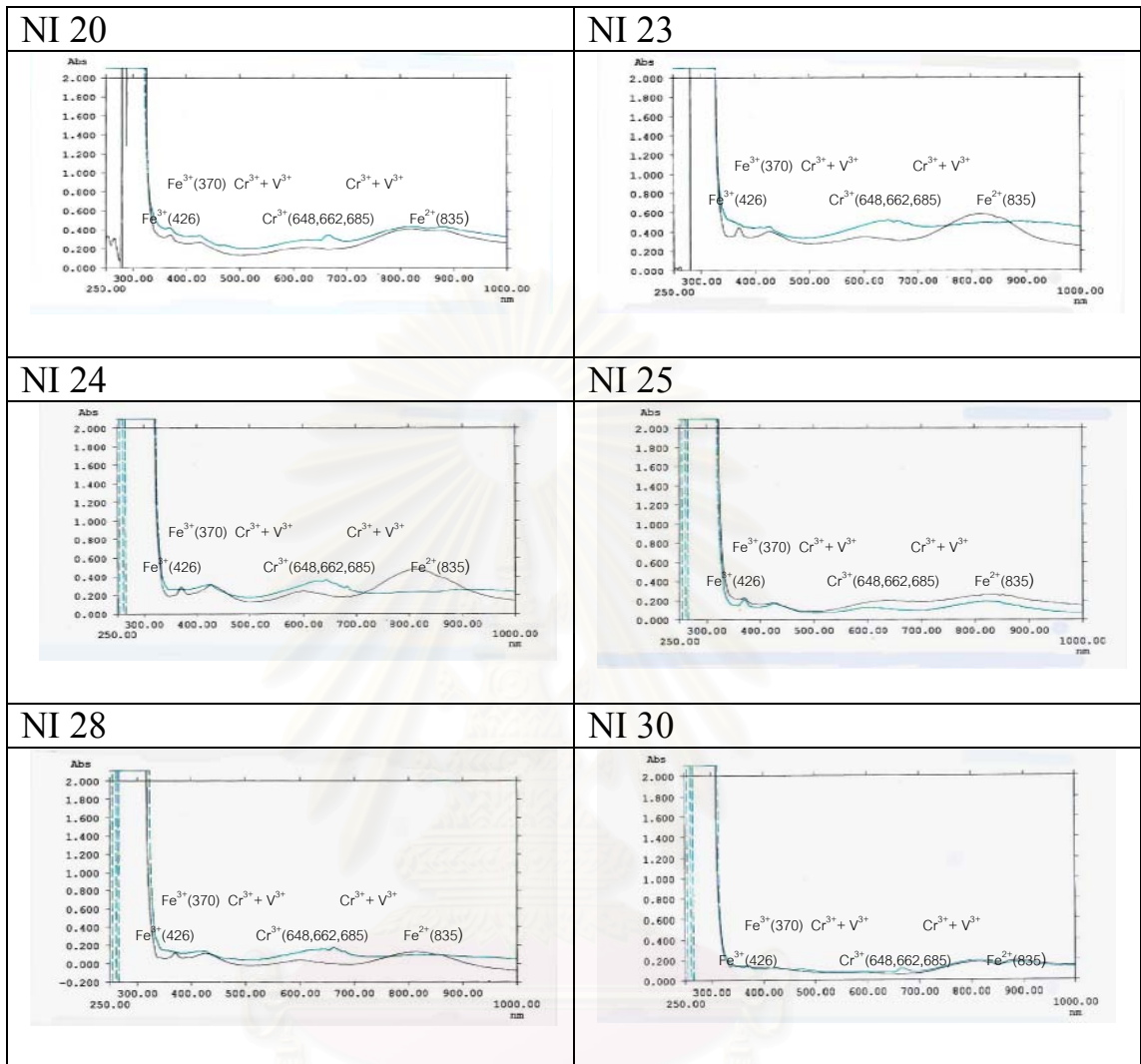
UV-Vis-NIR spectra of emerald samples from Nigeria, Colombia, Brazil, Madagascar and Zambia;

- 15 samples of Kaduna Plateau emeralds from Nigeria: NI3, NI4, NI5, NI11, NI12, NI14, NI15, NI17, NI10, NI20, NI23, NI24, NI25, NI28, NI30
- 12 samples of Cordillera Oriental emeralds from Colombia: CO1, CO2, CO4, CO6, CO8, CO11, CO13, CO18, CO20, CO22, CO26, CO27
- 4 samples of Santa Terezinha from Goais State, Brazil: SAN3, SAN7, SAN8, SAN10
- 14 samples of Manajary emeralds from Madagascar: MA1, MA3, MA9, MA15, MA17, MA18, MA21, MA23, MA24, MA26, MA27, MA28, MA30
- 14 samples of emeralds from Ndola Rural district: ZA3, ZA4, ZA5, ZA11, ZA12, ZA14, ZA15, ZA17, ZA20, ZA23, ZA24, ZA25, ZA28, ZA30
- 5 samples of Carnaiba/Socoto emeralds from Bahia State, Brazil: CAR2, CAR3, CAR4, CAR8, CAR9
- 5 samples of Itabira emeralds from Minas Gerias State, Brazil: ITA1, ITA2, ITA3, ITA7, ITA9

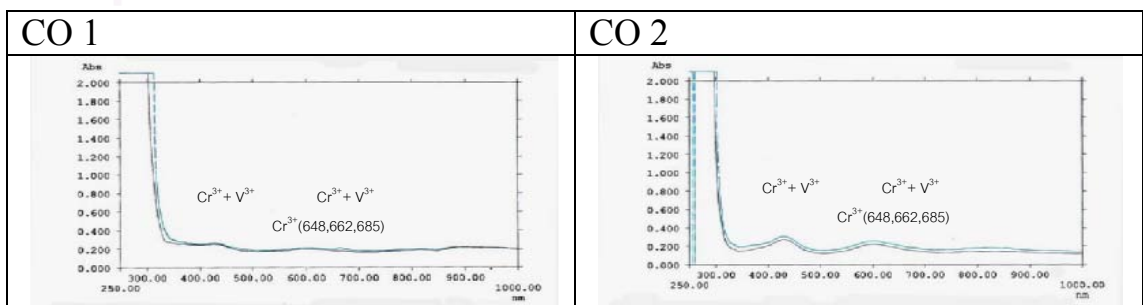
UV-Vis-NIR spectra of emerald samples from Nigeria (--- = O-ray, --- = E-ray)



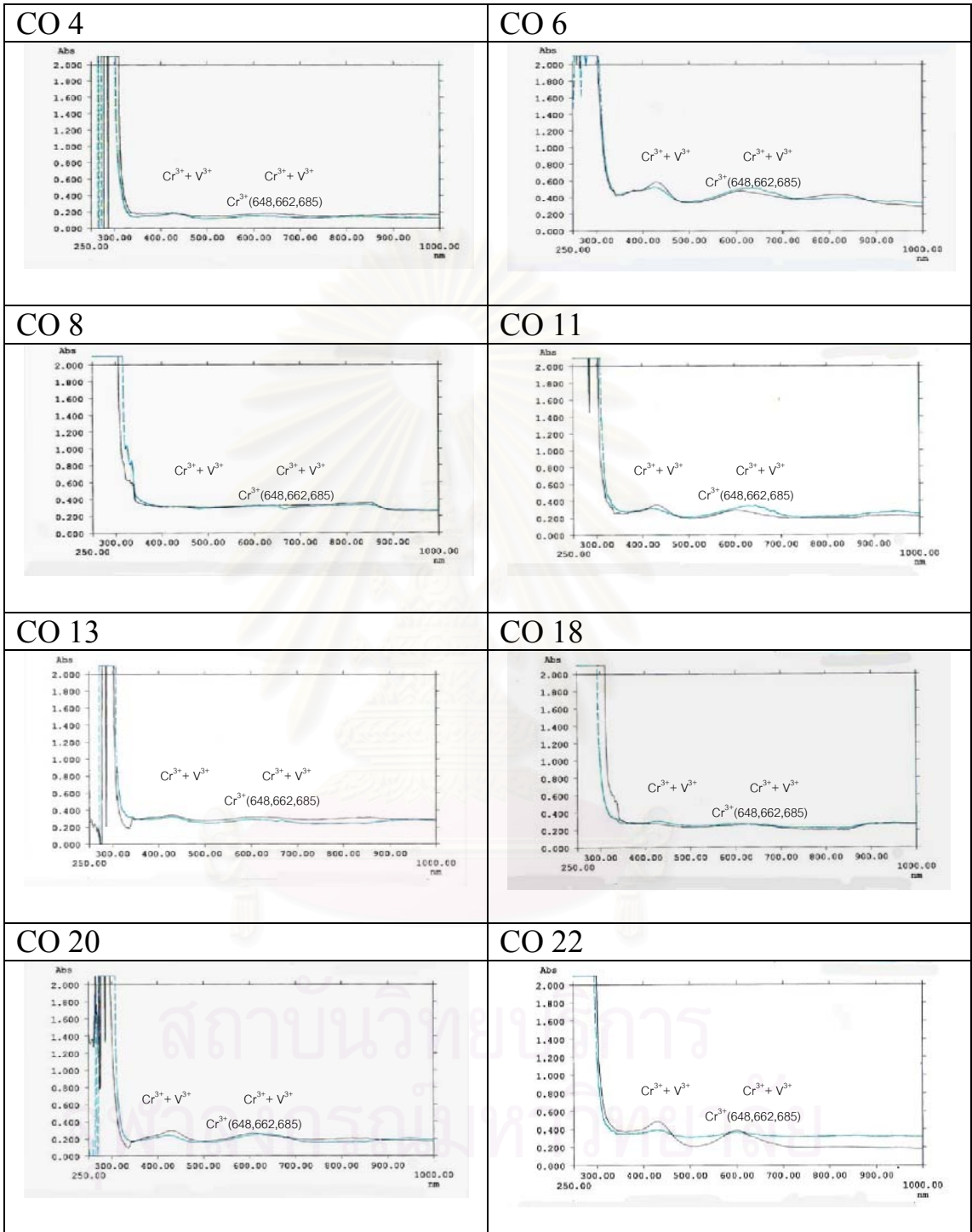
UV-Vis-NIR spectra of emerald samples from Nigeria (continue) (--- = O-ray, --- = E-ray)



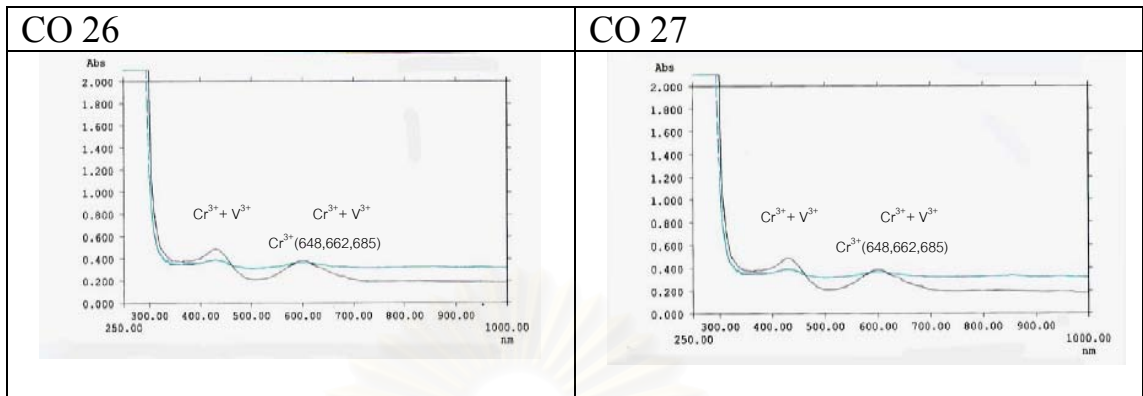
UV-Vis-NIR spectra of emerald samples from Colombia (--- = O-ray, --- = E-ray)



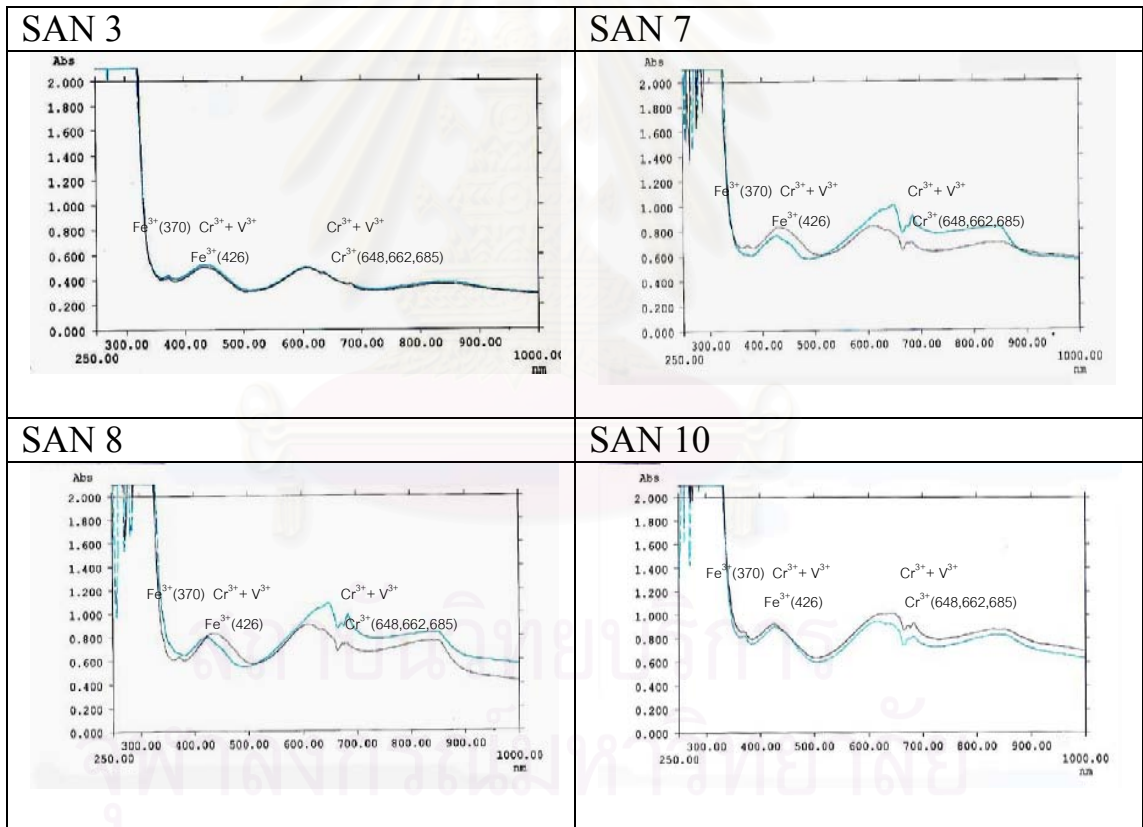
UV-Vis-NIR spectra of emerald samples from Colombia (continue) (--- = O-ray, --- = E-ray)



UV-Vis-NIR spectra of emerald samples from Colombia (continue) (--- = O-ray, --- = E-ray)

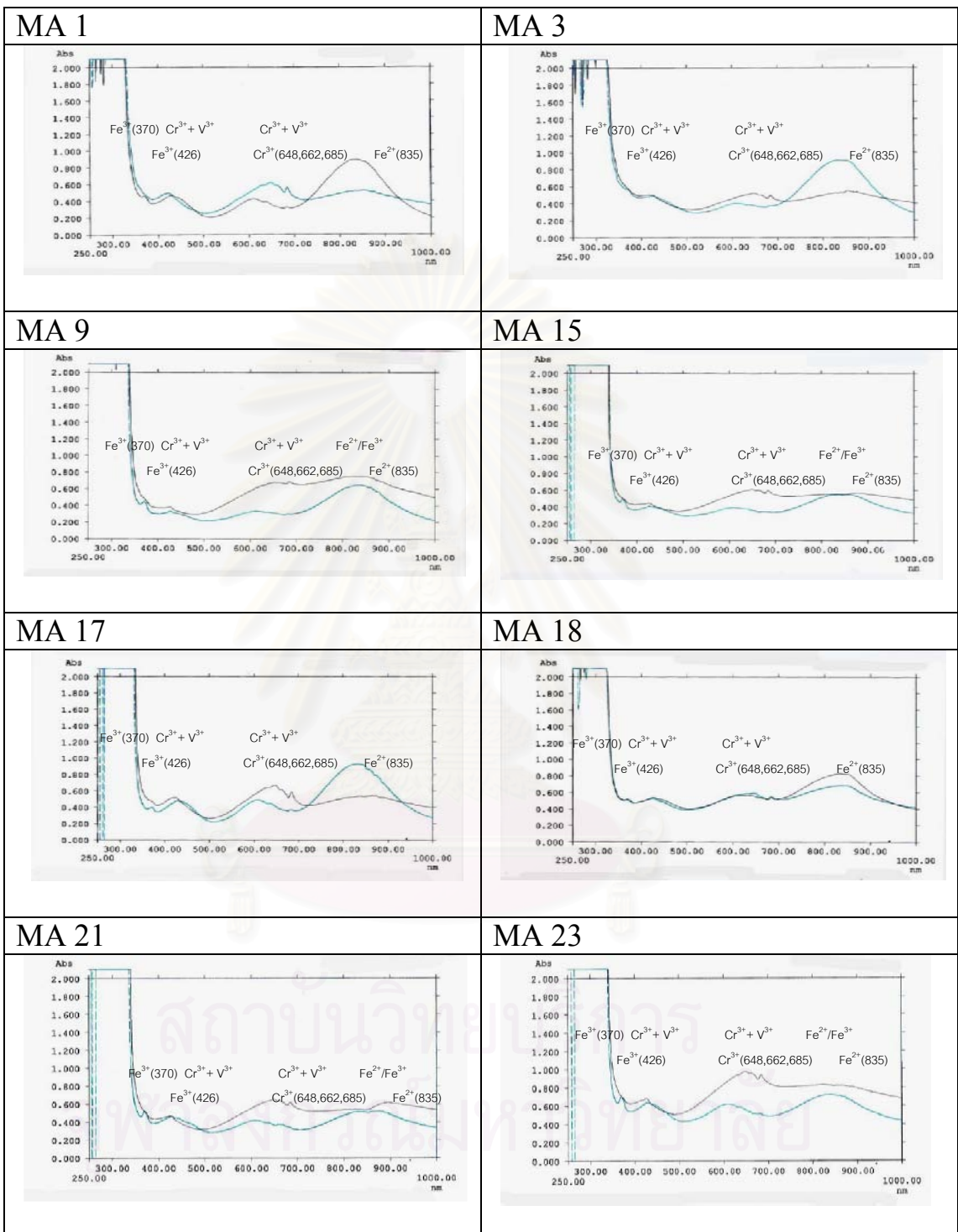


UV-Vis-NIR spectra of emerald samples from Santa Terezinha, Brazil (--- = O-ray, --- = E-ray)

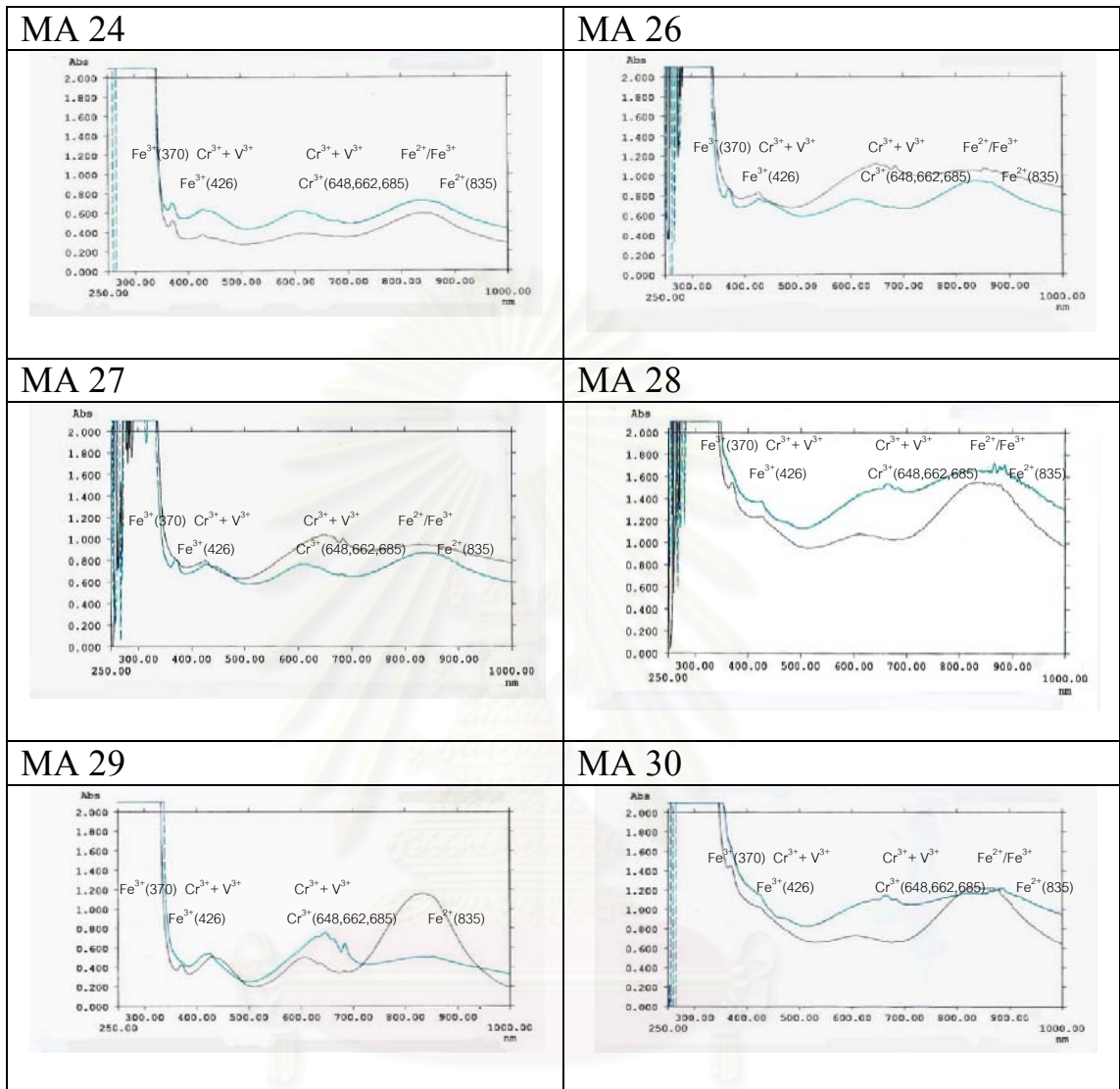




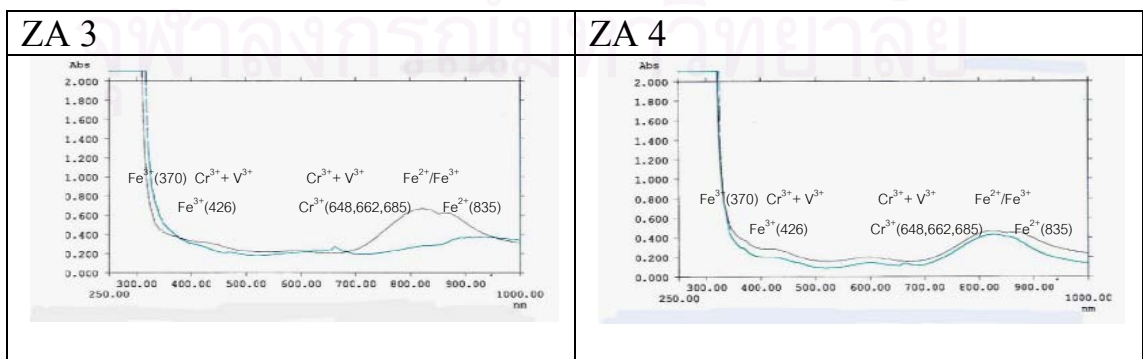
UV-Vis-NIR spectra of emerald samples from Madagascar (--- = O-ray, --- = E-ray)



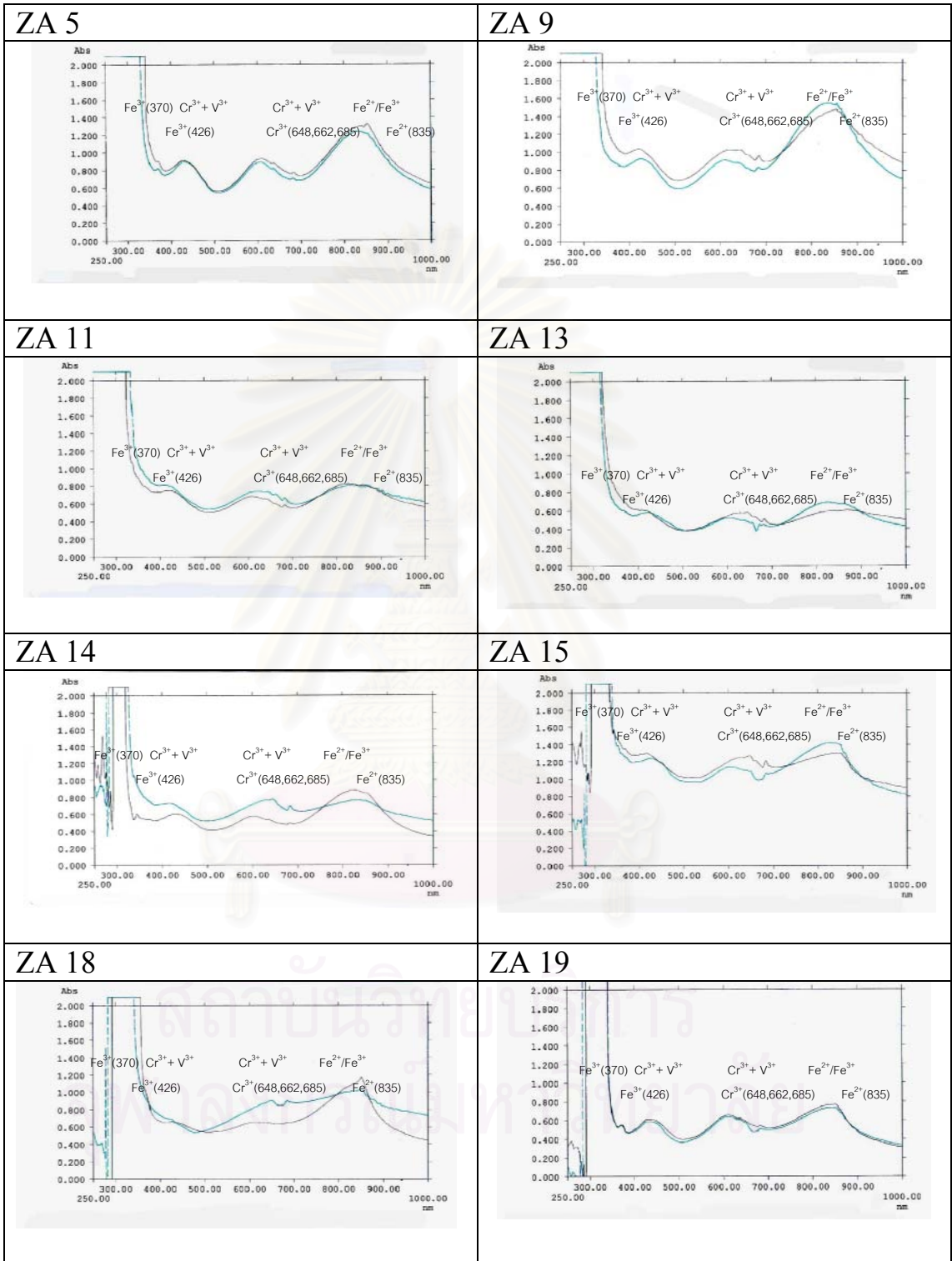
UV-Vis-NIR spectra of emerald samples from Madagascar (continue) (--- = O-ray, --- = E-ray)



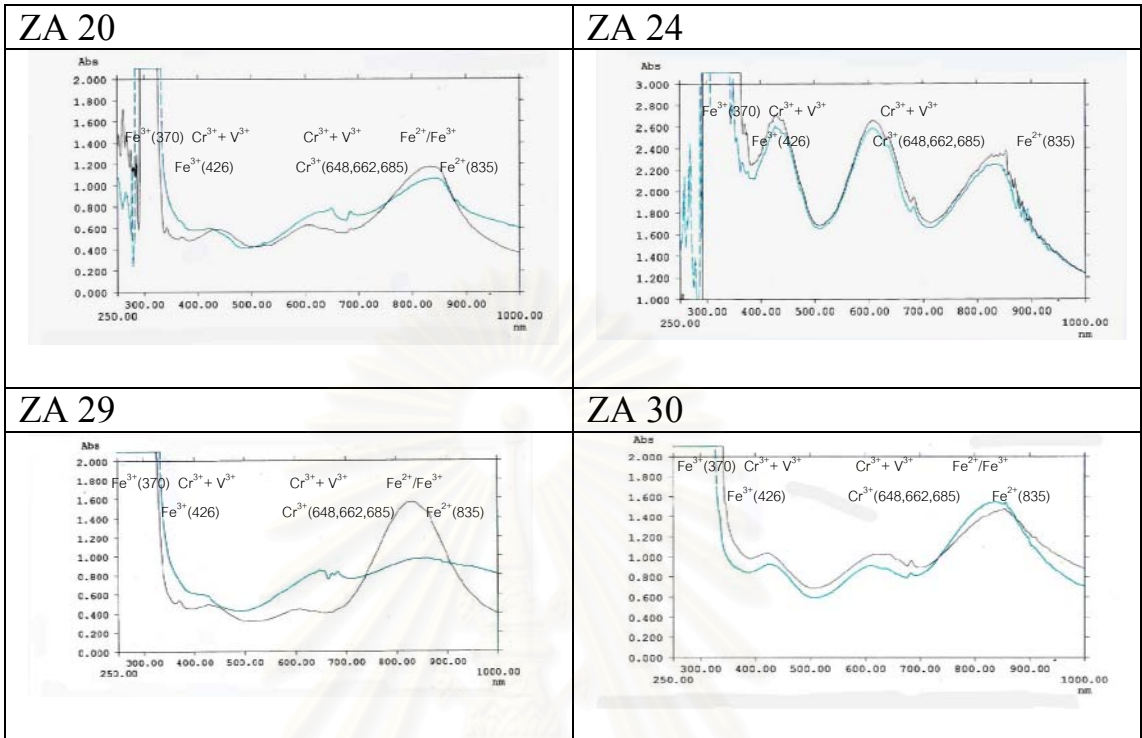
UV-Vis-NIR spectra of emerald samples from Zambia (--- = O-ray, --- = E-ray)



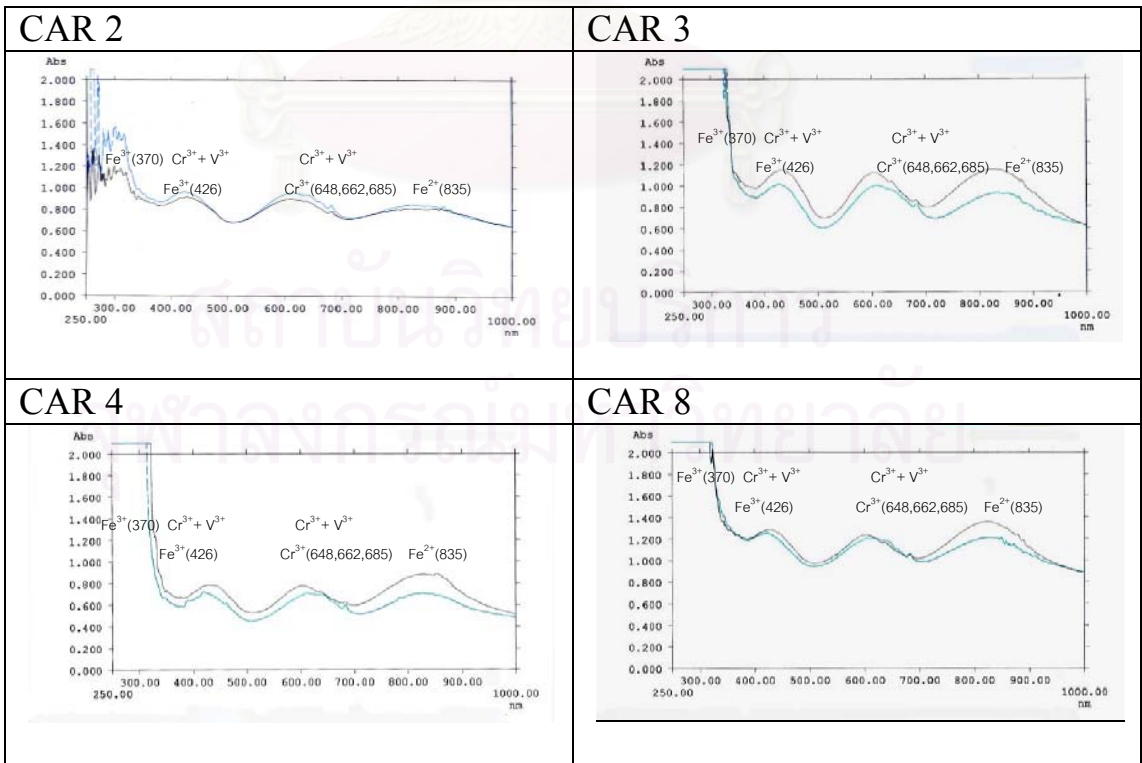
UV-Vis-NIR spectra of emerald samples from Zambia (continue) (--- = O-ray, --- = E-ray)



UV-Vis-NIR spectra of emerald samples from Zambia (continue) (--- = O-ray, --- = E-ray)

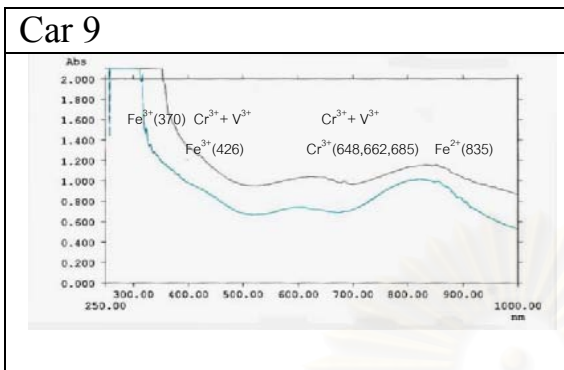


UV-Vis-NIR spectra of emerald samples from Carnaiba/Socoto, Brazil (--- = O-ray, --- = E-ray)

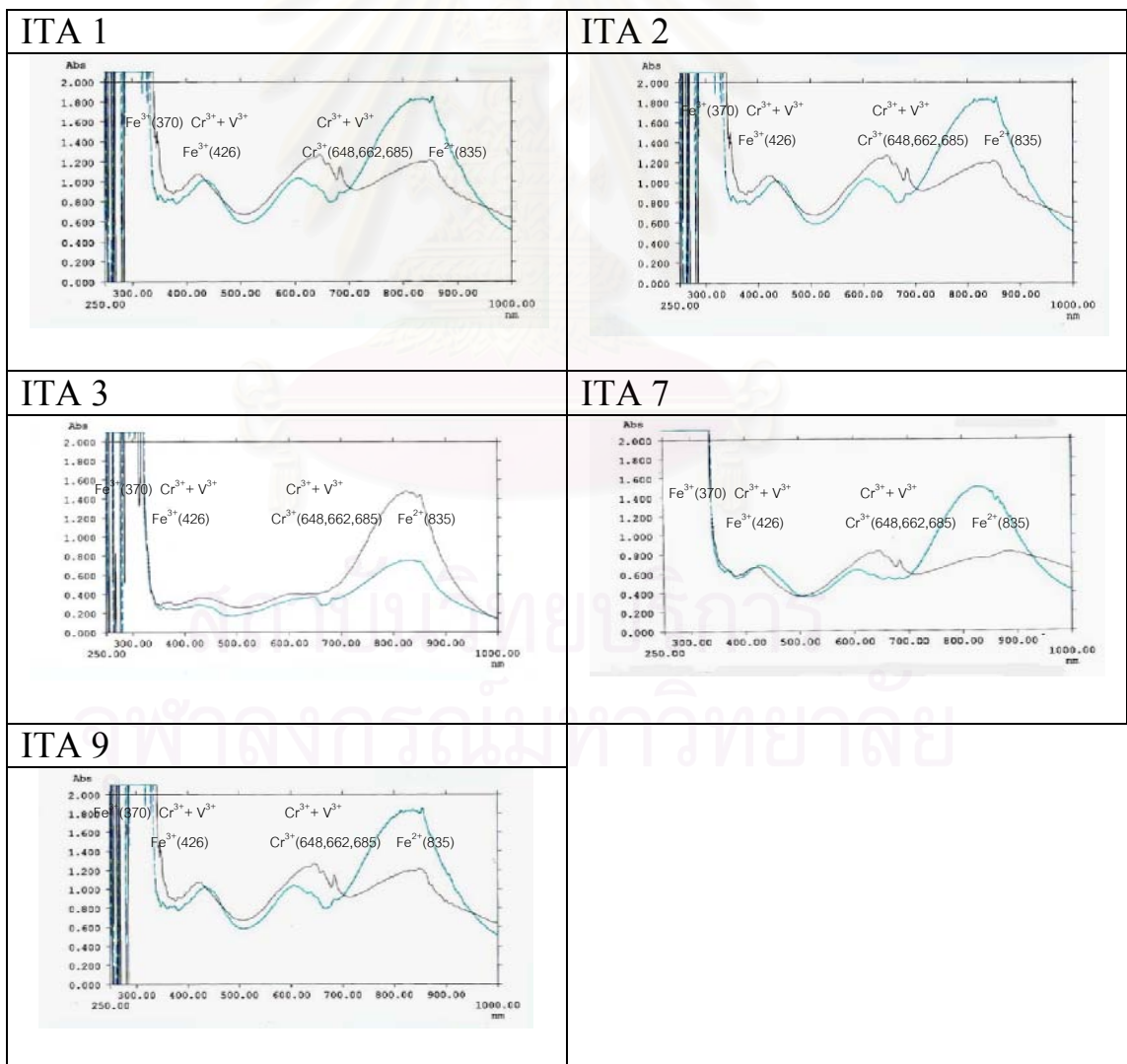




UV-Vis-NIR spectra of emerald samples from Carnaiba/Socoto, Brazil (continue) (--- = O-ray, --- = E-ray)



UV-Vis-NIR spectra of emerald samples from Itabira, Brazil (--- = O-ray, --- = E-ray)







## APPENDIX IV

Table IV.1 to IV.7: Chemical data show concentration (atom mole ppm) of 16 trace elements of emeralds from Nigeria (NI), Colombia (CO), Santa Terezinha (STA), Madagascar (MA), Zambia (ZA), Carnaiba/Socoto (CAR) and Itabira (ITA).

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## BIOGRAPHY

Miss Thitinthree Pavaro was born in February 23, 1975, in Bangkok. She graduated with a bachelor degree in Gemology from the Department of Materials Science, Faculty of Science, Srinakharinwirot University in 1997. At present, she works as a gemologist at The Gem and Jewelry Institute of Thailand, and also studies in a Master Program in Earth Science at the Department of Geology, Chulalongkorn University.



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