กำเนิดเซิงศิลาวิทยาของหินแปลกปลอมอุ้มพลอยทับทิม จากศิลาภูมิประเทศหินบะซอลต์ บริเวณเมืองซิมบาและเมืองอิมาลี ประเทศเคนยา



นายธวัชชัย เชื้อเหล่าวานิช

วิทยานิพนธ์นี้เป็นส่วนหนึ่งของการศึกษาตามหลักสูตรปริญญาวิทยาศาสตรดุษฎีบัณฑิต สาขาวิชาธรณีวิทยา ภาควิชาธรณีวิทยา คณะวิทยาศาสตร์ จุฬาลงกรณ์มหาวิทยาลัย ปีการศึกษา 2553 ลิขสิทธิ์ของจุฬาลงกรณ์มหาวิทยาลัย



PETROGENESIS OF RUBY-BEARING XENOLITH FROM SIMBA AND EMALI BASALTIC TERRANES, KENYA

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A Dissertation Submitted in Partial Fulfillment of the Requirements for the Degree of Doctor of Philosophy Program in Geology Department of Geology Faculty of Science Chulalongkorn University Academic Year 2010 Copyright of Chulalongkorn University

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งานวิจัยนี้มุ่งเน้นศึกษาการกำเนิดเชิงศิลาวิทยาของหินแปลกปลอมอุ้มพลอยทับทิมที่พบ มากในพื้นที่ศิลาภูมิประเทศหินบะขอลต์ขนาดเล็ก 2 บริเวณ คือ บริเวณกลุ่มเนินภูเขาไฟเขากู และ ้บริเวณกลุ่มเนินภูเขาไฟกูลัย ใกล้กับเมืองซิมบา ในภูมิภาคตะวันออกเฉียงใต้ของประเทศเคนยา ซึ่งศิลา ้ ภูมิประเทศทั้งสองนี้จัดเป็นส่วนหนึ่งของเขตหินภูเขาไฟไขอูลูตอนเหนือ พบว่าหินบะซอลต์บริเวณ นี้เป็นเนื้ออัลคาไลที่มีองค์ประกอบอยู่ในช่วง ฟอยไดต์ ถึง เทไฟรต์/บาซาไนต์ โดยมีลักษณะ องค์ประกอบทางเคมีที่บ่งชี้ถึงแหล่งต้นกำเนิดแมกมาที่มีแร่แอมฟิโบลและ/หรือโฟลโกไปต์เป็น ้ส่วนประกอบที่ผ่านการหลอมละลายบางส่วน จากผลของกระบวนการร่วมสัมพันธ์ระหว่าง mantle plume กับโครงสร้างที่ได้จากการคลายแรงที่เกิดขึ้นจากการขนกันของแผ่นเปลือกโลกที่ ก่อเกิดแนวเทือกเขาโมซัมบิกในอดีต(~600-800 ล้านปี) ทั้งนี้บะขอลต์ในบริเวณนี้ได้นำเอาหิน แปลกปลอมระดับลึกชนิดแกรนูไลต์ที่มีพลอยทับทิมผึงอยู่ภายในขึ้นมาสู่ผิวโลกพร้อมๆกับหิน แปลกปลอมอัลตร้าเมฟิกชนิดเพอริโดไทต์และไพรอกซิไนต์ ซึ่งหินแปลกปลอมนี้มีองค์ประกอบ ทางเคมีเทียบเคียงได้กับหินบะซอลต์ โดยมีเนื้อหินที่แสดงถึงลักษณะการแปรสภาพภายใต้แรง กดดันระดับสูงและการเกิดขอบปฏิกิริยารอบเม็ดแร่ที่ซับซ้อนประกอบร่วมกับลักษณะทางเคมีที่ ปรากฏ ที่บ่งชี้ถึงการแปรสภาพระดับแกรนูไลต์ของชั้นหินต้นกำเนิดที่น่าจะเทียบเคียงได้กับชั้นหิน อัคนีขนิดเมฟิกที่การสลับชั้นระหว่างหินเนื้อแกรโบกับหินเนื้ออนอร์โทไซต์ ภายใต้ช่วงสมดุล ภาวะ∼750-1500ºC และ 5-5-23 kb เทียบเท่ากับช่วงความลึกระดับเปลือกโลกชั้นล่าง (~20 km) ทะลูผ่านช่วงชั้นรอยต่อ Moho (~44 km) ลึกลงไปถึงในชั้นเนื้อโลกส่วนบน (~75 km) ซึ่ง คาบเกี่ยวกับช่วงความลึกที่แมกมาบะซอลต์ถือกำเนิด (~30-90 km) โดยมีภูเขาไฟปะทูและนำ พลอยจากที่ลึกขึ้นสู่ผิวโลกหลายครั้งในช่วงปลายไพลโอซีน ถึง ไพลสโตซีน (~ 2-0.8 ล้านปีที่ผ่านมา)

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TAWATCHAI CHUALAOWANICH: PETROGENESIS OF RUBY-BEARING XENOLITH FROM SIMBA AND EMALI BASALTIC TERRANES, KENYA. ADVISOR: ASST.PROF. CHAKKAPHAN SUTTHIRAT, Ph.D., CO-ADVISOR: ASSOC.PROF. VISUT PISUTTHA-ARNOND, Ph.D.,

Abundant ruby-bearing granulite xenoliths have been found in association with ultramafic xenoliths in the northern Chyulu Volcanic Province, within two small basaltic terranes of the Nguu Hills and the Ngulai Hills, located in the vicinity of Simba town, SE Kenya. The host basalts belong to an alkali affinity ranging in chemical composition from foidite to tephite/basanite. These chemical characteristics suggest that the magmas have been derived from partial melting of an amphibole- and/or phlogophite-bearing spinel lherzolite in the upper mantle source region that involves the interaction between pre-existing zones of lithospheric weak zones, caused by the evolution of the Mozambic Belt (~500-800 Ma), and slightly upwelling of mantle plume. The observed xenoliths, e.g. peridotite, pyroxenite and granulite, bear equivalent basaltic composition. However, complex reactions, as well as plastic deformation, combined with geochemical characteristics of these xenoliths may lead to progressive transformation of a layering mafic (gabbroic/anorthositic) cumulate protolith including corundum-bearing formation and other related rocks under a condition of granulite facies. Based on geothermobarometric constraints, these rocks appear to have undertaken high P-T ranges of ~750-1500°C and 5-23 kb equivalent to depths of the lower crust (~20 km) across the Moho (~44 km) down to the upper mantle (~75 km), corresponding to where the host magmas generated (~30-90 km). The corundum-related basalt eruptions had intermittently occurred during late Pliocene $(2.1\pm0.09 \text{ Ma})$ to middle Pleistocene (0.83±0.03Ma) times.

Department : _____ Geology_____ Student's Signature : ______ Field of Study : Geology Advisor's Signature : L.S.H.L. Academic Year : 2010 Co-Advisor's Signature : Vint Pruth MA

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- Figure 4.5. a) P-T constraints of foliated mafic granulite, *KNt01*, according to 5 different thermobarometers including (1) Cpx-Opx (Brey and Köhler, 1990), (2) Grt-Cpx (Ai, 1994) and (3) Grt-Opx (Aranovich and Berman, 1997) thermometers, and (4) Grt-Cpx-Pl-Qtz (Eckert *et al.*, 1991) and (5) Grt-Opx-Pl-Qtz (Lal, 1993) barometers. Star symbol marking the best-fit P-T condition within the constraint area. b) P-T constraints of banded mafic granulite, *Ng21*, according to 4 different thermobarometers including (1) Cpx-Opx (Brey and Köhler, 1990) and (2) Grt-Opx (Aranovich and Berman, 1997) thermometers, and (3) Grt-Cpx-Pl-Qtz (Eckert *et al.*, 1991) and (4) Grt-Opx-Pl-Qtz (Lal, 1993) barometers. Star symbol marking the best-fit P-T condition within the constraint area.
- Figure 4.6 P-T constraints of felsic granulite, *KMb01*, according to 4 different thermo-barometers including (1) Cpx-Opx (Brey and Köhler, 1990) and (2) Grt-Opx (Aranovich and Berman, 1997) thermometers, and (3) Grt-Cpx-Pl-Qtz (Eckert *et al.*, 1991) and (4) Grt-Opx-Pl-Qtz (Lal, 1993) barometers. Star symbol marking the best-fit P-T condition within the constraint area.
- Figure 4.7 Pseudosection diagrams of the periditite xenoliths represented by Spl lherzolite (*Ng07*: left) and the Spl wehrlite (*Ng18*: right).
- Figure 4.8 Pseudosection diagrams of the pyroxenite xenoliths represented by Spl-free websterite (*Ng34*: left) compared to the pyroxene-rich band of the crn-bearing granulite (*KNt02_2*: right).
- Figure 4.9 Pseudosection diagrams of the granulite xenoliths represented by foliated, Crn-beaing mafic granulite (*Ng21*: left) compared to the crn-barren felsic granulite (*KMb01*: right).
- Figure 5.1 Plots of K/Th* (=K2O*10,000/Th) versus Th (ppm) for the basaltic rocks from the Nguu Hills, Ngulai Hills, and Kiboko (Chyulu) lavas in comparison to the published data ranges from the Chyulu Hills and the Western Rift of Rungwe and Muhavura volcanic provinces. Dotted and dashed borders are from Furman (2007).

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- Figure 5.2 Covariation plot of ¹⁴³Nd/¹⁴⁴Nd vs. ⁸⁷Sr/⁸⁶Sr displaying a broad negative trend from less radiogenic values in northern section of the Eastern Kenya Rift (NRK) towards higher radiogenic compositions observed in the Western Kenya Rift of the Kivu, Rungwe, Karisimbi and Muhavura volcanic regions. The values from the Chyulu Hills are within the lower part of the NKR approaching towards the field of HIMU. Point data for Chyulu Hills and Nguu Hills lavas are from Späth *et al.* (2001).
- Figure 5.3 Covariation plot of ²⁰⁸Pb/²⁰⁴Pb vs. ²⁰⁶Pb/²⁰⁴Pb displaying a broad positive trend for northern Kenya Rift lavas (NKR) and southern Kenya Rift, both off-rift (Chyulu Hills) and within-rift (SKR), towards HIMU. Point data for Chyulu Hills and Nguu Hills lavas are from Späth *et al.* (2001).
- Figure 5.4 Showing Chondrite-normalized REE patterns for a) the cogenetic lavas from the Nguu Hills, Ngulai Hills, and Kiboka areas. b) equilibrium batch melts of an amphibole-bearing spinel lherzolite source. c) equilibrium batch melts of an amphibole-garnet-bearing spinel lherzolite source. d) fractional melts of an amphibole-bearing spinel lherzolite source. Normalizing values from Sun and McDonough (1989).
- Figure 5.5 P-T constraint diagrams of the xenoliths compared to the host basalts.
- Figure 5.6 The Ti-V discrimination diagram for the xenoliths and their host basalts from the Nguu Hills and Ngulai Hills areas. IAT is standed for island arc tholeite and BAB is back arc basin. Lines represent Ti/V ratio.
- Figure 5.7 Cartoon summarizing the major features of deep structure underneath the study area and the Chyulu Volcanic Province (modified after Novak et al., 1997). Filled numbered ellipses represent ultramafic layers where 1 is peridotite and 2 is pyroxenite (plagioclase-free websterite). Filled numbered rectangles represent granulite layers where 3.1 is corundumbearing mafic granulite and 3.2 is corundum-free felsic granulite. Filled triangles are basalt bodies.

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